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Seasonal simulation of water, salinity and nitrate dynamics under drip irrigated mandarin (*Citrus reticulata*) and assessing management options for drainage and nitrate leaching

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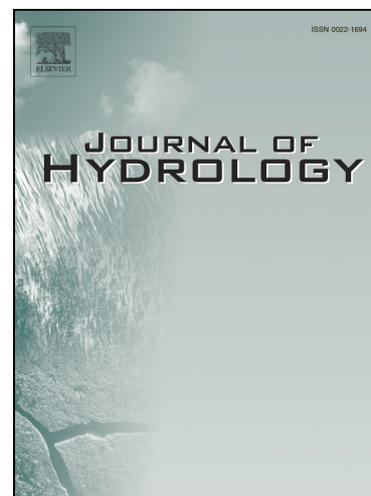
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1 **Seasonal simulation of water, salinity and nitrate dynamics under drip**  
2 **irrigated mandarin (*Citrus reticulata*) and assessing management options**  
3 **for drainage and nitrate leaching**

4  
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15  
16 **Summary**

17 Estimation of all water fluxes temporally and spatially within and out of the crop root zone,  
18 and evaluation of issues like salinity and nutrient leaching, are necessary to fully appraise the  
19 efficiency of irrigation systems. Simulation models can be used to investigate these issues  
20 over several seasons when the cost of long term monitoring is prohibitive. Model results can  
21 be used to advise growers if improvements are required to various aspects of irrigation  
22 system operations. In this study, HYDRUS-2D was used to evaluate data measured during  
23 one season in a young mandarin (*Citrus reticulata*) orchard, irrigated with an intensive  
24 surface drip fertigation system. Water contents, salinities, and nitrate concentrations  
25 measured weekly in the field were compared with model predictions.

26 The temporal mean absolute error (*MAE*) values between weekly measured and simulated  
27 water contents ranged from 0.01 to 0.04 cm<sup>3</sup>cm<sup>-3</sup>. However, modelling error (*MAE*) was  
28 slightly larger at 10 cm depth (0.04 cm<sup>3</sup>cm<sup>-3</sup>), as compared to greater depths (0.02-0.03  
29 cm<sup>3</sup>cm<sup>-3</sup>). Similarly, the errors were larger in the surface soil layer (25 cm depth) for nitrate-  
30 nitrogen, NO<sub>3</sub><sup>-</sup>-N (1.52 mmol<sub>(c)</sub> L<sup>-1</sup>), as compared to greater depths. The spatial and temporal  
31 soil solution salinity (*EC<sub>sw</sub>*) and NO<sub>3</sub><sup>-</sup>-N data showed accumulation of salts and nitrate within  
32 the soil up until day 150 of the simulation (December, 2006), followed by leaching due to  
33 high precipitation and over irrigation at later times. Only 49% of applied water was used by  
34 the mandarin trees, while 33.5% was leached. On the other hand, the simulation revealed that  
35 a significant amount of applied nitrogen (85%) was taken up by the mandarin trees, and the  
36 remaining 15% was leached. The results indicate that the irrigation and fertigation schedule  
37 needs modifying as there was overwatering from December onwards.

38 Different permutations and combinations of irrigation and fertigation scheduling were  
39 evaluated to optimise the water and nitrogen uptake and to reduce their leaching out of the  
40 crop root zone. Slightly higher nitrogen uptake (1.73 kg ha<sup>-1</sup>) was recorded when fertigation  
41 was applied second to last hour in an irrigation event, as compared to applying it earlier  
42 during an irrigation event. Similarly, a 20% reduction in irrigation and N application  
43 produced a pronounced reduction in drainage (28%) and N leaching (46.4%), but it also  
44 decreased plant N uptake by 15.8% and water uptake by 4.8%, and increased salinity by  
45 25.8%, as compared to the normal practice. This management would adversely impact the  
46 sustainability of this expensive irrigation system. However, reducing only irrigation by 30%  
47 during the 2<sup>nd</sup> half of the crop season (January to August) reduced drainage and N leaching  
48 by 37.2 and 50.5%, respectively, and increased N uptake by 6.9%. Such management of  
49 irrigation would be quite promising for the sustainability of the entire system. It is concluded  
50 that judicious manipulations of irrigation and fertilizer applications can be helpful in

51 designing drip irrigation schedules for perennial horticultural crops to achieve improved  
52 efficiency of irrigation and fertigation applications and reduced contamination of receiving  
53 water bodies.

54

55 **Keywords:** Modelling, HYDRUS, mandarin, fertigation, nitrate leaching, soil salinity

56

## 57 **Introduction**

58

59 Micro-irrigation has become the optimal standard for irrigation and fertigation of  
60 horticultural crops in Australia, due to increased water scarcity and higher costs of fertilizers  
61 over the last decade. Intensive fertigation schedules have been developed to increase yield  
62 and quality of many permanent horticultural crops, including mandarin. This combines drip  
63 irrigation and fertigation to deliver water and nutrients directly to the roots of the crop, with  
64 the aim of synchronizing the applications with crop demand (Assouline, 2002; Gärdenäs et  
65 al., 2005) and maintaining the desired concentration and distribution of ions and water in the  
66 soil (Bar-Yosef, 1999). The overall aim of these interventions is to develop an irrigation and  
67 nutrient management program that increases yield and fruit quality, while reducing leaching.  
68 The fundamental principle of drip fertigation is to apply water and nutrients regularly to a  
69 small volume of soil at a low application rate and at a high frequency to closely meet crop  
70 demand (Falivene et al., 2005). However, the potential for movement of water and mineral  
71 nutrients, especially nitrogen (as nitrate), below the root zone and into the ground- and then  
72 surface-waters using these approaches is still high. This is due to a number of factors: amount  
73 and intensity of precipitation, the large amounts of water and nutrients being applied, the  
74 limited capacity of roots to take up these nutrients, and to the ability of irrigators to manage  
75 drainage and hence leaching.

76 Citrus is one of the important horticultural crops being grown under intensive fertigation  
77 systems in Australia. The vast majority of citrus plantings are oranges (73%), with the rest  
78 split between mandarins (20%), lemons and limes (5%), and grapefruit (2%) (Horticulture  
79 Australia Limited, 2008). About 75% of the Australian citrus industry is located in the  
80 Murray-Darling Basin, utilising the lighter-textured free-draining soils adjacent to the  
81 Murray, Darling and Murrumbidgee rivers, and thus potential off-site effects of poorly  
82 managed fertigation may have wider implications.

83 Irrigated horticulture has, in general, been identified as the major source of nitrogen in  
84 drainage waters in the Murray Darling Basin (Harrison, 1994). A significantly high nitrate  
85 level has been reported in drainage water ( $60 \text{ mg L}^{-1}$ ) and soil solution ( $100 \text{ mg L}^{-1}$ ) under  
86 grapevines (Correll et al., 2010) in the Murray Darling Basin. These values are significantly  
87 higher than the Australian environmental trigger value ( $0.5 \text{ mg L}^{-1}$  for lowland rivers) for  
88 nitrate (ANZECC and ARMCANZ, 2000). Leaching of nitrates from soils under perennial  
89 horticulture may pose a potential threat to groundwater.

90 The main sources of nitrate in mandarin production are mineral fertilizers. Nitrate is  
91 removed from the soil by plant uptake or through decomposition by micro-organisms in the  
92 process of denitrification. In well-aerated soils typical of this region, denitrification is often  
93 negligible because of a lack of favourable conditions (Alva et al., 2006). Nitrate, being an  
94 anion, moves freely in these mineral soils, and hence has the potential to leach into  
95 groundwater and waterways if fertigation is not well scheduled (Paramasivam et al 2002;  
96 Gärdenäs et al., 2005; White, 2006). Several researchers have reported substantial leaching (6  
97 to 45%) of applied N in citrus cultivation under field conditions (Wang and Alva, 1996;  
98 Paramasivam and Alva, 1997; Paramasivam et al., 2002; Sluggett, 2010). Syvertsen and Jifon  
99 (2001) found that N leaching was higher under weekly fertigated orange trees than under  
100 daily or monthly fertigated trees. Syvertsen and Sax (1999) reported that increasing the

101 number of fertigation events could significantly reduce N leaching. However, they observed  
102 38 to 52% leaching of N from fertilizer, and the nitrogen use efficiency ranging between 25  
103 and 44% in Hamlin orange trees. Other researchers (Clothier et al., 1988; Li and Liu, 2011)  
104 have reported that nitrate accumulates toward the boundary of the wetted volume for most  
105 combinations of drip emitter discharge, input concentrations, and volumes applied. These  
106 studies suggest that there is a need for efficient tools, capable of describing and quantifying  
107 nitrate leaching, as well as nitrate uptake by crops, which in turn would help in designing and  
108 managing drip irrigation systems and achieving a high N fertilizer use efficiency, thereby  
109 limiting the export of this nutrient as a pollutant to downstream water systems.

110 In addition to nitrate leaching, salinity is also an important factor influencing the  
111 sustainability of the citrus production worldwide, as citrus species are relatively salt sensitive.  
112 The reported value of the average threshold electrical conductivity of saturation extract ( $EC_e$ )  
113 and slope for oranges (*Citrus sinensis*) are  $1.7 \text{ dS m}^{-1}$  and 16%, respectively (Maas and  
114 Hoffmann, 1977). Salt damage is usually manifested as leaf burn and defoliation, and is  
115 associated with accumulation of toxic levels of  $\text{Na}^+$  and/or  $\text{Cl}^-$  in leaf cells. Under drip  
116 irrigation there are many factors influencing the distribution of soil water and salts, and hence  
117 the water use efficiency (WUE), such as water quality, dripper discharge rate (Liu et al.,  
118 2012), irrigation water depth (Hanson et al., 2006), and irrigation frequency (El-Hendawy et  
119 al., 2008).

120 Simulation models have been valuable research tools in studies involving complex and  
121 interactive processes of water flow and solute transport through the soil profile, as well as the  
122 effects of management practices on crop yields and the environment (Pang and Letey, 1998;  
123 Li et al., 2003). HYDRUS-2D (Šimůnek et al., 2011) has been used extensively in evaluating  
124 the effects of soil hydraulic properties, soil layering, dripper discharge rates, irrigation  
125 frequencies, water quality, and timing of nutrient applications on wetting patterns and solute

126 distribution (e.g., Cote et al., 2003; Lazarovitch et al., 2005; Gärdenäs et al., 2005; Hanson et  
127 al., 2006; Ajdary et al., 2007; Phogat et al., 2009; Šimůnek and Hopmans, 2009; Li and Liu,  
128 2011; Phogat et al., 2012ab, 2013ab; Ramos et al., 2011, 2012). Although these studies  
129 demonstrate well the importance of numerical modelling in the design and management of  
130 irrigation and fertigation systems for various crops, most studies involving salinity and nitrate  
131 leaching are based on either an analysis of hypothetical scenarios, or are carried out for  
132 annual crops. Hence, there is a need to carry out modelling studies for perennial horticultural  
133 crops such as mandarin, using experimental results from field studies involving modern  
134 irrigation systems such as drip.

135 The objectives of the present investigation were to evaluate water, salt ( $EC_{sw}$ ), and nitrate  
136 ( $NO_3^-$ -N) movement in soil below young mandarin tree using HYDRUS-2D, and to evaluate  
137 various irrigation and fertigation strategies for controlling deep drainage and nitrate leaching,  
138 whilst maintaining soil salinity below the threshold for mandarin. This approach will help us  
139 understand the best irrigation and fertigation management practices to be adopted in future  
140 practical applications, with the goal to increase root water and nutrient uptake.

141

## 142 **2. Materials and Methods**

143

### 144 *2.1. Field experiment*

145

146 The field experiment was conducted at the Dareton Agricultural and Advisory Station  
147 (34.10°S and 142.04°E), located in the Coomealla Irrigation Area, 3 km from Dareton and 10  
148 km from Wentworth in New South Wales (NSW). The research station forms part of the  
149 Sunraysia fruit growing district of NSW and Victoria located in the Murray Darling Basin.

150 An experimental site with an intensive fertigation system, consisting of various mandarin  
151 (*Citrus reticulata*) varieties budded onto a number of rootstock varieties (Volkameriana, C35,  
152 Cleopatra Mandarin, Trifoliata, Swingle Citrumelo and Citrange), was established in October  
153 2005. The trees were planted at a spacing of 5 m x 2 m. The actual monitoring and  
154 measurements were initiated in August 2006. The trees were managed and fertilized  
155 following current commercial practices, although the amounts of applied fertilizer varied.  
156 The soils of the site are alkaline (Class IIIA), with red sandy loam from the surface to 90-cm  
157 depth, and loam below (90 to 150 cm). The total organic carbon content is very low (0.4%) in  
158 the first 30 cm, and below 0.25% in the remainder of the root zone. The climate is  
159 characterized as dry, with warm to hot summers and mild winters. The total rainfall during  
160 the experimental period from 21 August 2006 (DOY 233) to 20 August 2007 (DOY 232) was  
161 187 mm (Fig. 1), which was slightly below average for the area. Potential evapotranspiration  
162 is normally high and equal to 1400 mm per year. Mild frost conditions occur during the  
163 winter months. Weather data were collected from an automated weather station located  
164 within the research station.

165

## 166 2.2. Irrigation, fertigation and measurements

167

168 Irrigation water was supplied through a surface drip irrigation system, with drip lines  
169 placed on both sides of the tree line at a distance of 60 cm. Laterals had 1.6 Lh<sup>-1</sup> pressure  
170 compensating online drippers spaced at 40 cm, resulting in 10 drippers per tree. Irrigation  
171 was performed weekly/ bi-weekly, depending on the plant requirement, and the total seasonal  
172 irrigation was 432.8 mm.

173 The crop was irrigated to replace estimated crop evapotranspiration ( $ET_C$ ) for previous  
174 days. Reference crop evapotranspiration ( $ET_0$ ) was calculated using the FAO 56 method  
175 (Allen et al., 1998).  $ET_C$  was calculated using the equation:

$$176 \quad ET_C = ET_0 \cdot K_c \cdot A_c \quad (1)$$

177 where  $K_c$  is the crop coefficient and  $A_c$  is the crop age coefficient. The  $K_c$  values were  
178 compiled by the Irrigated Crop Management Service (ICMS) at Rural Solutions, South  
179 Australia.  $K_c$  values were taken from the FAO 56 report and adjusted for the Southern  
180 Hemisphere.  $A_c$  was used to correct  $ET_0$  for the age of the crop and its impact on canopy area  
181 (RMCWMB, 2009). Mandarin is an evergreen tree that requires nitrogen throughout the year.  
182 Nitrogen fertilizer was applied as ammonium nitrate and mono ammonium phosphate. The  
183 amount and timing of fertilizers injected into the irrigation water during the crop growth  
184 season is shown in Fig. 2. Total seasonal amounts of applied ammonium nitrate and mono  
185 ammonium phosphate fertilizers were equal to 508.1 and 139.4 kg ha<sup>-1</sup>, respectively. While  
186 irrigation was applied continuously during multiple hours, fertigation was applied during a  
187 one hour interval.

188 Water for irrigation was pumped directly from the Murray River. The salinity of the  
189 irrigation water ( $EC_w$ ) was monitored daily, and ranged between 0.09 and 0.19 dS m<sup>-1</sup>, well  
190 below the  $EC_w$  threshold for irrigation of orange, a close relative of mandarin (1.1 dS m<sup>-1</sup>;  
191 Ayers and Westcot, 1989).

192 Daily soil water content measurements were collected using Sentek® EnviroSCAN®  
193 logging capacitance soil water sensors, installed adjacent to the drip line (approximately 10  
194 cm away from the dripper) at depths of 10, 25, 50, 80, and 110 cm. The EnviroSCAN probes  
195 were calibrated for the experimental site by the gravimetric method.

196 Soil water was sampled on a weekly basis using SoluSAMPLERS™ (Biswas, 2006;  
197 Biswas and Schrale, 2007). The SoluSAMPLER is a porous ceramic cup connected to a PVC

198 sample reservoir and the tubing from the reservoir to the soil surface, which is used to apply  
199 suction and then extract soil solution within 24 hours. The experimental site had  
200 SoluSAMPLERs located at depths of 25, 50, 100, and 150 cm at a horizontal distance of 10  
201 cm from the drip emitter. The SoluSAMPLERs used in this study were developed at the  
202 South Australian Research and Development Institute (SARDI) and are distributed by Sentek  
203 Pty, Ltd.

204 The extracted soil solution was analysed to determine  $EC_{sw}$  and the  $NO_3^-$ -N content.  
205 Nitrate was determined by the Auto-analyser (cadmium reduction) procedure of Maynard and  
206 Kalra (1993).

207

### 208 *2.3. Modelling software*

209

210 The HYDRUS-2D software package (Šimůnek et al., 2011) was used to simulate the  
211 transient two-dimensional movement of water and solutes in the soil. This program  
212 numerically solves the Richards' equation for variably-saturated water flow, and advection-  
213 dispersion equations for both heat and solute transport. The model additionally allows  
214 specification of root water uptake, which affects the spatial distribution of water, salts and  
215 nitrate between irrigation cycles. The solute transport equation considers the advective-  
216 dispersive transport in the liquid phase, as well as diffusion in the gaseous phase. The  
217 theoretical part of the model is described in detail in the technical manual (Šimůnek et al.,  
218 2011) and in Šimůnek et al. (2008).

219

### 220 *2.4. Input parameters*

221

#### 222 *2.4.1. Soil hydraulic properties*

223

224 Soil hydraulic properties were described using the van Genuchten-Mualem constitutive  
 225 relationships (van Genuchten, 1980). The parameters for these constitutive relationships  
 226 (except for the 120-150 cm soil depth) were optimised using data from a lysimeter  
 227 experiment (Phogat et al., 2013b) (Table 1) involving similar soils as in the current study.

228

#### 229 2.4.2. Root water uptake

230

231 The spatial root distribution is defined in HYDRUS-2D according to Vrugt et al. (2001a):

$$232 \quad \Omega(x, z) = \left[ 1 - \frac{z}{z_m} \right] \left[ 1 - \frac{x}{x_m} \right] e^{-\left( \frac{p_z}{z_m} |z^* - z| + \frac{p_x}{x_m} |x^* - x| \right)} \quad (2)$$

233 where  $x_m$  and  $z_m$  are the maximum width and depth of the root zone (cm), respectively,  $z^*$  and  
 234  $x^*$  describe the location of the maximum root water uptake, from the soil surface in the  
 235 vertical direction ( $z^*$ ) and from the tree position in the horizontal direction ( $x^*$ ), and  $p_x$  and  $p_z$   
 236 are empirical coefficients.

237 We considered a simple root distribution model, in which the roots of young mandarin  
 238 trees expanded horizontally into all available space between tree lines ( $x_m = 200$  cm), were  
 239 concentrated mainly below the drip emitter ( $x^* = 60$  cm,  $z^* = 20$  cm) where water and  
 240 nutrients were applied, and extended to a depth of 60 cm ( $z_m = 60$  cm). The parameters  
 241 defining the maximum root water uptake in vertical and horizontal directions ( $z^*$  and  $x^*$ )  
 242 were also based on our earlier experience in similar studies (Phogat et al. 2012ab, 2013ab).  
 243 No significant volume of roots was found outside of the specified area in field observations.

244 The reduction of root water uptake due to the water stress,  $\alpha_1(h)$ , was described using the  
 245 well-known piecewise linear relation, developed by Feddes et al. (1978):

$$246 \quad \alpha_1(h) = \begin{cases} 0, & h > h_1 \text{ or } h \leq h_4 \\ \frac{h-h_1}{h_2-h_1}, & h_2 < h \leq h_1 \\ 1, & h_3 < h \leq h_2 \\ \frac{h-h_4}{h_3-h_4}, & h_4 < h \leq h_3 \end{cases} \quad (3)$$

247 where  $h_1$ ,  $h_2$ ,  $h_3$ , and  $h_4$  are the threshold parameters. Water uptake is at the potential rate  
 248 when the pressure head is between  $h_2$  and  $h_3$ , decreases linearly when  $h > h_2$  or  $h < h_3$ , and  
 249 becomes zero when  $h < h_4$  or  $h > h_1$ . The following parameters of the Feddes et al. (1978)  
 250 model were used:  $h_1 = -10$ ,  $h_2 = -25$ ,  $h_3 = -200$  to  $-1000$ ,  $h_4 = -8000$  cm, which were taken  
 251 from Taylor and Ashcroft (1972) for orange.

252 The reduction of root water uptake due to the salinity stress,  $\alpha_2(h_\phi)$ , was described by  
 253 adopting the Maas and Hoffmann (1977) salinity threshold and slope function. The salinity  
 254 threshold ( $EC_7$ ) for orange (closely related to mandarin) corresponds to a value for the  
 255 electrical conductivity of the saturation extract ( $EC_e$ ) of  $1.7 \text{ dS m}^{-1}$ , and a slope ( $s$ ) of 16%.  
 256 As required by HYDRUS-2D, these values were converted into  $EC_{sw}$ , assuming that the  
 257  $EC_{sw}/EC_e$  ratio was 2, which is a common approximation used for soil water contents near  
 258 field capacity in light-textured soils (U.S. Salinity Laboratory Staff, 1954; Skaggs et al.,  
 259 2006).

260 Plant uptake of non-adsorbing nutrients such as nitrate is controlled mainly by mass flow  
 261 of water uptake (Barber, 1995). Therefore, it was assumed that nitrate was either passively  
 262 taken up by the tree with root water uptake (Šimůnek and Hopmans, 2009) or moved  
 263 downward with soil water.

264

265 *2.4.3. Solute parameters*

266

267 Soil solution salinity ( $EC_{sw}$ ) distribution in soil was modelled as a non reactive solute (e.g.,  
268 Skaggs et al., 2006; Ramos et al., 2011; Wang et al., 2014). These studies demonstrated that  
269 this approach can be successfully used in environments under intensive irrigation and  
270 fertigation management. Additionally, Ramos et al. (2011) reported that similar salinity  
271 distributions were obtained when this simple approach of  $EC$  modelling using HYDRUS was  
272 compared with much more complex predictions involving consideration of  
273 precipitation/dissolution and ion exchange as done with UNSATCHEM, particularly when  
274 the soil solution is under-saturated with calcite and gypsum.

275 Nitrogen transport was simulated by means of a sequential first-order decay chain,  
276 implemented in HYDRUS-2D. Hence, N reaction or transformation processes, other than  
277 nitrification, were not considered. Similar assumptions have also been made in previous  
278 studies involving modelling of the nitrate transport in soil (Ramos et al., 2011, 2012). We also  
279 assumed that inherent soil organic N was mineralised directly into  $NO_3^-$ -N, consistent with  
280 other studies (Wang et al., 2010; Tafteh and Sepaskhah, 2012).

281 Nitrate ( $NO_3^-$ -N) was assumed to be present only in the dissolved phase (with the  
282 distribution coefficient,  $K_d = 0 \text{ cm}^3 \text{ g}^{-1}$ ). Ammonium ( $NH_4^+$ -N) was assumed to adsorb to the  
283 solid phase with a  $K_d$  value of  $3.5 \text{ cm}^3 \text{ g}^{-1}$  (e.g., Hanson et al., 2006; Ramos et al., 2012). The  
284 nitrification of  $NH_4^+$ -N to  $NO_3^-$ -N thus acts as a sink for  $NH_4^+$ -N and as a source for  $NO_3^-$ -N.  
285 First-order rate constants for solutes in the liquid and solid phases were set to be  $0.2 \text{ d}^{-1}$ .  
286 These were taken from a review of published data presented by Hanson et al. (2006), and  
287 represent the centre of the range of reported values.

288 The longitudinal dispersivity ( $\epsilon_L$ ) was considered to be 20 cm and the transverse  
289 dispersivity ( $\epsilon_T$ ) was taken as one-tenth of  $\epsilon_L$ . These values have been optimised in similar  
290 studies involving solute transport in field soils (e.g., Cote et al., 2003; Mallants et al., 2011).

291

## 292 2.4.4. Initial and boundary conditions

293

294 A time-variable flux boundary condition was applied to a 20 cm long boundary directly  
295 below the dripper, centred on 60 cm from the top left corner of the soil domain (Fig. 3). The  
296 flux boundary condition with a flux  $q$  was defined as:

$$297 \quad q = \frac{\text{volume of water applied / day}}{\text{surface wetted area}} \quad (4)$$

298 where the volume of water applied ( $L^3$ ) varied for different irrigation events and was  
299 calculated by multiplying the dripper discharge rate by irrigation time, and the surface wetted  
300 area ( $L^2$ ) was approximately  $800 \text{ cm}^2$  (i.e.,  $20 \text{ cm} \times 40 \text{ cm}$ ). The length of the boundary was  
301 selected to ensure that all water could infiltrate into the soil without producing positive  
302 surface pressure heads, because positive pressure heads at the flux boundary could make the  
303 numerical code unstable. During irrigation, the drip line boundary was held at a constant  
304 water flux,  $q$ . The atmospheric boundary condition was assumed for the remainder of the soil  
305 surface during periods of irrigation, and for the entire soil surface during periods between  
306 irrigation. A no-flow boundary condition was established at the left and right edges of the soil  
307 profile, to account for flow and transport symmetry. A free drainage boundary condition was  
308 assumed at the bottom of the soil profile. All these boundary conditions are illustrated in Fig.  
309 3. The mathematical details of applying the boundary conditions to a domain similar to the  
310 current one can be obtained from Phogat et al. (2012a).

311 The initial soil water content distribution was based on EnviroSCAN measured values and  
312 varied from 0.1 to  $0.25 \text{ cm}^3 \text{ cm}^{-3}$  in the soil domain (0-150 cm). Measured values of  $EC_{sw}$  and  
313  $\text{NO}_3^-$ -N in the soil were used as initial conditions in the model. The  $EC_{sw}$  varied from 0.8-1.5  
314  $\text{dS m}^{-1}$  and  $\text{NO}_3^-$ -N concentrations ranged between 0.16-1.07  $\text{mmol}_{(c)} \text{ L}^{-1}$  in the soil profile  
315 (0-150 cm).

316 The third-type Cauchy boundary conditions were imposed at the soil surface and at the  
317 free drainage boundary for solute transport ( $EC_{sw}$ ,  $NH_4^+$ -N and  $NO_3^-$ -N) and no flux  
318 boundary was imposed on the sides of the domain.

319

#### 320 *2.4.5. Flow domain and simulation*

321

322 In this approach, the drip tubing can be considered as a line source (Fig. 3), because in a  
323 twin line drip irrigation system with closely spaced drippers the wetted pattern from adjacent  
324 drippers merges to form a continuous wetted strip along the drip lines (Falivene et al., 2005).  
325 Water movement was therefore treated as a two-dimensional (in the vertical plane) process  
326 (Skaggs et al., 2004). Our field observations of the wetting pattern on the soil surface during  
327 experiments also supported this approach. The transport domain was set as a rectangle with a  
328 width of 250 cm (half of the lateral spacing between tree rows) and a depth of 150 cm. The  
329 transport domain was discretised into 2172 finite element nodes, which corresponded to 4191  
330 triangular elements (Fig. 3). Observation nodes corresponded to the locations where  
331 EnviroSCAN probes (depths of 10, 25, 50, 80, 100, and 110 cm) and SoluSAMPLERs  
332 (depths of 25, 50, 100, and 150 cm) were installed, at a distance of 10 cm from the emitter  
333 source (Fig. 3).

334

#### 335 *2.4.6. Estimation of potential evaporation and transpiration*

336

337 HYDRUS-2D requires daily estimates of potential evaporation ( $E_s$ ) and transpiration ( $T_p$ ).  
338 In this study, these parameters were obtained by combining the daily values of reference  
339 evapotranspiration ( $ET_0$ ), determined by the FAO Penman–Monteith method, and the dual  
340 crop coefficient approach (Allen et al., 1998, Allen and Pereira, 2009), as follows:

$$341 \quad ET_C = (K_{cb} + K_e) ET_0 \quad (5)$$

342 where  $ET_C$  is the evapotranspiration ( $LT^{-1}$ ),  $K_{cb}$  is the basal crop coefficient, which represents  
343 the plant transpiration component, and  $K_e$  is the soil evaporation coefficient. Standard  
344 mandarin  $K_{cb}$  values (Allen et al., 1998) were adjusted for the local climate, taking into  
345 consideration crop height, wind speed, and minimum relative humidity averages for the  
346 period under consideration. The values of daily potential transpiration ( $T_p$ ) and soil  
347 evaporation ( $E_s$ ) thus obtained (Fig. 4) were used as time-variable boundary conditions (see  
348 Fig. 3) in the model, along with the precipitation received at the site during the experimental  
349 period. The seasonal  $T_p$  amounted to 696 mm and  $E_s$  to 174 mm. The maximum  $T_p$  of 4.4 mm  
350 occurred on 10<sup>th</sup> January 2007 (DOY 10), when the most adverse weather conditions  
351 occurred.

352

### 353 *2.5. Scenario analysis for controlling deep drainage and N losses*

354

355 The nitrogen balance for the mandarin crop was evaluated for two fertigation strategies.  
356 First, the fertigation pulse was applied at the beginning of each irrigation event (Fert A).  
357 Second, the fertigation pulse was applied near the end of each irrigation event (Fert B). It is a  
358 common practice that irrigation water is initially and at the end free of fertilizer, to ensure a  
359 uniform fertiliser application and flushing of the drip lines (Gärdenäs et al., 2005). Therefore,  
360 fertigation applications were simulated to either start one hour after irrigation started or to  
361 end one hour before irrigation stopped.

362 Nitrate management strategies also include a judicious manipulation of irrigation and N  
363 fertilizer applications, and increasing or decreasing the frequency of applications. These  
364 interventions should improve N uptake by plants and reduce N leaching out of the plant root  
365 zone (Harrison, 1994). The evaluated scenarios are described in Table 2. Scenario, S1,  
366 illustrates the impact of applying the same volume of water in small irrigation events (<5

367 mm). Scenarios S2 and S3 then represents the reduction of the irrigation volume application  
368 by 10 and 20%, respectively. Scenarios S4 and S5 are based on decreasing the nitrogen  
369 application by 10 to 20%, respectively, while scenarios S6 and S7 represent a combined  
370 reduction in irrigation and fertigation by 10 to 20%, respectively. Five scenarios (S8 to S12)  
371 were executed, in which irrigation was reduced during the second half of the crop season, i.e.,  
372 between January and August, by 10, 20, 30, 40, and 50%, respectively.

373

#### 374 2.6. Statistical analysis

375

376 A mean absolute error (*MAE*) has been reported (Willmott and Matsuura, 2005) to be a  
377 good parameter for comparing modelling results with observed values. It was calculated by  
378 comparing weekly measured (*M*) and corresponding HYDRUS-2D simulated (*S*) values of  
379 water contents, electrical conductivities of soil solution ( $EC_{sw}$ ), and nitrate concentrations  
380 ( $NO_3^-$ -N) in soil as follows:

$$381 \quad MAE = \frac{1}{N} \sum_{i=1}^N |M_i - S_i| \quad (6)$$

382 Here, *N* is the number of comparisons.

383

### 384 3. Results and Discussion

385

#### 386 3.1. Moisture distribution

387

388 The water contents measured weekly by EnviroSCAN at different depths (10, 25, 50, 80  
389 and 100 cm) at a horizontal distance of 10 cm from the dripper, and corresponding values  
390 simulated by HYDRUS-2D during the entire growing season are illustrated in Fig. 5. The  
391 measured water contents remained similar at 10 ( $0.2 \text{ cm}^3 \text{ cm}^{-3}$ ) and 80 cm ( $0.1 \text{ cm}^3 \text{ cm}^{-3}$ ) cm,

392 fluctuated between 0.1 and 0.2  $\text{cm}^3\text{cm}^{-3}$  at 25 and 50 cm, and stayed higher than 0.2  $\text{cm}^3\text{cm}^{-3}$   
393 at 110 cm soil depths throughout the growing season, indicating a favourable moisture regime  
394 in the crop root zone. However, the simulated water contents were lower than the measured  
395 values during the initial period at a depth of 10 cm and during the mid period at a depth of  
396 110 cm. The simulated values matched the measured values more closely at soil depths of 25  
397 and 50 cm, which is the most active root zone for water and nutrient uptake for citrus  
398 (Mikhail and El-Zeftawi, 1979). However, the profile average water distribution matched  
399 well.

400 The *MAE* between weekly measured and simulated moisture content values across all  
401 locations varied from 0.01 to 0.04  $\text{cm}^3\text{cm}^{-3}$ , indicating a good agreement between the two sets  
402 of values (Table 3). Slightly higher temporal *MAE* values during the mid-season agreed well  
403 with the variation shown in Fig. 5. Similarly, the *MAE* values at 10, 25, 50, 80, and 110 cm  
404 soil depths (Table 3) at a 10 cm lateral distance from the dripper also revealed that the  
405 variation between measured and simulated water contents remained between 0.02 to 0.04  
406  $\text{cm}^3\text{cm}^{-3}$ . However, the differences were slightly higher at 10 cm depth (0.04  $\text{cm}^3\text{cm}^{-3}$ ) as  
407 compared to greater depths (0.02-0.03  $\text{cm}^3\text{cm}^{-3}$ ). Higher variations at the surface depth (10  
408 cm) are to be expected because this part of the soil profile is influenced by soil evaporation,  
409 which peaks in day time and is low at night time, while the assumption of a constant  
410 atmospheric boundary flux for daily time steps in the model (Ramos et al., 2012) deviated  
411 from the actual transient conditions existing at the surface boundary. Other studies (Vrugt et  
412 al., 2001b; Skaggs et al., 2010; Phogat et al., 2012ab; Phogat et al., 2013ab; Ramos et al.,  
413 2012) also showed a similar magnitude of variations between measured and predicted water  
414 contents.

415

416 *3.2. Soil solution salinity distribution*

417

418 Comparison of simulated electrical conductivities of soil solution ( $EC_{sw}$ ) with weekly  
419 measured values at different depths (25, 50, 100 and 150 cm) are shown in Fig. 6. Despite of  
420 low irrigation water salinity ( $0.09\text{-}0.2\text{ dSm}^{-1}$ ) and low initial soil salinity ( $0.8\text{-}1.5\text{ dS m}^{-1}$ ), the  
421 measured  $EC_{sw}$  increased in the soil with the onset of irrigation at all depths, except at 150 cm  
422 where the increase in salinity occurred only after Dec 2006. Subsequently, a decreasing trend  
423 was observed in  $EC_{sw}$  later in the season. The higher amount of irrigation compared to  $ET_C$   
424 and an significant amount of precipitation (Fig. 1) during this period resulted in a reduction in  
425 soil solution salinity.

426 On the other hand, the model over-predicted  $EC_{sw}$  at a depth of 25 cm from Oct to Dec  
427 2006 and under-predicted it at a depth of 100 cm during the same period. However, at a depth  
428 of 150 cm, simulated values remained constant till Jan 2007, indicating a delayed response.  
429 The increase in simulated  $EC_{sw}$  values was delayed at 100 and 150 cm depths as compared to  
430 measured values. Both set of values matched well at a depth of 50 cm and the profile average  
431 of  $EC_{sw}$  also showed a close match.

432 It is significant to note that irrigation with good quality water ( $EC_w < 0.2\text{ dS m}^{-1}$ ) in our  
433 study led to the development of significant levels of measured  $EC_{sw}$  ( $0.34\text{ to }2.32\text{ dS m}^{-1}$ ;  
434 mean  $1.17\text{ dS m}^{-1}$ ). However, the  $EC_{sw}$  values remained below the threshold of salinity  
435 tolerance ( $EC_e = 1.7\text{ dS m}^{-1}$  or  $EC_{sw} = 3.4\text{ dS m}^{-1}$ ) of orange throughout the season (Ayers  
436 and Westcot, 1989; ANZECC and ARMCANZ, 2000).

437 The (temporal) MAEs between weekly measured and simulated  $EC_{sw}$  in the soil ranged  
438 from  $0.08\text{ to }0.76\text{ dS m}^{-1}$  (Table 3), which are acceptable for a complex and highly dynamic  
439 soil system, with the exception of a few divergent values obtained between mid October and  
440 December (DOY 290-365). The disagreement in  $EC_{sw}$  values during this period was  
441 correlated with corresponding fluctuations and low values of water contents, especially at soil  
442 depths of 10 and 25 cm and this variability was transferred to the  $EC_{sw}$  values. Differences

443 between measured and simulated  $EC_{sw}$  values at 50 cm depth were relatively higher ( $MAE$   
444  $=0.47$  dS  $m^{-1}$ ) than at other depths (Table 3). The mean  $MAE$  at 25, 100, and 150 cm depths  
445 ranged from 0.19 to 0.36 dS  $m^{-1}$ , showing a good agreement with the measured values at  
446 these depths.

447 The spatial distribution of  $EC_{sw}$  in the soil profile at various dates is depicted in Fig. 7. It  
448 can be seen that salts remained restricted to roughly the upper 50 cm of the soil profile until  
449 December (between 28/11/2006 and 17/01/2007 in Fig. 7). The salts mass was later pushed  
450 deeper due to high rainfall (55 mm in January, 29% of seasonal rain). The downward  
451 movement of salts continued in February and March (8/03/2007 in Fig. 7), because in March  
452 the amount of irrigation was higher than  $ET_C$  (Fig. 2). It is pertinent to note here that the  $EC_{sw}$   
453 distribution under the dripper remained lower as compared to the adjoining soil at all times,  
454 because a continuous water application in this region pushes the salts towards the outer  
455 boundary of the wetting front. The drainage flux during and after March transported salts  
456 vertically downwards, thereby making the soil directly beneath the dripper relatively salt free  
457 by the end of the season. Applying additional water at the end of the season could be a  
458 strategy to create a salt free rootzone which may encourage vigorous root development, and  
459 assist the plant growth in the ensuing season.

460

### 461 3.3. Nitrate nitrogen distribution

462

463 Comparison of weekly measured and daily simulated nitrate-nitrogen ( $NO_3^-$ -N)  
464 concentrations at different depths (25, 50, 100 and 150 cm) in the soil profile is illustrated in  
465 Fig. 8. Over-prediction was observed at a depth of 25 cm from Oct to Nov 2006, which  
466 coincided with similar over-prediction for salinity. Similarly, both measured and simulated  
467 values matched well at a depth of 50 cm, while a delayed response in predicted nitrate

468 contents was observed at lower depths. However, a fairly good correspondence was observed  
469 between profile averaged  $\text{NO}_3^-$ -N contents. The temporal *MAE* values for  $\text{NO}_3^-$ -N ranged  
470 from 0.1 to 1.97  $\text{mmol}_{(\text{c})} \text{L}^{-1}$  (Table 3). Similar differences between measured and HYDRUS-  
471 2D simulated values were also reported in another study (Ramos et al. (2012) involving  
472 simulations of nitrogen under field cropped conditions. Additionally, *MAE* at a 25 cm depth  
473 (Table 3) had a higher value (1.52  $\text{mmol}_{(\text{c})} \text{L}^{-1}$ ) than at greater depths (0.63 to 0.73  $\text{mmol}_{(\text{c})} \text{L}^{-1}$ ).  
474 A similar match of nitrate distributions has been reported in other studies as well (Ajdary  
475 et al., 2007; Ramos et al., 2012; Tournebize et al., 2012).

476 The reason for differences in  $EC_{\text{sw}}$  and  $\text{NO}_3^-$ -N values may be partially due to the fact that  
477 model reports point values, whereas the SoluSAMPLER draws in solution from a sampling  
478 area of a certain volume, the size of which depends on the soil hydraulic properties, the soil  
479 water content, and the applied suction within the ceramic cup (Weihermuller et al., 2005;  
480 Ramos et al., 2012; Phogat et al., 2012a). Hence the measured parameters considered in  
481 modelling may not represent the inherent spatial variability of the soil. In addition, while a  
482 homogeneous soil environment is assumed by the model, the field site could be far more  
483 heterogeneous and anisotropic. Also, the model simulations considered only a 2D movement  
484 of nitrogen and the nitrification process, while more complex nitrate processes (e.g.,  
485 mineralisation, ammonification, denitrification, immobilization through carbon-nitrogen  
486 complex formation and microbial interaction) were not taken into account. Ramos et al  
487 (2012) documented numerous factors influencing the correspondence between measurements  
488 and simulations of water contents and solute concentrations in the soil under drip irrigation  
489 conditions and these factors are relevant also for the present investigation. These factors,  
490 including those mentioned above, may modify the error in the simulated  $\text{NO}_3^-$ -N values.

491 The simulated movement of nitrate-nitrogen ( $\text{NO}_3^-$ -N) in the soil under a mandarin tree at  
492 various dates is shown in Fig. 9. Nitrate fertigation increased the nitrogen content in the soil

493 with time, as is evident from an increasing size of the concentration plume below the dripper  
494 as the season progressed. This indicates that the plant was not able to take up all nitrogen  
495 added through fertigation, and thus nitrogen built up in the soil over time, leading to a  
496 maximum concentration values in January (17/01/07 in Fig. 11). Ultimately, nitrogen started  
497 moving downwards after late January, when there was high rainfall and total water additions  
498 exceeded  $ET_C$ . Alva et al. (2006) also detected greater variations in  $\text{NO}_3^-$ -N concentrations in  
499 the 0-15 cm depth horizon, as compared to greater depths in a field experiment involving  
500 citrus. The seasonal  $\text{NO}_3^-$ -N concentrations in the domain varied from 0.01- 7.03  $\text{mmol}_{(c)} \text{L}^{-1}$ .  
501 Hutton et al. (2008) reported higher mobilization of nitrate at a shallower depth under drip  
502 irrigation of grapevine, and seasonal root zone nitrate concentrations ranging between 0-  
503 11.07  $\text{mmol}_{(c)} \text{L}^{-1}$  in the Murrumbidgee Irrigation Areas in Australia.

504 As the season continued and plant uptake was reduced, excess water further mobilised  
505 nitrate-nitrogen out of the root zone, as is evident from 27/04/07 and beyond (Fig. 9). At the  
506 end of the crop season, little nitrogen remained in the soil system, and what did remain was  
507 well beyond the reach of the plants. This nitrogen is expected to continue leaching  
508 downwards over time and become a potential source of nitrate-nitrogen loading to the ground  
509 water.

510 Additionally, peak  $\text{NO}_3^-$ -N concentrations in the soil profile (7.03  $\text{mmol}_{(c)} \text{L}^{-1}$ ) and in  
511 drained water ( $\text{NO}_3^-$ -N concentration at the 150 cm depth, 2.14  $\text{mmol}_{(c)} \text{L}^{-1}$ ) were  
512 significantly higher than the Australian environmental standard (ANZECC and ARMCANZ,  
513 2000) for protection of 80% (17 mg  $\text{NO}_3 \text{L}^{-1} = 0.27 \text{mmol}_{(c)} \text{NO}_3^- \text{N L}^{-1}$ ) and 95% of species  
514 (0.7 mg  $\text{NO}_3 \text{L}^{-1} = 0.01 \text{mmol}_{(c)} \text{NO}_3^- \text{N L}^{-1}$ ). The  $\text{NO}_3^-$ -N concentrations in the soil solution  
515 also occasionally exceeded the level of Australian drinking water quality standard (NRMMC,  
516 2011) for nitrate (100 mg  $\text{NO}_3 \text{L}^{-1} = 1.61 \text{mmol}_{(c)} \text{NO}_3^- \text{N L}^{-1}$ ). High levels of nitrate-nitrogen  
517 below the crop root zone are undesirable, as some recharge to groundwater aquifers can

518 occur, in addition to flow into downstream rivers, which are used for drinking water and  
519 irrigation. These findings are consistent with other studies (Barlow et al., 2009; Correll et al.,  
520 2010), in which high nitrate concentrations in drainage water under drip and furrow fertigated  
521 irrigation systems have been reported.

522

### 523 3.4. Water and nitrogen balance

524

525 The seasonal water balance was computed from cumulative fluxes calculated by  
526 HYDRUS-2D. Estimated water balance components above and below the soil surface under a  
527 mandarin tree are presented in Table 4. It can be seen that in a highly precise drip irrigation  
528 system, a large amount of applied water (210.9 mm) drained out of the rootzone, even though  
529 the amount of irrigation applied was based on estimated  $ET_C$ . This drainage corresponded to  
530 33.5% of applied water, and occurred because highly permeable light textured soils, such as  
531 those found in this study, are prone to deep drainage whenever the water application exceeds  
532  $ET_C$ . The drainage amount in our study falls within the range of recharge fluxes to  
533 groundwater reported by Kurtzman et al. (2013) under citrus orchards in a semiarid  
534 Mediterranean climate. Mandarin root water uptake amounted to 307.3 mm, which  
535 constitutes about 49% of applied water. Root water uptake slightly increased (3.5%) when the  
536 model was run without considering solute (salt) stress (not shown here), which is not a  
537 significant difference. It further substantiates the results obtained for seasonal  $EC_{sw}$  in Fig. 6,  
538 where salinity remained below threshold ( $3.4 \text{ dS m}^{-1}$ ) over the season. Evaporation accounted  
539 for 17.7% of the total water applied through irrigation and rainfall. The modelling study over-  
540 estimated the sink components of the water balance by 4.79 mm (0.77%, Table 4).

541 There were major differences between water input and output from January 2007 onwards  
542 (Fig. 10). During this period, irrigation (I) and precipitation (P) significantly exceeded tree

543 water uptake ( $S\_W$ ), which eventually resulted in deep drainage ( $Dr\_W$ ) from March 2007  
544 onwards. Therefore, current irrigation scheduling requires adjustment during this period. This  
545 illustrates how simulations were helpful in evaluating the overall water dynamics in soil  
546 under the mandarin tree.

547 The nitrogen balance is presented in Table 5. The nitrogen fertilizer was applied either in  
548 the form of  $NH_4^+$  or  $NO_3^-$ , but  $NH_4^+$  transforms quickly to  $NO_3^-$  through the process of  
549 nitrification. Model simulations showed that nitrification of  $NH_4^+$  was very rapid and most of  
550 the  $NH_4^+$ -N converted to  $NO_3^-$  before it moved to a depth of 20 cm, and no traces of  $NH_4^+$   
551 were observed below this depth. It is apparent that the nitrification of  $NH_4^+$  took place in the  
552 upper soil layer, which contains organic matter and moisture that supports microorganisms  
553 (*Nitrosomonas* and *Nitrobacter*), facilitating the nitrification of  $NH_4^+$ . Though  $NH_4^+$  was  
554 initially nitrified to  $NO_2^-$  and consequently to  $NO_3^-$ ,  $NO_2^-$  was short-lived in the soil and  
555 decayed to  $NO_3^-$  quickly. Therefore, the simulated plant  $NH_4^+$ -N uptake was only  $0.71 \text{ kg ha}^{-1}$   
556 <sup>1</sup>. Hence, the  $NO_3^-$ -N form was responsible for most of the plant uptake, corresponding to  
557 about 85% of the applied nitrogen. The monthly N applications were slightly higher than  
558 plant uptake during the flowering (August-October) and fruit growth (January-March)  
559 periods (Fig. 11). However, the monthly uptake was slightly higher than the N application  
560 between these periods.

561 High frequency of N applications in small doses resulted in similar nitrogen uptake  
562 efficiency (61 to 75%) in citrus as in other studies (Syvertsen and Smith, 1995; Quinones et  
563 al., 2007). Similarly, Scholberg et al. (2002) reported doubling of nitrogen use efficiency as a  
564 result of frequent application of N in a dilute solution. Slightly higher uptake ( $1.73 \text{ kg ha}^{-1}$ )  
565 was recorded when fertigation was applied in second last hour of an irrigation event (Fert B),  
566 as compared to when it was applied early in the irrigation event (Fert A, Table 5). Hence, it  
567 can be concluded that timing of fertigation does not have a major impact in a normal

568 fertigation schedule with small and frequent N doses within an irrigation event in light  
569 textured soils. Similar results were also obtained in our earlier study in a lysimeter planted  
570 with an orange tree (Phogat et al., 2013b), which revealed that timing of fertilizer N  
571 applications in small doses in an irrigation event with a low emitter rate had little impact on  
572 the nitrogen uptake efficiency.

573 Nitrate-nitrogen leaching accounted for only 15% of the applied nitrogen (Table 5).  
574 Monthly N balance (Fig. 11) revealed that most of the N leaching happened between March  
575 2007 and August 2007, which was correlated with the extent of deep drainage occurring  
576 during this period.  $\text{NO}_3^-$ -N losses ranging from 2-15% were illustrated by Paramasivam et al.  
577 (2002) and Alva et al. (2006), attributable in part to an improved management of N, which  
578 could be a contributor in the current estimation.

579

### 580 *3.5. Strategies for controlling water and nitrogen losses*

581

582 In our study, it is evident that there were significant deep drainage (33%) and nitrate-  
583 nitrogen leaching losses (15%), which could be reduced by appropriate management. Hence,  
584 different simulations involving the reduction of irrigation and fertigation applications during  
585 the whole or part of the crop season were conducted, to optimize water and nitrogen uptake  
586 and to reduce their losses from the soil (Table 6).

587 Increasing the irrigation frequency with short irrigation events (S1) while maintaining the  
588 same irrigation volume, had no impact on deep drainage (Dr\_W) and N leaching (Dr\_N).  
589 However, the seasonal salinity increased by 11% compared to the standard practice. This  
590 confirms that the current irrigation schedule followed with respect to the irrigation frequency  
591 seems to be optimal under the experimental conditions. In S2, Dr\_W and Dr\_N were reduced  
592 by 14.4 and 19%, respectively, but salinity increased by 11%. However, a sustained reduction

593 in irrigation by 20% (S3) eventually reduced the  $Dr_W$  and  $Dr_N$  by 28.1 and 38.3%,  
594 respectively, at the expense of a 4.9% decline in plant water uptake, but with a 4% increase in  
595 N uptake. However, salinity increased by 25.8% compared to the normal practice, which  
596 would likely have a significant impact on plant growth.

597 Scenarios S4 and S5 were based on decreasing the nitrogen application by 10 and 20%,  
598 resulting in a decrease in N leaching by 7.4 and 14.8%, respectively, along with a much  
599 higher reduction in plant N uptake (10.4% in S4 and 19.7% in S5), suggesting that the  
600 reduction in the fertilizer application alone is not a viable option to control N leaching under  
601 standard conditions. A combined reduction in irrigation and fertigation by 10% (S6) further  
602 reduced N leaching by 5.5%, compared to reducing irrigation alone (S2), but at the same time  
603 plant N uptake was reduced by 5% more than in S2. Similarly, reducing irrigation and N  
604 application by 20% (S7) produced a pronounced reduction in N leaching (46.4%) and water  
605 drainage (28%), but it also resulted in a decrease in plant N uptake by 15.8% and water  
606 uptake by 4.8%, compared to normal practice. At the same time, salinity increased by 25.8%,  
607 which is similar to S3. The reduction in plant water and N uptake would have a major impact  
608 on plant growth and yield, and would adversely impact the sustainability of this expensive  
609 irrigation system. Hence, reducing fertilizer applications does not seem to be a good  
610 proposition under the current experimental conditions, as it results in an appreciable decline  
611 in plant N uptake. However, Kurtzman et al. (2013) reported that a 25% reduction in the  
612 application of N fertilizer is a suitable agro-hydrological strategy to lower the nitrate flux to  
613 groundwater by 50% under different environmental conditions. Rather, reducing irrigation  
614 alone seems to be a better option to control the deep drainage and N leaching losses under the  
615 conditions encountered at the experimental site.

616 Additionally, it is worth noting that in S3 and S7 the salinity ( $EC_{sw}$ ) during a period  
617 between October and December at a depth of 25 cm, and during December at a depth of 50

618 cm, increased considerably, and was higher than the threshold level (Fig. 12), confirming that  
619 a sustained reduction in irrigation (S3) and fertigation (S7) is not a viable agro-hydrological  
620 option for controlling water and N leaching under the mandarin orchard.

621 However, it seems unnecessary to reduce irrigation applications uniformly across the  
622 season as suggested by Lidon et al. (2013). Rather, irrigation could more profitably be  
623 reduced only during a particular time period when excess water was applied. The water and N  
624 balance data in our study revealed that an imbalance between water applications and uptake  
625 happened during the second half of the crop season, i.e., from January till August 2007,  
626 resulting in maximum drainage (Fig. 10) and N leaching (Fig. 11), coinciding with the fruit  
627 maturation and harvesting stage. Hence, there is a need to reschedule irrigation within this  
628 period, rather than reducing water applications throughout the entire season. Keeping this in  
629 mind, the following 5 scenarios (S8 to S12, Table 6) were executed, in which irrigation was  
630 reduced during the second half of the crop season, i.e., between January and August, by 10,  
631 20, 30, 40, and 50%, respectively.

632 Scenarios S10, S11, and S12 showed an enormous potential for reducing water and N  
633 losses. In S10, Dr\_W and Dr\_N were reduced by 8 and 4% more than in S7, N uptake was  
634 increased by 6.9% (compared with a reduction in S7), and salinity was also 4% less than in  
635 S7, which seems quite promising. On the other hand, in S11 and S12, the Dr\_W and Dr\_N  
636 were reduced to a greater extent (50-58% water and 70-80% N leaching) than in S10, and soil  
637 salinity increased substantially (40.3 and 58.7% higher than normal practice), due to a  
638 considerable reduction in the leaching fraction. This is also shown in Fig. 12, which shows  
639 that monthly soil solution salinity ( $EC_{sw}$ ) in S11 and S12 at the 25 and 50 cm soil depths  
640 increased dramatically between January and August. Although  $EC_{sw}$  remained below the  
641 threshold level, except at a 50 cm depth in S12 during March 2007, there is a significant  
642 likelihood of it increasing further in subsequent seasons, which would ultimately impact the

643 growth and yield of mandarin trees. Hence, under current conditions, Scenario S10 represents  
644 the best option to control excessive water and N losses, and high salinity, and to increase the  
645 water and N efficiency for mandarin trees. Other permutations and combinations, involving  
646 fertilizer reductions along with S10, did not provide further improvements in controlling  
647 water and N leaching. It is concluded that simulations of irrigation and fertilizer applications,  
648 using HYDRUS, can be helpful in identifying strategies to improve the water and N  
649 efficiency for drip irrigation systems of perennial horticultural crops.

650

#### 651 **4. Conclusions**

652

653 This study demonstrates the importance of combining strategic monitoring with numerical  
654 modelling to assess water movement, salinity distribution, and nitrogen management under  
655 drip irrigation systems in young mandarin orchards in Australia. HYDRUS-2D was used to  
656 predict seasonal water, salt, and nitrate dynamics in soils. Modelling results were compared  
657 with measured values of moisture content, soil solution salinity ( $EC_{sw}$ ), and nitrate-nitrogen  
658 ( $NO_3^-$ -N) in the soil profile during the complete season.

659 Graphical and statistical comparisons of measured and simulated values of water contents,  
660  $EC_{sw}$ , and  $NO_3^-$ -N concentrations in the soil under a mandarin tree showed a consistent  
661 performance of HYDRUS-2D for modelling water, salinity, and nitrogen transport. The  
662 temporal mean absolute errors (*MAE*) for water contents,  $EC_{sw}$ , and  $NO_3^-$ -N concentrations  
663 were within acceptable limits. However, *MAE* showed divergent values at shallow depths  
664 (10-25 cm) due to the assumption of a constant surface boundary flux during a particular  
665 daily time step, which deviated from normal diurnal fluctuations in the real-time evaporation  
666 flux. Other reasons for deviations between predicted and observed  $NO_3^-$ -N contents were  
667 attributed to the model considering only a simple linear movement of nitrogen, rather than

668 considering all complex processes (e.g., mineralisation, ammonification, denitrification,  
669 immobilization through carbon-nitrogen complex formation, and microbial interactions).

670 The simulated water and nutrient balances showed that the irrigation scheduling at the  
671 experimental site from December onwards needed to be modified in order to control deep  
672 drainage (33.5% of applied water) and nitrate leaching (15% of applied  $\text{NO}_3^-$ -N). Sustained  
673 reduction of irrigation and/or fertilization by 10 to 20% reduced water (14-28%) and  $\text{NO}_3^-$ -N  
674 (19-46%) losses appreciably, but these strategies reduced the leaching fraction and/or plant N  
675 uptake to a level where root zone  $EC_{sw}$  increased to substantially higher values than the  
676 recommended threshold ( $3.4 \text{ dS m}^{-1}$ ) and plant N uptake was reduced (7-20%), both of which  
677 may affect plant growth and yield, and in turn would adversely impact the sustainability of  
678 expensive irrigation systems.

679 Other evaluated scenarios focused on reducing irrigation (by 10-50%) between January  
680 and August, when a mismatch between irrigation applications and plant uptake was observed.  
681 A 30% reduction in irrigation during this period provided the best scenario, in which both  
682 water and  $\text{NO}_3^-$ -N leaching were reduced by 37 and 52%, respectively, and plant N uptake  
683 was increased by 7%, compared to the normal practice. However, a further reduction in  
684 irrigation by 40 and 50% reduced the water (50-58%) and  $\text{NO}_3^-$ -N (70-80%) losses to a great  
685 extent, but increased salinity in the root zone to a level much higher than the tolerance  
686 threshold of mandarin.

687 This study forms the basis for future evaluation of irrigation, salinity, and nitrate-nitrogen  
688 dynamics under drip fertigation systems in fields with horticultural trees, and for future  
689 exploration of ways to fine-tune irrigation schedules in order to better control excessive  
690 drainage and N losses. However, there is a need to further improve the modelling estimates  
691 by considering all processes of the nitrogen cycle in the soil system. It is concluded that such  
692 studies would help in improving irrigation and fertigation programs for horticultural crops

693 irrigated with drip irrigation systems, and would lead to more efficient and less  
694 environmentally detrimental crop management practices.

695

#### 696 **Acknowledgements**

697

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702

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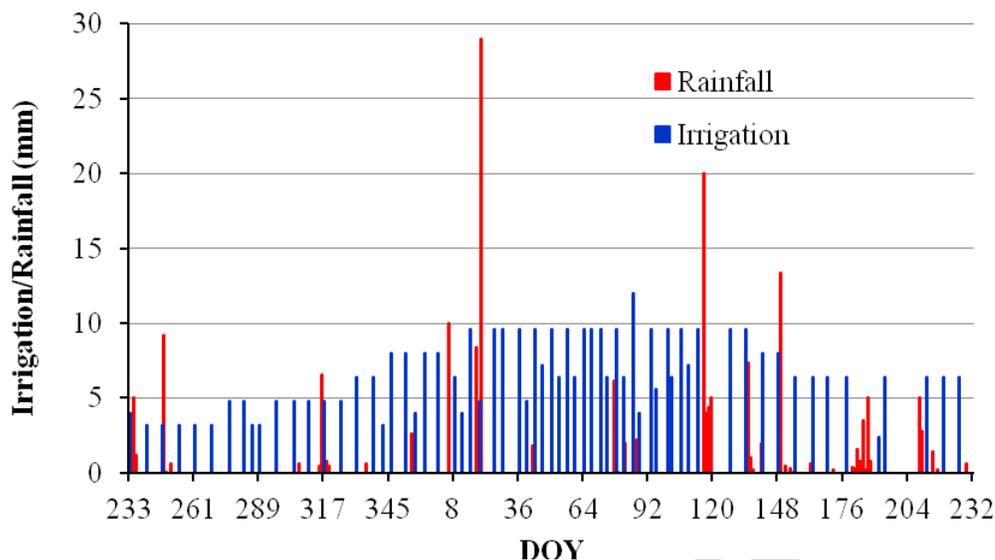
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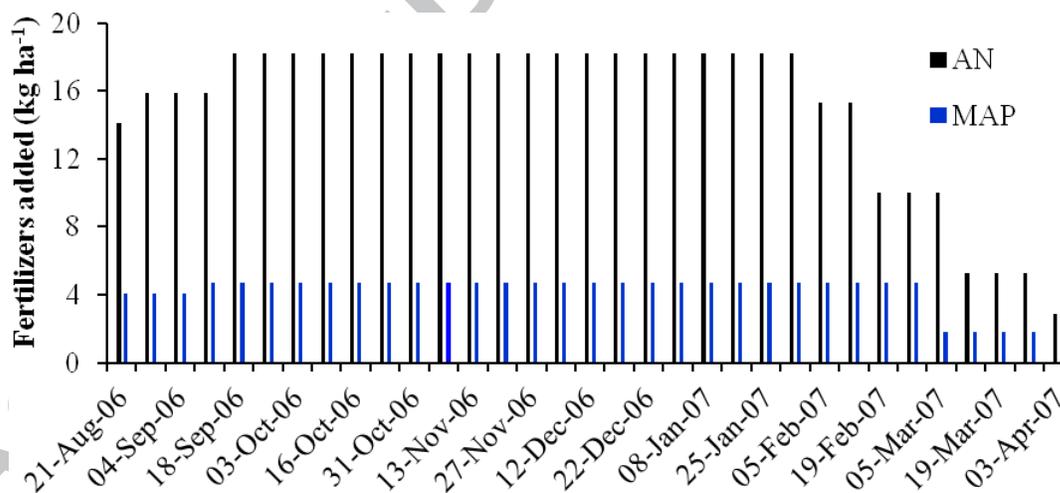
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908 **Fig. 1.** Rainfall received (red bars) and irrigation applied (blue bars) during the experimental  
 909 period (21 August 2006 to 20 August 2007). DOY represents the Julian day of the year.

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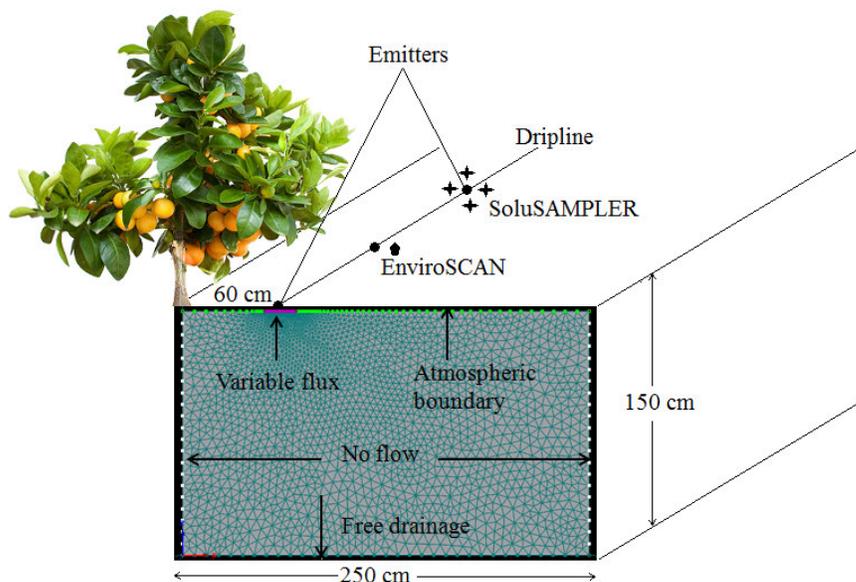
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914 **Fig. 2.** Fertigation schedule followed during the experimental period (21 August 2006 to 20  
 915 August 2007) (AN represents Ammonium nitrate, black bars; MAP represents Mono-  
 916 ammonium phosphate, blue bars).



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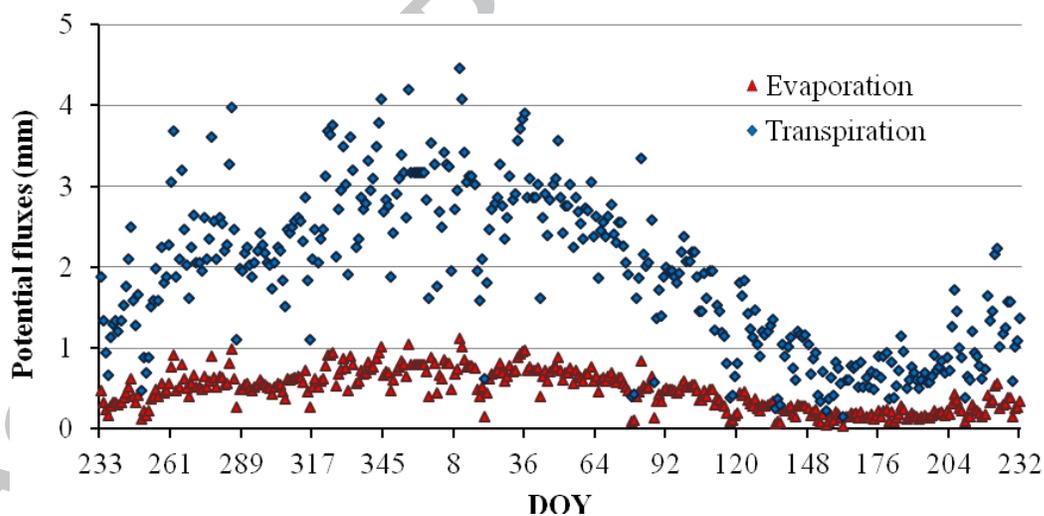
919 **Fig. 3.** A schematic view of the model domain (2D) showing considered boundary conditions

920 based on the experimental layout, plant and drip spacing, and locations of monitoring

921 equipments.

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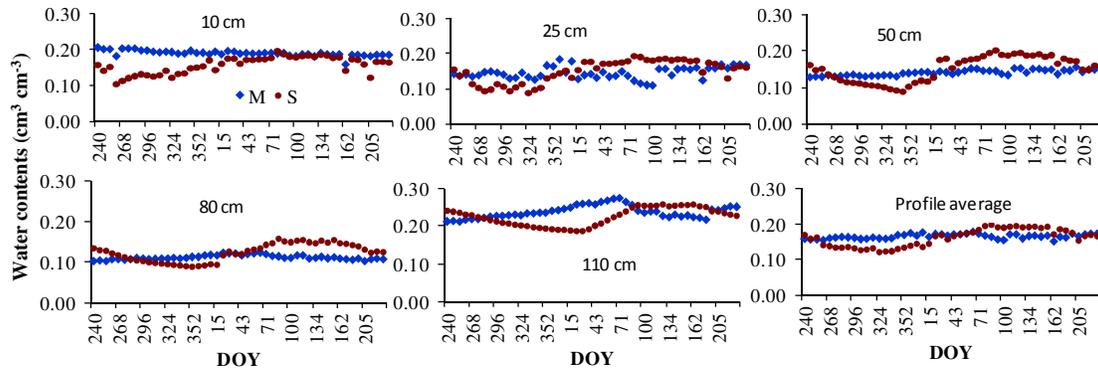


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925 **Fig. 4.** Daily potential transpiration ( $T_p$ ) and potential soil evaporation ( $E_s$ ) estimated using

926 the dual crop coefficient approach during the study period. DOY represents the Julian day of

927 the year 2006-07.

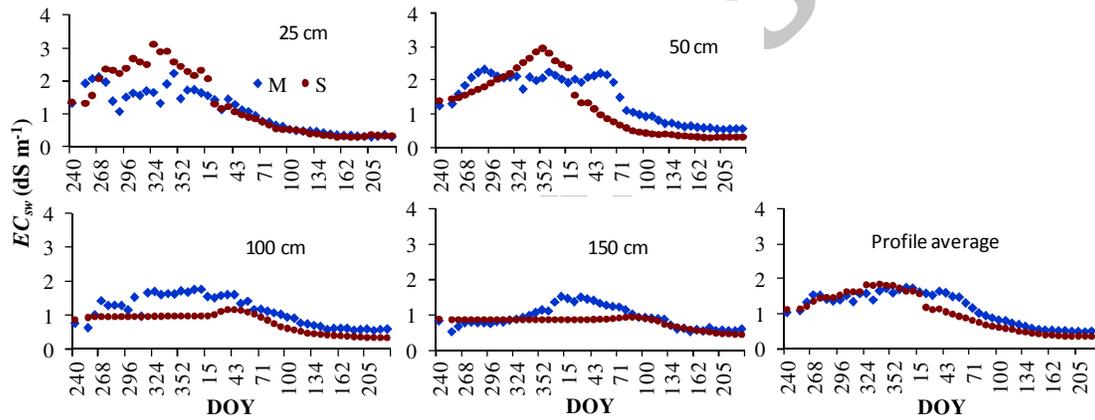


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929 **Fig. 5.** Comparison of weekly measured (M) and simulated (S) water contents at indicated  
 930 depths in the soil profile under a mandarin tree.

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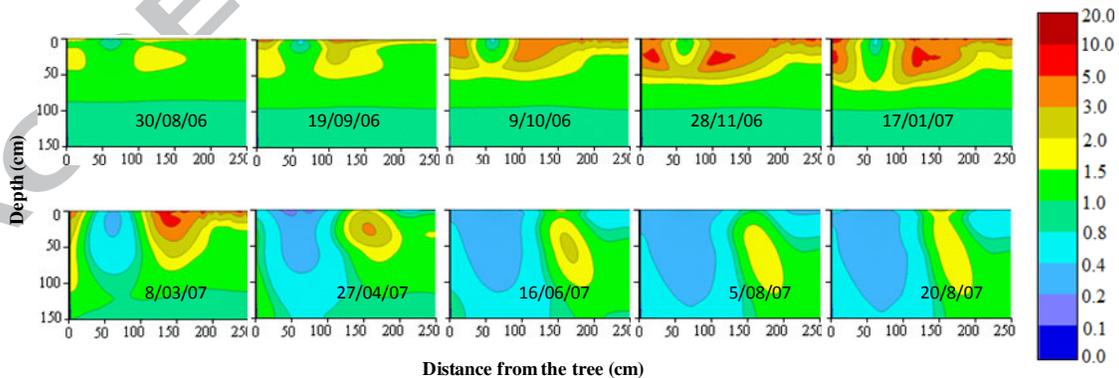
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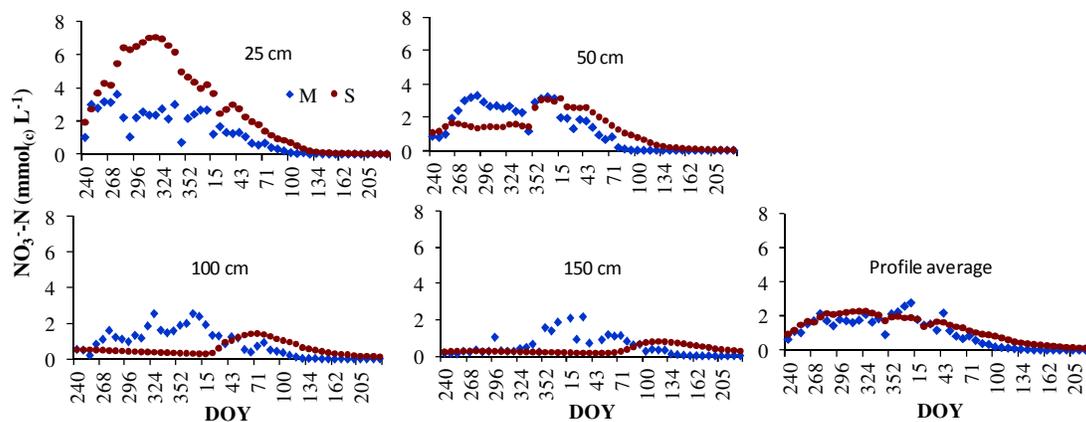
934 **Fig. 6.** Comparison of measured (M) and simulated (S) values of soil solution salinity ( $EC_{sw}$ )  
 935 at indicated depths in the soil profile under a mandarin tree.

936



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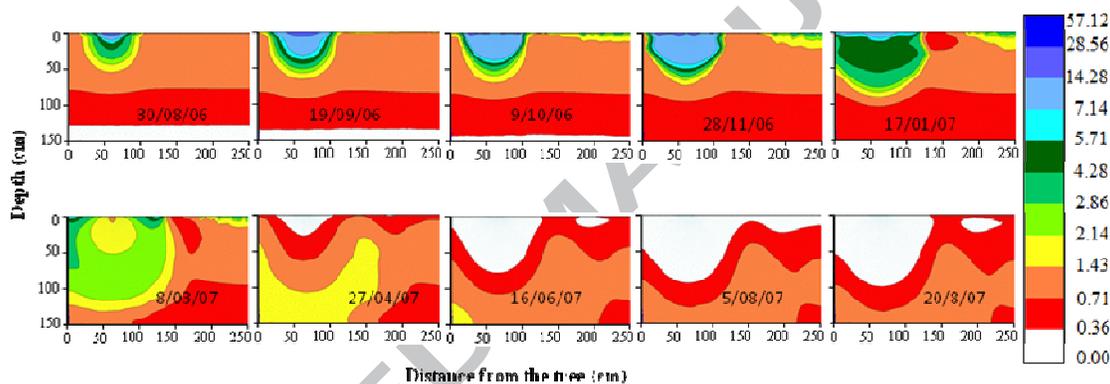
938 **Fig. 7.** Spatial distribution of simulated soil solution  $EC$  ( $EC_{sw}$ ,  $dS\ m^{-1}$ ) in the soil profile at  
 939 indicated times.



940

941 **Fig. 8.** Comparison of measured (M) and simulated (S) values of soil solution nitrate-nitrogen  
 942 ( $\text{NO}_3^-$ -N) at indicated depths in the soil profile under a mandarin tree.

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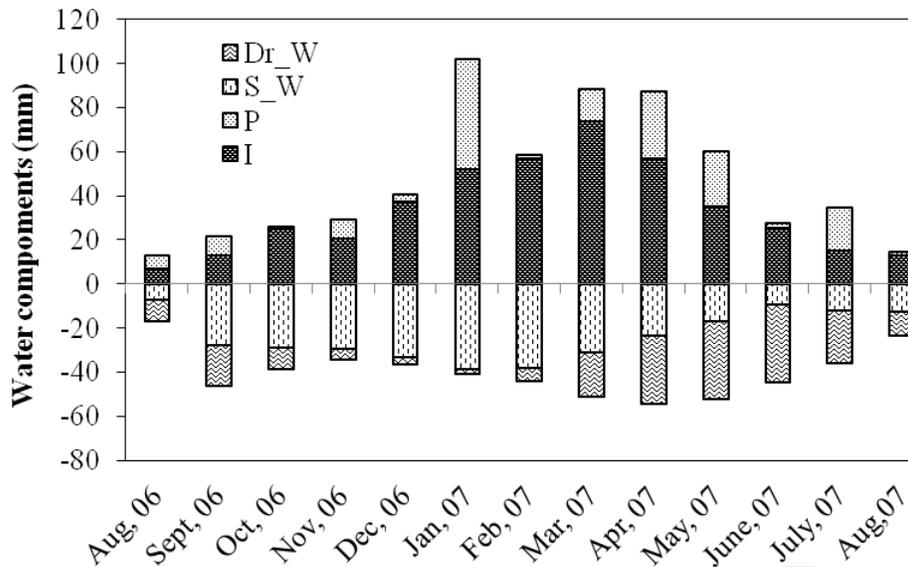
945 **Fig. 9.** Spatial distribution of simulated soil solution  $\text{NO}_3^-$ -N ( $\text{mmol}_{(c)} \text{L}^{-1}$ ) in the soil profile  
 946 at indicated times.

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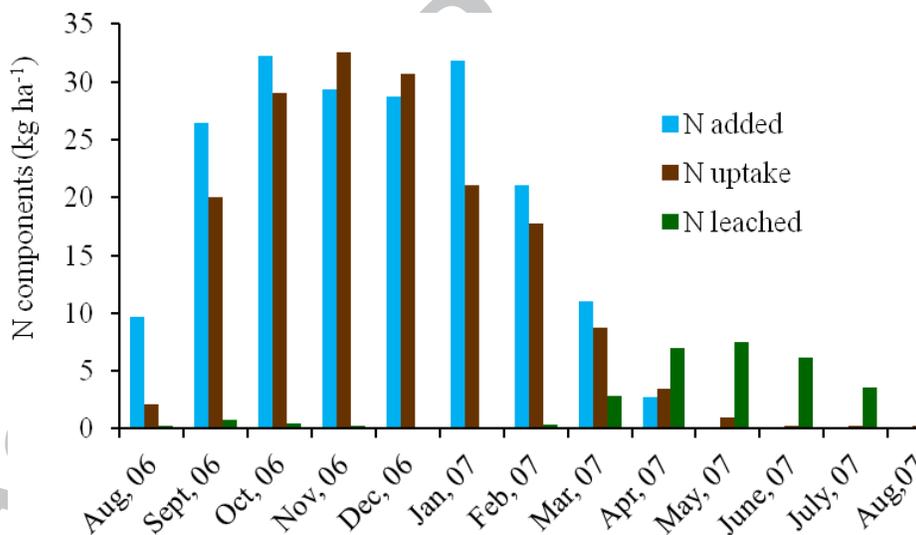
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952 **Fig. 10.** Monthly irrigation (I, mm), precipitation (P), water uptake (S\_W) by a mandarin  
 953 tree, and deep drainage (Dr\_W) from the soil during the study period (from Aug 2006 to Aug  
 954 2007).

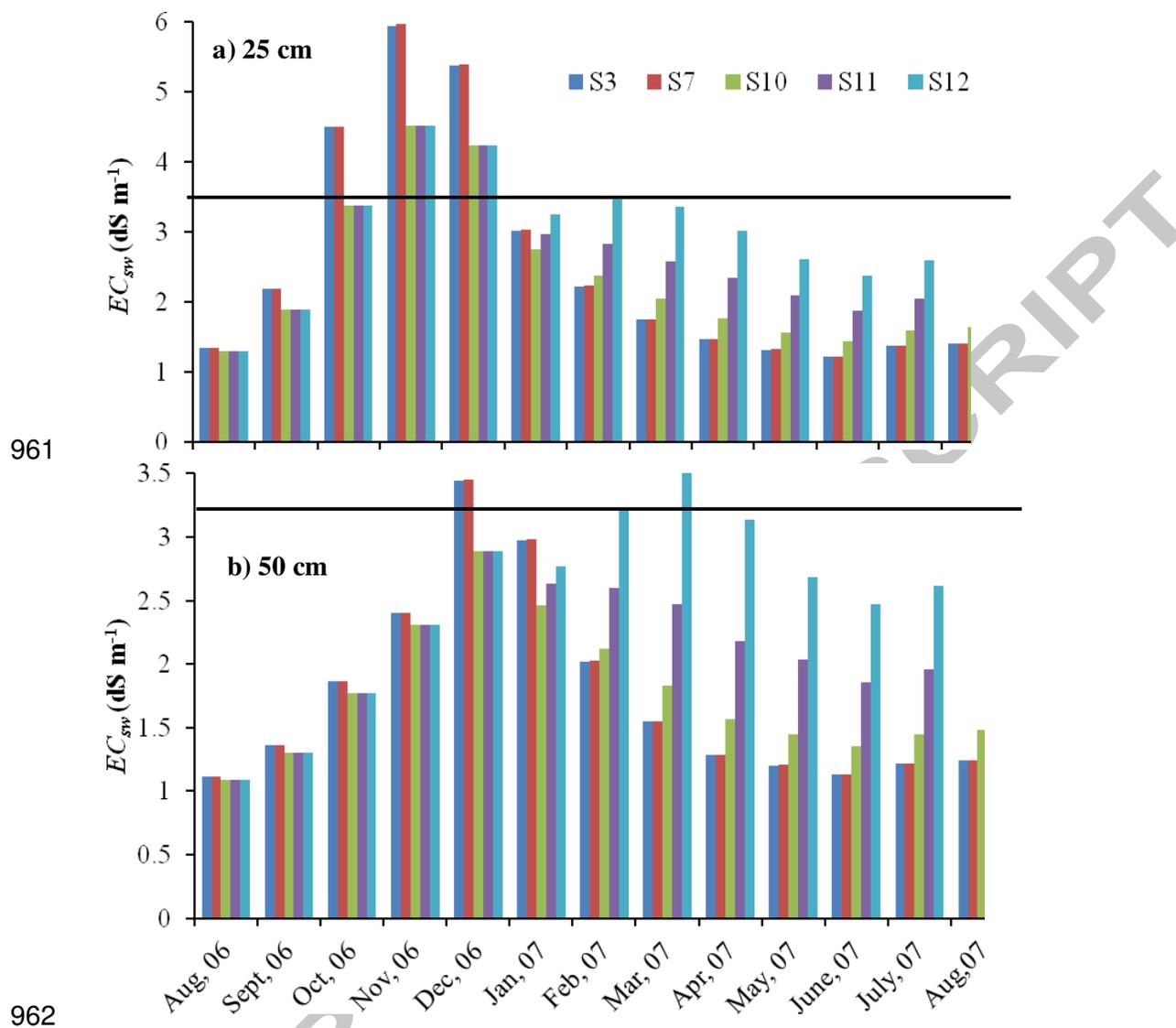
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956

957 **Fig. 11.** Simulated monthly values of nitrogen added (N added), nitrogen uptake (N uptake)  
 958 by a young mandarin tree, and nitrogen leached (N leached) from the soil during the study  
 959 period (from Aug 2006 to Aug 2007).

960



963 **Fig. 12.** Monthly average soil solution salinity ( $EC_{sw}$ ,  $dS m^{-1}$ ) at a) 25 cm  
 964 depth in the soil profile under different scenarios (see Table 6). Horizontal lines show the  
 965 threshold salinity level for citrus ( $3.4 dS m^{-1}$ ).

966

967 **Table 1**

968 Soil hydraulic parameters used in the modelling study (the residual water content  $\theta_r$ , the  
 969 saturated water content  $\theta_s$ , van Genuchten shape parameters ( $\alpha$ ,  $n$  and  $l$ ), and the saturated  
 970 hydraulic conductivity  $K_s$ ).

Soil depth (cm)	Texture	$\theta_r$ (cm <sup>3</sup> cm <sup>-3</sup> )	$\theta_s$ (cm <sup>3</sup> cm <sup>-3</sup> )	$\alpha$ (cm <sup>-1</sup> )	$n$	$K_s$ (cm day <sup>-1</sup> )	$l$
0-30	loamy sand	0.060	0.37	0.0294	1.92	116.88	0.5
30-60	loamy sand	0.060	0.36	0.0268	1.91	107.04	0.5
60-90	loamy sand	0.050	0.34	0.0308	1.99	113.28	0.5
90-120	loam	0.050	0.33	0.0300	1.85	79.20	0.5
120-150	loam	0.046	0.36	0.0346	1.41	27.89	0.5

971

972 **Table 2**

973 Various scenarios evaluated for optimising irrigation and fertigation of a mandarin orchard.

Scenario	Reduction in Irrigation (I) and/or Fertigation (F)
S1	All irrigation events $\leq 5$ mm
S2	10% less I during the entire season
S3	20% less I during the entire season
S4	10% less F during the entire season
S5	20% less F during the entire season
S6	10% less I & F during the entire season
S7	20% less I & F during the entire season
S8	10% less I during Jan-Aug, 07
S9	20% less I during Jan-Aug, 07
S10	30% less I during Jan-Aug, 07
S11	40% less I during Jan-Aug, 07
S12	50% less I during Jan-Aug, 07

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977

978 **Table 3**

979 Temporal and spatial mean absolute error (*MAE*) values between measured and simulated  
 980 water contents, soil solution salinities ( $EC_{sw}$ ), and nitrate-nitrogen ( $NO_3^-$ -N) concentrations.

Water content			$EC_{sw}$			$NO_3^-$ -N		
Temporal <i>MAE</i> values*								
<i>N</i> **	Mean	Range	<i>N</i> **	Mean	Range	<i>N</i> **	Mean	Range
	(cm <sup>3</sup> cm <sup>-3</sup> )			(dS m <sup>-1</sup> )			(mmol <sub>(c)</sub> L <sup>-1</sup> )	
48	0.03	0.01-0.04	47	0.34	0.08-0.76	48	0.89	0.10-1.97
Spatial <i>MAE</i> values								
<i>n</i> ***	Depth	Error	<i>n</i> ***	Depth	Error	<i>n</i> ***	Depth	Error
	(cm)	(cm <sup>3</sup> cm <sup>-3</sup> )		(cm)	(dS m <sup>-1</sup> )		(cm)	(mmol <sub>(c)</sub> L <sup>-1</sup> )
353	10	0.04	47	25	0.36	48	25	1.52
353	25	0.03	47	50	0.47	48	50	0.64
353	50	0.02	47	100	0.36	48	100	0.73
353	80	0.02	47	150	0.19	48	150	0.63
353	110	0.03	--	--	--	--	--	--

981 \*MAE for temporal data were calculated across 5 depths (i.e.  $n = 5$ ) at weekly interval of the  
 982 trial.

983 \*\* represents the number of weekly comparisons

984 \*\*\*represents number of values in each error calculation

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987 **Table 4**

988 Simulated components of the seasonal water balance under a young mandarin tree.

	Components	(mm)	(%)
Sources	Irrigation	432.68	69.16
	Rainfall	171.13	27.35
	Soil depletion	21.80	3.48
Sinks	Root water uptake	307.3	48.80
	Drainage	210.94	33.50
	Evaporation	111.56	17.70
	Water balance error	-4.79	-0.77 <sup>a</sup>

989 <sup>a</sup>Water balance error (%) =  $\left( \frac{\sum W_{source} - \sum W_{sink}}{\sum W_{source}} \right) \times 100$

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991 **Table 5**

992 Components of the nitrogen balance under a mandarin crop for fertigation at the beginning  
 993 (Fert A) and at the end (Fert B) of an irrigation event.

N source	N balance (kg ha <sup>-1</sup> )	Fert A	Fert B
NH <sub>4</sub> <sup>+</sup> -N	Soil <sub>initial</sub>	0	0
	Added	105.6	105.6
	Adsorbed on soil	0	0
	Uptake	0.71	0.71
	Leached	0	0
	Nitrification	104.9	104.9
	Soil <sub>end</sub>	0	0
NO <sub>3</sub> <sup>-</sup> -N	Soil <sub>initial</sub>	22.8	22.8
	Added	88.8	88.8
	Uptake	167.11	168.84
	Leached	31.3	31.1
	Soil <sub>end</sub>	20.21	20.09
<sup>a</sup> Mass balance error		-0.98	-1.63

994 <sup>a</sup>Mass balance error (%) =  $\left( \frac{\sum W_{input} - \sum W_{output}}{\sum W_{input}} \right) \times 100$

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1007 **Table 6**

1008 Percent increase (+)/decrease (-) in water uptake (S\_W), drainage (Dr\_W), N uptake (S\_N),

1009 N leaching (Dr\_N), and electrical conductivity of the soil solution (EC\_sw) in different

1010 scenarios of water and fertilizer applications, compared to the normal practice.

Scenario	Reduction in Irrigation (I)/ Fertigation (F)	S_W	Dr_W	S_N	Dr_N	EC_sw
S1	Irrigation events $\leq$ 5 mm	-0.25	-1.16	-1.01	0.06	11.29
S2	10% I, full season	-2.17	-14.40	2.64	-19.07	11.29
S3	20% I, full season	-4.88	-28.15	4.07	-38.29	25.81
S4	10% F, full season	0.18	0.06	-10.38	-7.40	0.00
S5	20% F, full season	0.21	0.03	-19.72	-14.76	0.00
S6	10% I & F, full season	-2.06	-14.25	-7.05	-24.60	11.29
S7	20% I & F, full season	-4.83	-28.18	-15.66	-46.43	25.81
S8	10% I, Jan-Aug, 07	-0.30	-12.74	1.90	-15.53	5.65
S9	20% I, Jan-Aug, 07	-0.87	-25.41	4.36	-32.97	13.71
S10	30% I, Jan-Aug, 07	-1.66	-37.16	6.89	-50.52	21.77
S11	40% I, Jan-Aug, 07	-2.68	-49.89	9.79	-69.53	40.32
S12	50% I, Jan-Aug, 07	-4.11	-57.91	12.76	-80.51	58.87

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1020 **Research Highlights**

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1022 • The seasonal water, salinity and nitrate distribution in soil was simulated by  
1023 HYDRUS-2D1024 • Deep drainage accounts for 33.5% and water uptake 49% of applied water by  
1025 mandarin crop

1026 • Model simulation predicted 15% leaching of applied nitrate as fertilizer

1027 • Higher N uptake recorded for fertigation during one hour before the last hour in an  
1028 irrigation event1029 • Irrigation cut by 30% during 2<sup>nd</sup> half of the crop season reduced drainage and N  
1030 leaching significantly

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