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Modeling floods in a dense urban area using 2D shallow water equations

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Summary A code solving the 2D shallow water equations by an explicit second-order scheme is used to simulate the severe October 1988 flood in the Richelieu urban locality of the French city of Nîmes. A reference calculation using a detailed description of the street network and of the cross-sections of the streets, considering impervious residence blocks and neglecting the flow interaction with the sewer network provides a mean peak water elevation 0.13 m lower than the measured flood marks with a standard deviation between the measured and computed water depths of 0.53 m. Sensitivity analysis of various topographical and numerical parameters shows that globally, the results keep the same level of accuracy, which reflects both the stability of the calculation method and the smoothening of results. However, the local flow modifications due to change of parameter values can drastically modify the local water depths, especially when the local flow regime is modified. Furthermore, the flow distribution to the downstream parts of the city can be altered depending on the set of parameters used. Finally, a second event, the 2002 flood, was simulated with the calibrated model providing results similar to 1988 flood calculation. Thus, the article shows that, after calibration, a 2D model can be used to help planning mitigation measures in a dense urban area.

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Introduction

Surface flood modeling in urban environment is a challenging task for a number of reasons: the presence of a large number of obstacles of varying shapes and length scales, the storages in the buildings, the complex geometry of the city, etc. However, if the intensity of the event is strong enough and the domain is a dense urban zone, it can be

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Nomenclature

dh	difference between computed peak water depth and flood mark measurement	n	Manning's roughness coefficient
\overline{dh}	average of dh	u	velocity along horizontal x-axis
F	local Froude number	u^*	bottom friction velocity
g	gravity acceleration	v	velocity along horizontal y-axis
h	water depth	z_b	bottom level
k	non-dimensional coefficient in viscosity calculation	ν_{eff}	effective kinematic viscosity
		σ	standard deviation

assumed that the majority of the flow passes by the streets and the junctions. The flow in the streets is mostly one-dimensional except near the junction but it was shown that at junctions (Weber et al., 2001) and bifurcations (Neary et al., 1999), the flow is basically 3D. Thus (Huang et al., 2002) performed a 3D calculation of flow at a junction and found that the results agreed with the measurements. However, (Khan et al., 2000) and (Shettar and Murthy, 1996) showed that 2D models can simulate the water surface and velocity field near the crossing and predict the distribution at a subcritical bifurcation.

Looking at the available literature in surface urban flood simulations, this statement seems accepted by many authors as it appears that most studies use two-dimensional models (Ishigaki et al., 2004). For an experimental urban flood case study and for real cases use full shallow water equation models or simplified models neglecting the inertial terms (Chen et al., 2004) to simulate urban floods. For these real cases, the origin of the flood can be the overflow of a river (Calenda et al., 2003), the invasion of land drainage flow from the upstream hills (Aronica and Lanza, 2005; Haider et al., 2003), the rain storm event on the modeled area or sewer drainage failure (Hsu et al., 2000; Mark et al., 2004; Schmitt et al., 2004). The building blocks can be either considered impervious or numerical methods for introduction of flow in the building blocks are proposed (Inoue et al., 2000). In most of the urban flooding numerical simulation references, only few validation data was available and so the quality of the simulations could not be precisely assessed. Furthermore, calibration processes and the influence of the chosen parameters were not proposed in the previously cited references.

In order to examine the conditions of use of a 2D model, the authors decided to investigate the case of the dense city of Nîmes that experienced two recent dramatic flood events in a topographical context of steep slopes and high velocities. The explicit finite volume numerical code "Rubar 20" (Paquier, 1995, 1998) was used to solve the full 2D shallow water equations. The October 1988 flood was used as a base simulation in order to discuss the sensitivity of such a calculation to particular modeling features: the level of detail in the description of the streets, the description of specific urban structures, the boundary conditions, the values of some numerical parameters. Then modeling of a second event (the September 2002 flood) was performed in order to validate the calibration of the model. Finally, a conclusion about the calibration process of such a model and its technical application for planning emergency measures is undertaken.

The governing equations and the numerical scheme

The system of 2D shallow water equations consists of three equations: one equation for continuity and two equations for the conservation of momentum in the two orthogonal directions.

$$\frac{\partial h}{\partial t} + \frac{\partial(hu)}{\partial x} + \frac{\partial(hv)}{\partial y} = 0 \quad (1)$$

$$\begin{aligned} \frac{\partial(hu)}{\partial t} + \frac{\partial(hu^2)}{\partial x} + \frac{\partial(huv)}{\partial y} + gh \frac{\partial h}{\partial x} \\ = -gh \frac{\partial z_b}{\partial x} - gn^2 \frac{u\sqrt{u^2 + v^2}}{h^{1/3}} + \nu_{\text{eff}} \frac{\partial}{\partial x} \left(h \frac{\partial u}{\partial x} \right) + \nu_{\text{eff}} \frac{\partial}{\partial y} \left(h \frac{\partial u}{\partial y} \right) \end{aligned} \quad (2)$$

$$\begin{aligned} \frac{\partial(hv)}{\partial t} + \frac{\partial(huv)}{\partial x} + \frac{\partial(hv^2)}{\partial y} + gh \frac{\partial h}{\partial y} \\ = -gh \frac{\partial z_b}{\partial y} - gn^2 \frac{v\sqrt{u^2 + v^2}}{h^{1/3}} + \nu_{\text{eff}} \frac{\partial}{\partial x} \left(h \frac{\partial v}{\partial x} \right) + \nu_{\text{eff}} \frac{\partial}{\partial y} \left(h \frac{\partial v}{\partial y} \right) \end{aligned} \quad (3)$$

where h is water depth, u and v are velocities along horizontal x- and y-axes, z_b is bottom level, n is Manning's roughness coefficient, g is gravity acceleration and ν_{eff} is effective cinematic viscosity.

Usually, it is assumed that the viscosity is constant throughout the flow field. However, the model (simpler than the k- ϵ model) defined by the following relation was also tested:

$$\nu_{\text{eff}} = khu^* \quad (4)$$

where k is a non-dimensional coefficient and the bottom friction velocity u^* is approximated by (5) instead of the usual relation from the friction coefficient in order to enhance the variations of the water surface elevation

$$u^* = \sqrt{gh \sqrt{\left(\frac{\partial(z_b + h)}{\partial x} \right)^2 + \left(\frac{\partial(z_b + h)}{\partial y} \right)^2}} \quad (5)$$

In the "Rubar 20" code (Paquier, 1995, 1998), the conservative form of Eqs. (1)–(3) above are solved on a grid constituted of quadrilaterals and triangles using an explicit second-order numerical scheme that is adapted from MUSCL approach (VanLeer, 1979). The selected scheme is able to calculate highly unsteady flows and can treat the transitions between subcritical and supercritical flows as ordinary points. Drying and wetting of cells are treated specifically

in the following way: a cell is considered as dry as long as the water volume entering an originally dry cell during one time step provides an average water depth in the cell below a minimal value ($0.01 \mu\text{m}$ in present study). During drying process, a similar option is used and leads to a mass conservation error that is usually less than 0.01% of the total mass. This numerical scheme is generally stable under the Courant Friedrichs Levy condition that limits the Courant number to values below unity. However, in the case of steep slopes, the time step should be further restricted; thus, modeling requires high calculation power.

The validation of the numerical scheme on very unsteady flows (against analytical solutions and experimental data sets) was extensively performed through the IAHR group on dam-break wave and during the CADAM and IMPACT European research projects that included comparisons with other numerical codes. Some of these experiments included dam break wave calculation in simple rectangular channels, in more irregular topography domains with introduction of blocks to simulate obstacles or buildings (Mignot and Paquier, 2003a,b) or on a real urban flood situation (Mignot and Paquier, 2004). All these tests proved the relevance of "Rubar 20" code to solve 2D shallow water equations in rapid flow and complex topographic situations. To complete the code validation in the domain of urban floods, some experimental tests of four rectangular channel junction flows were performed in the INSA Fluid Mechanics Laboratory in Lyon (France). Precise measurements of the observed flow patterns of supercritical junction flows were performed and compared to the water depths computed

with "Rubar 20" (Mignot et al., 2005b). It appeared that the code was able to predict with fair agreement the structure of the flow and the locations of the specific flow structures such as the hydraulic jumps, the recirculation zones and the eddying (Mignot et al., 2005a).

The 1988 flood and the presentation of the studied area

Nîmes is a French city located in a plain just downstream from seven hills surrounding its northern area. Since 1350, many records of severe floods have been recorded in the city and the flood of October, third 1988 in Nîmes was among the most dramatic observed events. It was caused by a rainstorm that generated up to 420 mm of rain in about 8 h on the northern hills (Desbordes et al., 1989) and led to the overflow of the watercourses to the north of the city and the subsequent inundation of various localities with water depths up to 3 m, causing extensive damages to property and human life (Bonneaud, 2002). (Desbordes et al., 1989) evaluated the return period of this event to 150–250 years.

The studied area known as "Richelieu", located in the north-eastern part of the city of Nîmes was one of the most dramatically affected zones. The dimensions of the area are about 1400 m along the north–south axis that is also the main principal flow direction, and a variable east–west width with a maximum and minimum of about 1050 and 220 m, respectively. A railway embankment runs

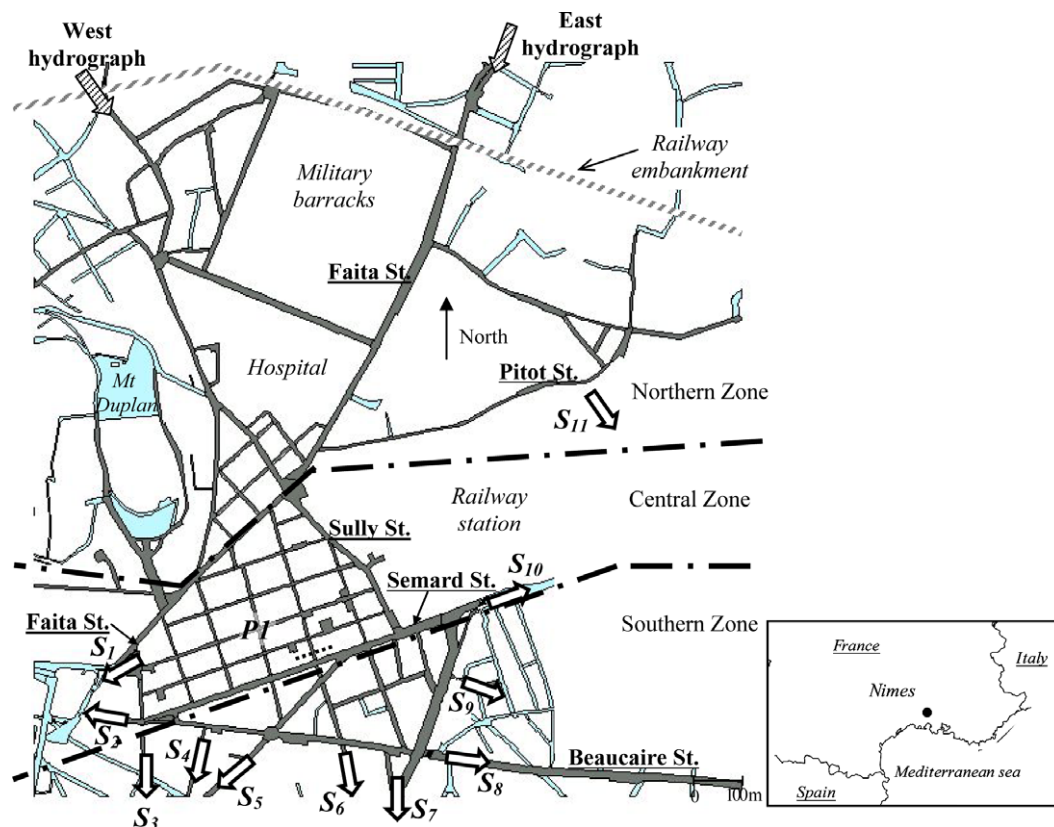


Figure 1 "Richelieu" area of the city of Nîmes with S_1 to S_{11} the outlets from the area.

all the way along the northern side, the western flank is formed by hills and the eastern flank is constituted by an impervious railway station (Fig. 1). One particularity of the city of Nîmes is that no river crosses the city, indeed when it rains on the region, the flow drained by the temporary watercourses enters the urban area and flows in the underground watercourses until reaching the Vistre river downstream from the city. However, when the discharges in the watercourses becomes larger than their conveyances, overflows occur upstream from the city and the water enters its northern parts following the natural slope (north to south), flowing within the streets. "Richelieu" is one of these localities, where the flow can enter through several street underpasses located in the railway embankment. The drop in ground level in the studied area is, respectively, 20 and 15 m from the western and eastern underpasses (Fig. 1) to the southern boundary, which represents an average slope higher than 1%. The upstream part of the city (Fig. 1) is composed of large building blocks areas such as the military barracks and the hospital with wide streets and few crossroads. The central part is a rather regular network with narrow streets (5–8 m wide) and right angle crossings between them. This area is interesting due to the relative abundance of the flood marks that demonstrates an average peak water depth of about one and a half meter. However, photos show that, in some specific points, waves occurred and water surface rose along the buildings, which means some uncertainties in the peak water level, confirmed by differences between couples of close water marks. Finally, the urban structure of the southern part is intermediate between the northern and central part, with average width streets and still a high main north–south slope.

Available data for the 1988 flood and their preliminary processing

The Public works department of the Gard County and the Technical Services of Nîmes city supplied the necessary primary data for the simulations. It included:

- About 200 cross-sectional profiles for the 60 streets located in the studied area. A typical street profile (Fig. 2) contained 11 points including the top and foot of both side walls (points 1 and 2), both pedestrian ways (points 3), both gutters (points 4 and 5) and the mid point of the road section (point 6).
- A map of the area showing the limits of the flooded area, the land use (built-up places, etc.) and some complementary ground elevations. The 99 flood mark

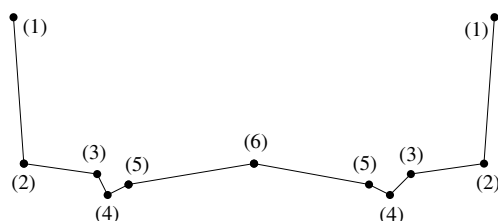


Figure 2 11-points street section profile.

elevation points measured immediately after the 1988 flood were also noted on the map showing the measurement locations. 85% of these flood marks are located in crossroads at the corner between two building walls and the remaining 15% are located within the streets on the façade of a building.

- Rainfall data measured at various locations around the area.

BCEOM French consulting firm performed an hydrological modeling (BCEOM et al., 2004) on the upstream basins (10 and 3.5 km²) starting from rainfall measurements in order to compute hydrographs at the upstream border of the Richelieu domain. The model is based on a conceptual linear tank approach similar to GR4 model (Perrin et al., 2003); it considers separately the hydrological processes on the urban areas and the runoff on grass land of the upstream basins. Then, for various hydrological events, the model was calibrated to fit with the observations of flood marks at the upstream end of the Richelieu area (railway line) and more upstream.

In this step, the sewage network was taken into account, the discharge capacity of the sewage network (4 and 7 m³/s for the pipes at the western and eastern domain entrance, respectively) was subtracted from the calculated surface inflow hydrographs. Nevertheless, it appears that the maximum capacity of the sewage network is less than 10% of total peak discharge in 1988, which is of the order of the hydrological results uncertainties. Furthermore, when the upstream hydrograph reached the studied area, the sewage network was already almost full due to the rain falling on the city. The result of this hydrological modeling consists in two discharge hydrographs that can be used for inputs at the northwest and northeast upstream boundary of the model. The hydrographs span a period of 13 h with a time step of 15 min (Fig. 3).

The cross-sectional street profiles (Fig. 2) constituted the basic structure of the calculation mesh. In order to obtain a complete description of the city, a linear interpolation between the intersecting streets permitted to obtain the input and output cross-sections of each street junction. Then the generation of the 121 points of each crossroad (11 points from both joining streets) was achieved by interpolating between the input and output sections. Additional points and information were introduced for the mesh of the "complex" topographical crossings. Building blocks were considered as watertight and details about them were

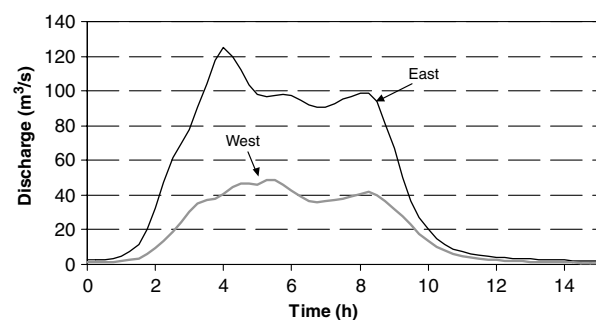


Figure 3 East and west hydrographs at the studied area border for the 1988 flood.

not introduced. Furthermore, the sewage network was not taken into account in the hydraulic study, assuming that the exchanges between the surface and the network could be neglected compared to the flow rates and water depths in the streets. From this construction, the DEM contains about 25,000 points for the streets, each one described by its three coordinates x , y , z .

Reference calculation

The basic calculation mesh is generated keeping all the points from the DEM and adding extra points along the streets. The specification of an average space step of 25 m along the streets generally leads the mesh generator to interpolate less than five additional sections along the streets. This choice of a rather long space step came from the hypothesis that, except in (or close to) the crossings, the flow is predominantly 1D. Furthermore, this reduces the calculation time but restricts the possibilities to include small scale urban structures and to simulate the complexity of the flow in the streets in the vicinity of the junctions. The mesh consists of 100 quadrangles in each crossroad and between 30 and 60 quadrangles in each street reach (10 across each street multiplied by 3–6 along the street).

The parameters were set to their default values that a priori was considered as the best initial way to obtain suitable results. Friction was estimated to have uniform Manning's coefficient on the whole area accounting for the actual flow friction on the road pavement and the building walls and for the interaction of the flow with various small scale obstacles (trees, cars, bus stops, etc.). The corresponding information available in the literature for 2D simulations in street networks can vary much: Manning coefficients used were $n = 0.01$ (Aronica and Lanza, 2005), $n = 0.015$ (Nania, 1999), $n = 0.025$ (Gourbesville and Savioli, 2002), $n = 0.043$ (Inoue et al., 2000) and $n = 0.08$ (Calenda et al., 2003). Values of friction coefficient $n = 0.025 \text{ s/m}^{1/3}$ and diffusion coefficient of $0.1 \text{ m}^2/\text{s}$ were chosen for our case. The calculation was performed with a fixed time step of 0.01 s. The hydrographs presented in Fig. 3 constitute the upstream boundary conditions at both streets passing below the railway embankment at the northern end of the domain. In order to limit the calculation time, only the initial 10 h of the hydrographs were considered for the simulation. This duration is sufficient to capture the flood dynamics and to obtain the peak water levels at any point in the modeled area. Also, the rain on the domain is neglected. A critical flow boundary condition (Froude number equal to 1) was fixed at the downstream ends of the model, i.e. the six downstream streets (S_3 to S_7 in Fig. 1) oriented from north to south, the three streets (S_8 to S_{10}) oriented west to east and the three (S_1 and S_2 in Fig. 1) oriented east to west. An additional outlet with free flow condition was added in a street at the eastern end of the domain (S_{11}) in order to take into account overflows to a railway line located more than 1 m below the model boundary.

The domain is initially dry and at the beginning of the event, both hydrographs invade the area on each side of the northern part of the domain. Part of the eastern discharge flow turns towards the eastern limit of the area and then leaves the calculation domain through the exit

S_{11} located in *Pitot street*. On the other hand, a large part of both upstream flows joins at the *Faita/Sully* crossroad and spreads in the Central zone narrow streets and into the southern part of the domain before leaving the area at the southern limits. Most of the streets are flooded by the flow from the upstream to the downstream reach of the street, however, some streets such as the western part of *Pitot street* are filled through their downstream reach due to backwater effects. The time of maximum water depth in most places corresponds to the peak of the eastern hydrograph (Fig. 3), which intensity is much larger than the western hydrograph. At the flood peak ($t \sim 4.1 \text{ h}$), high velocities (3–4 m/s) and supercritical regime flow occur along the main north–south slope directions especially in the wide streets and low velocities (0.5–0.7 m/s) with subcritical regimes appear in the streets aligned west to east (Fig. 4). In the crossroads, the flow is generally complex and both regimes can be observed. On the other hand, the maximum depths appear in two locations: in the *Faita* street at the sudden street narrowing and in the *Faita/Sully* crossroad (Fig. 5). Those two locations also correspond to low velocity zones.

Comparison of calculation results with observations

First it appears that the extent of the flooded area (Fig. 5) is in agreement with the limits supplied by the officials of the Gard County as it is mostly determined by topographical limits (hills and building walls).

In Nîmes, the flood marks were recorded on the building walls. However, in our numerical model, no water depth calculation is performed on the cell edges and the 99 flood marks had to be compared with the peak water depths calcu-

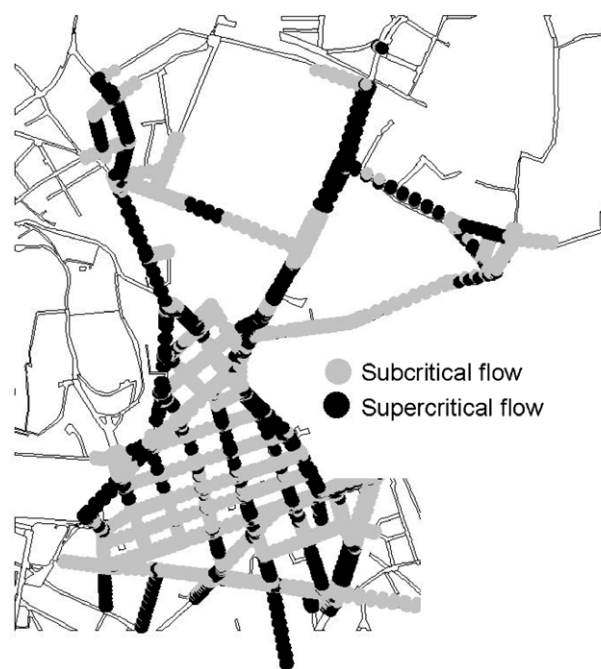


Figure 4 Supercritical and subcritical flow regimes at the peak of the flood.

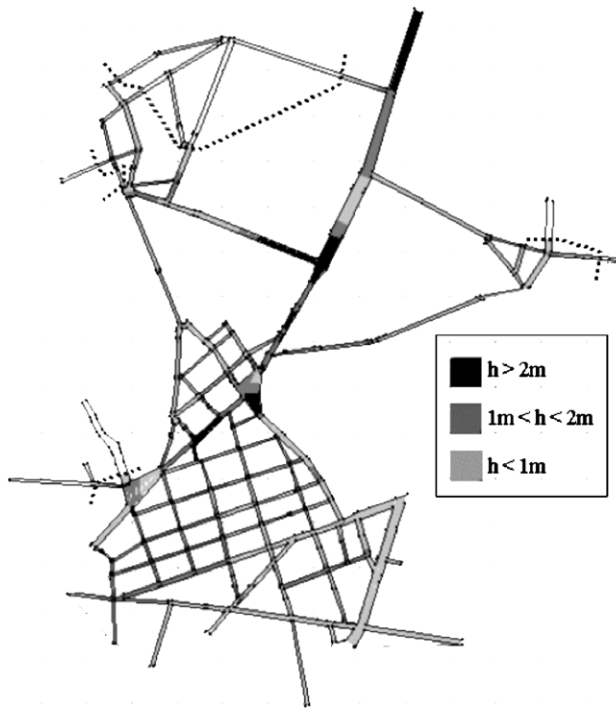


Figure 5 Water depths at the flood peak and limit of flooded area (in dot lines).

lated at the center of the nearest cell, which is on the “pedestrian walk” cell. These cells are named “comparison cells” in the following paragraphs. The results of the peak water depths comparison are given in Table 1. About 40%

of the peak water depths are overestimated and 60% underestimated by the calculation, the largest difference being about 1.6 m and the average absolute difference is 41 cm.

The model appears to be calibrated on the northern part of the domain (Table 1) where the average difference between the computed peak water depths and the flood marks is close to 0, even though the standard deviation is large (56 cm). The peak water depths at two main locations appear not to be fairly estimated: the eastern entry (zone A) where the water depth of the flow below the railway embankment is underestimated and upstream from one of Fanta crossroads (zone B) where the sudden reduction of the street width may not be represented accurately in the interpolated mesh. In the Central part of the domain, the computed peak water depths in the narrow streets are strongly underestimated (zone C) by an average of 43 cm (Table 1). This underestimate may be due to the street characteristics: narrow streets with many parked cars reducing the flowing section and increasing locally the friction. Finally, the peak water depths in the southern area is slightly overestimated ($\overline{dh} = 13$ cm in Table 1); however this overestimate is mainly due to the overestimate of the water depth in one street (zone D in Fig. 6) where topographical uncertainties or other flow features may have occurred.

The histogram (Fig. 7) showing the peak water depth differences between calculation results and on-site measurement is centered around 0 proving that no specific bias is encountered.

Sensitivity test cases

The influence of various parameters was tested on the 1988 flood event by altering just one parameter in each

Table 1 Comparison between computed and measured peak water depth for the reference case and some important parameters modification

	Northern zone	Central zone	Southern zone	Whole domain
Reference case				
\overline{dh} (m)	0.04	−0.43	0.13	−0.13
σ_{dh} (m)	0.56	0.54	0.43	0.53
Case 2B $n = 0.033$				
\overline{dh} (m)	0.17	−0.35	0.17	−0.04
σ_{dh} (m)	0.57	0.54	0.44	0.53
Case 2C $n = 0.025$ and $n = 0.05$				
\overline{dh} (m)	0.06	−0.23	0.14	−0.04
σ_{dh} (m)	0.54	0.46	0.39	0.48
Case 3B S_1, S_2, S_{10} off				
\overline{dh} (m)	0.05	−0.19	0.49	0.05
σ_{dh} (m)	0.54	0.68	0.58	0.61
Case 6B 5-points/section				
\overline{dh} (m)	0.01	−0.36	0.29	−0.08
σ_{dh} (m)	0.54	0.51	0.54	0.53
Case 6C 4-points/section				
\overline{dh} (m)	0.07	−0.45	0.22	−0.11
σ_{dh} (m)	0.64	0.7	0.56	0.65

\overline{dh} is the average difference between computed and measured peak water depths: $\overline{dh} = \frac{\sum dh}{n} = \frac{\sum (H_{\text{computed}} - H_{\text{measured}})}{n}$ with H_{measured} the flood mark measurement, H_{computed} the computed peak water depth in the nearest cell and n is the number of flood marks (99 in this case). σ_{dh} is the standard deviation of dh : $\sigma_{dh} = \sqrt{\frac{\sum (dh - \overline{dh})^2}{n}}$.

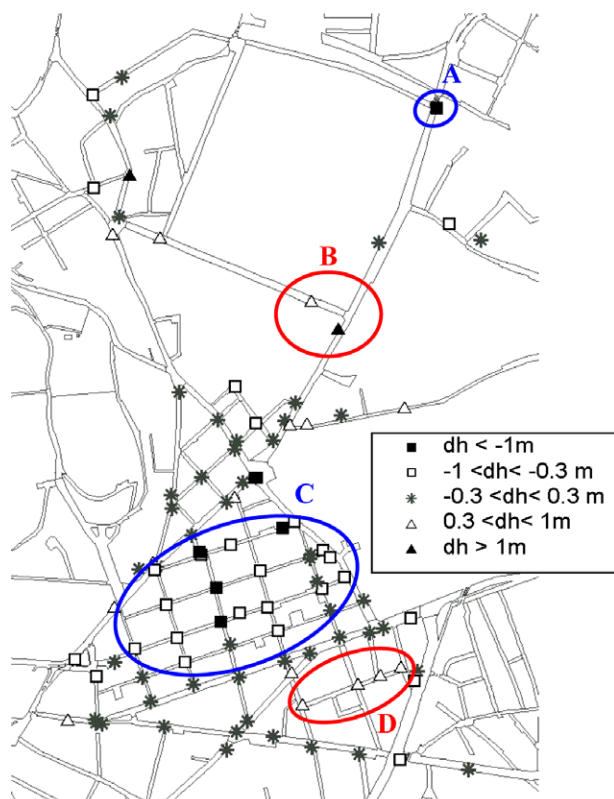


Figure 6 Map of comparison between computed peak water depths and measured flood marks with $dh = H_{\max_{\text{computed}}} - H_{\max_{\text{measured}}}$.

calculation compared to the reference case simulation, by checking the induced modifications in the numerical results and explaining the origins of the changes. From a general

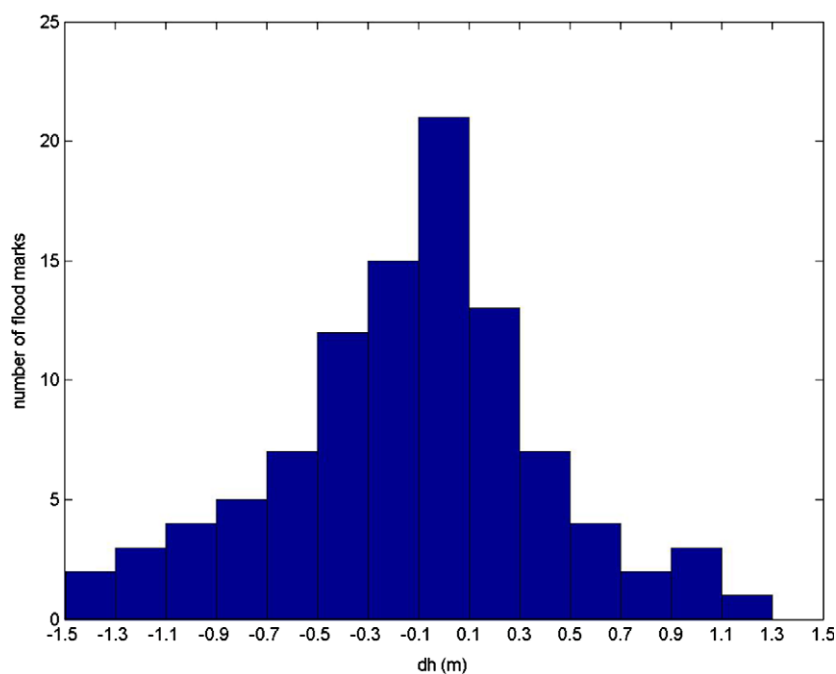


Figure 7 Histogram of peak water depth prediction error.

point of view, Fig. 8 shows that the shapes of the computed limnigrams remain all quite similar but, when looking at the error table (Table 1), it appears that the prediction of the peak water depth at the 99 comparison locations (flood mark locations) can be strongly affected by the choice of the parameters set.

Case 1: Inflow and storage of water

Case 1A. In order to check the influence of the hydrological calculation uncertainties, the input discharge hydrographs were increased by an amount of 20% at each time step compared to the reference calculation. The consequence of such increase in upstream input condition is an average increase of water depth (of 12.5 cm), with a larger increase in the upstream zone (17 cm) than the Central and downstream zones (10 cm each) because of higher damping of peak discharge downstream.

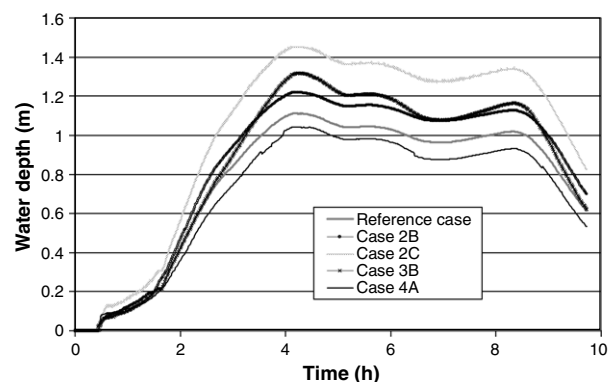


Figure 8 Limnigrams at point P1 (Fig. 1) for various parameter set cases.

Case 1B. In the reference case, the only water input considered is the drainage from the upstream watersheds. In Case 1B, the rain falling on the streets and building blocks (supposed to reach instantaneously the closest street cells) was added. Although the intensity of the rain is very high (maximum of 61 mm/h), the rain water volume remains low (212,000 m³) compared to the input hydrographs volume (3,600,000 m³) and its influence is very limited: the peak water depths are increased everywhere by about 4 cm, i.e. 4%. The strongest influences are observed in the downstream part of the domain (due to cumulative effects) and in the streets filled by backwater effects at the downstream reach (*Pitot street*).

Case 1C. In the reference case, it was assumed that all the buildings were impervious and that water could not be stored in caves, courtyards, dead end streets, etc. However, witnesses and photographs reveal that the water actually invaded most of these areas reducing the volume of water flowing in the streets. Among the storage areas in the domain, the military barracks area represents the largest storage volume available and a calculation with its gate completely open was launched. The computed volume stored at the flood peak is about 80,000 m³ which is very small compared to the input volume of 3,600,000 m³. Consequently, apart from local flow modifications near the gate, only a very small reduction of computed peak water depths (about 1 cm) occurs in the areas downstream from the military barracks.

Case 2: Friction and viscosity numerical coefficients

Case 2A. Energy is dissipated numerically through the bottom friction and diffusion that should mainly reflect both the effect of turbulence and the heterogeneity along the vertical direction. In order to check the influence of the diffusion coefficient ν_{eff} in Eqs. (2) and (3), we simulated the event applying, on the one hand a 0 diffusion coefficient and on the other hand a coefficient depending on the water depth as presented in Eq. (4) with $k = 10^{-2}$. For this latter value 60% of the cells have an equivalent diffusion coefficient ν_{eff} larger than 0.1 m²/s at the flood peak. The results from these two calculations show that the diffusion parameter has a very small influence on the computed flow, the average absolute peak water depth modification being about 3 cm on the studied area.

Case 2B. Identifying the exact friction coefficient is almost impossible as this coefficient stands for the sum of var-

ious sources of resistances such as bottom friction, friction on the walls or due to their irregularities, small scale obstacles to the flows. The $n = 0.025$ Manning coefficient first chosen was changed to $n = 0.033$.

- (1) Overall, this larger friction tends to increase all the computed peak water depths of about 10 cm.
- (2) Nevertheless, it also modifies strongly the flow structure in the vicinity of *Faita/Sully* crossroad with the consequence of altering the flow distribution at this junction. Indeed this larger friction coefficient lowers the flow discharge – thus decreasing the peak water depths – in the downstream *Faita* street and modifies the flow distribution in the downstream boundary streets. Consequently, the computed flood in the whole southern area is affected.
- (3) Finally, the change in friction coefficient slightly alters the predicted flow structure in most crossroads and, in some cells, the flow regime can be changed between supercritical and subcritical conditions, hence strongly modifying the local computed water depths. For instance Fig. 9 presents the flow regimes computed at one crossroad at the flood peak for the reference case and Case 2B and shows that the flow in the comparison cell is supercritical for the reference case and subcritical for Case 2B. Consequently, for Case 2B the water depth at this time is the maximum water depth computed during the event but for the reference case, the water depth is lower than before $t = 3.8$ h and after $t = 4.8$ h (Fig. 10).

In the end, it appears that considering the $n = 0.033$ friction coefficient, the underestimation of peak water depths is decreased in the Central zone but then overestimations appear in the northern zone and increase in the southern zone. Globally, the computed average peak water depth is less underestimated than the reference case (Table 1) but then the model is not any more calibrated in the northern zone ($\bar{d}h = 17$ cm) and becomes more overestimated in the southern zone.

Case 2C. The improvement of the results in the Central zone presented in the previous paragraph leads us to consider different friction coefficients for the various zones of the urban area. The friction coefficient in the narrow streets of the Central zone (all except *Faita* and *Semard* streets) is increased to $n = 0.05$ (and $n = 0.025$ remains in the rest of the domain). Indeed, as the streets are narrower in this zone, the relative importance of the building walls

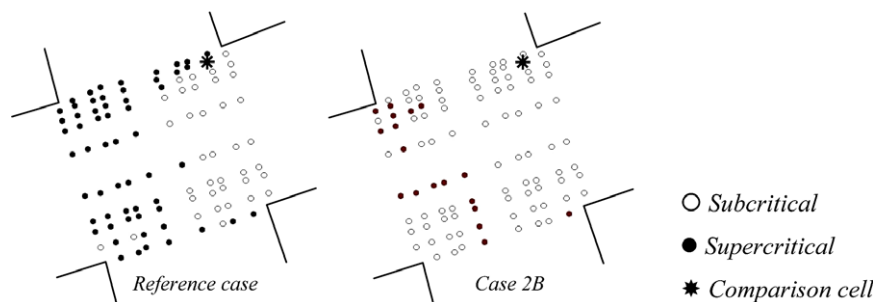


Figure 9 Flow regime patterns at the flood peak ($t = 4.1$ h) in one crossroad for reference case and Case 2B.

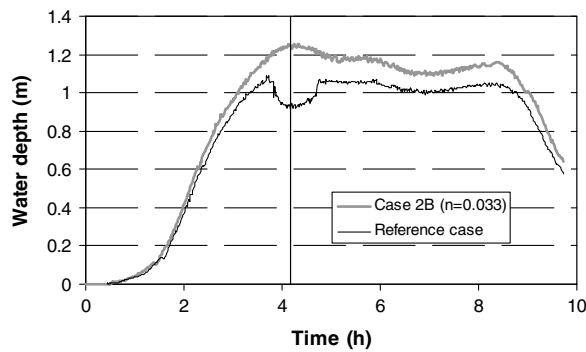


Figure 10 Limnigrams computed at the comparison cell of Fig. 9 for reference case and Case 2B.

friction increases; moreover, the presence of many parked cars creates a stronger resistance to the flow. As was expected, in the Central zone, the velocities are decreased and the peak water depths increased, improving their comparison with the flood marks (Table 1). This test proves that the friction parameter should be calibrated for each homogeneous urban zone according to the potential friction due to the fixed and mobile (non represented) obstacles and the street dimensions.

Case 3: Downstream boundary conditions

Case 3A. The change of the numerical treatment of the downstream boundary condition from a critical flow condition (Froude number = 1) to a free condition (corresponding to a local zero discharge gradient hypothesis in case the flow regime is subcritical, and no flow influence if the local flow is supercritical) strongly alters the flow in the close vicinity of the domain downstream limit: the computed flow depths are strongly increased and the flow distribution to the downstream streets is highly modified.

Case 3B. Besides, it was considered, in the reference case, that the flow could leave the studied domain without being influenced by the further downstream parts of the city. However, a large part of the city was flooded during the event and some backwater effects from the downstream neighbor streets may have occurred especially around the flood peak. Then, a calculation was performed assuming very strong downstream south-eastern and south-western urban zone influences: the flow was prevented from leaving

the domain through streets S_1 , S_2 and S_{10} (Fig. 1). This drastic change of downstream boundary condition has a strong effect on the flow in the whole southern part of the city: the peak water depths are increased in the closed streets (S_1 , S_2 and S_{10}) due to a local storage of water volume and also near the other downstream streets as more outflow discharge occurs. As a result, the computed peak water depth becomes strongly overestimated both in the southern part of the Central zone (due to backwater effects) and in the southern zones (Table 1 and Fig. 11).

To sum up, whereas the choice of numerical method to calculate the downstream boundary condition has only a space influence on the computed flow limited to the last downstream streets, the choice of the streets allowing outflow from the domain to the downstream zones can alter strongly the flow in the whole downstream area up to the fourth street towards the upstream direction.

Case 4: Topographical modifications without altering the mesh

Case 4A. In order to simplify the data processing or if the available topographical data are not precise enough, one can assume the street sections as being flat, thus neglecting the real profile of the streets. In order to check the influence of this simplification, the bottom elevation of each street node was set equal to the corresponding street mid-point bottom elevation. Two tendencies are then observed for the 99 comparison cells:

- (1) Globally, the water depths are increased by about 10 cm in the whole studied area. Indeed, we remark that the water volume in the street is similar to the reference case and generally, on the comparison cells both the water depth and water level are increased (see Fig. 12).
- (2) The topographical modification of the crossroads topography can affect strongly the local flow and the flow regime can change with the consequence of strong local water depth modification (see Fig. 10).

Overall, it appears that the comparison between computed peak water depths and the flood marks is improved in the underestimated zones, namely the Central zone and worsened in the northern and southern zones.

Case 4B. Another important question in urban flood modeling is whether it is worth representing precisely the topog-

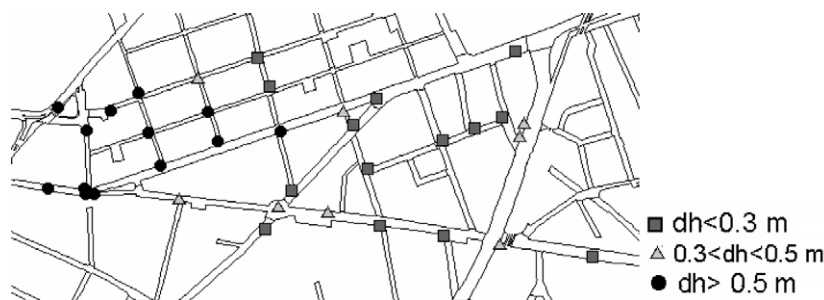


Figure 11 Comparison between the reference case and Case 3B computed water depths in the central and southern zones of the domain with $dh = h_{3B} - h_{reference}$.

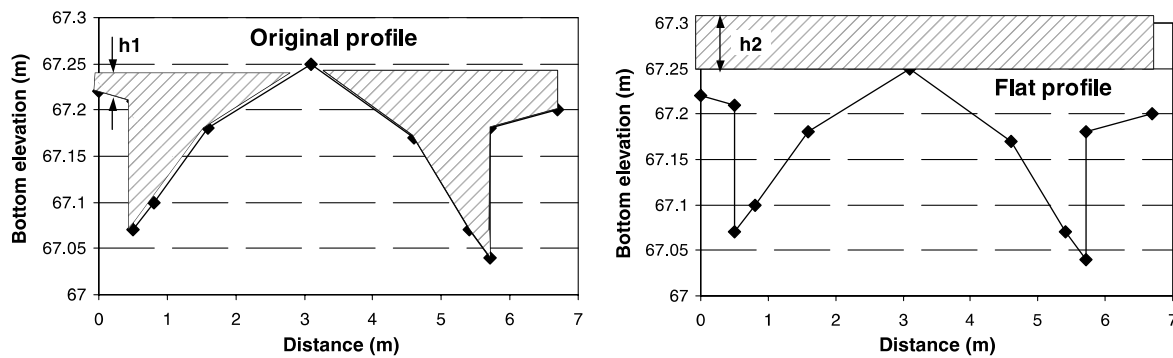


Figure 12 Increase in water depth using a flat profile description.

raphy of the complex crossroads (or squares). A calculation was made by simplifying drastically the topography of the *Faita/Sully* crossroad (to which more than 75% of the flood flow converges) considering that the junction between both streets is a simple quadrilateral area obtained by interpolating both streets sections as for the other simple crossroads. As could be expected the flow structures and thus the computed water depths within the crossroad are much affected. Moreover, it appears that:

- (1) The water depths in the upstream streets are increased as the reduction of the flowing section creates backwater effects towards the upstream area.
- (2) The flow distribution in the downstream streets of the crossroad is modified due to a change of the flow pattern at the junction: the discharge in *Faita* street is increased by 31% whereas the discharge in *Sully* street is decreased by 10%. As a consequence, the peak water depths are globally increased in *Faita* street and slightly decreased in *Sully* street.

Case 4C. Finally, when a large amount of small scale debris (cars, fences, etc.) is available in a zone, some important obstacles can form downstream. In order to check the influence of such an obstacle, an impervious wall was set at the downstream end of a main slope street of the central area (dot line in 1). The following consequences are observed. First, a storage of water just upstream from the obstacle causes a large increase of water depth in the street and backwater effects increasing the water depths in the upstream streets. Also, the original discharge of this street is re-directed towards the neighbor streets, also increasing the water depths. Finally, the water volume downstream from the obstacle is reduced and the computed water depths are lowered. Overall, it appears that the strong influence of the obstacle is spatially restricted, the maximum discharges modification in the downstream boundary streets is only $2 \text{ m}^3/\text{s}$ (i.e., 8%). Nevertheless, for flood events as strong as the 1988 flood, several such obstacles can happen in various places and in the end can have a strong influence on the peak water depths on the whole area.

Case 5: mesh density in a longitudinal direction (same profile description)

Two tests were performed to check the influence of the mesh density in the longitudinal street direction without

altering the mesh density of the streets profile and at the crossroads: the first test (Case 5A) mesh consists on only one calculation cell in each street portion between two crossroads, with the interest of reducing the calculation time. For the second test (Case 5B), the longitudinal mesh comprises an average of 4–5 cells in each street. The comparison between the computed peak water depths and the flood marks located in the streets has no meaning as the calculation points are too different. Concerning the flood marks in the crossroads (85% of the marks), the absolute average peak water depths are almost not affected by the mesh refinement (absolute average change of 2.5 cm) and only slightly affected (by about 8.5 cm) when using the coarser longitudinal mesh. The causes for these modifications are: first, some local flow structure modifications in the junction along with some changes of flow regimes as for Cases 2B and 4A (Fig. 10); then a modification of flow distribution at *Faita/Sully* crossroad, increasing the south-western flow rates and decreasing it in the south-eastern area and finally a strong change in backwater effect in Pitot street flooded by a downstream filling.

Case 6: simplification of street profile mesh

Altering the number of points used in the street profiles leads to the generation of different meshes and to a variation of the profile topography (Fig. 13) seven (Case 6A), five (Case 6B), and four points (Case 6C) were considered. In

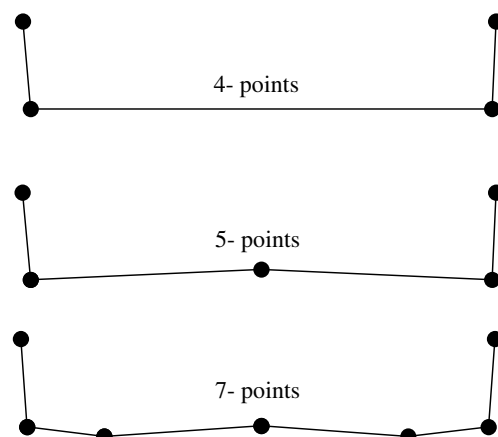


Figure 13 7–5- and 4-points simplified profiles.

going from an 11-points cross-section to a 4-point cross-section, the number of wetted calculation cells at the crossroads decreases from 64 to 1, and the amount of data required to mesh the urban area and the calculation time is also reduced.

For all the simplifications of the street section (Cases 6A, 6B and 6C), a tendency similar as for Case 4A (to a smaller extent) appears: the computed water depths are slightly increased on the whole domain due to an increase of the average bottom elevation of the street section. Moreover, for the 4-points cross-sections case (and also 5-points in a smaller extent), another phenomenon is observed: the small amount of cells in the crossroads simplifies drastically the flow structure. Then the local high or low velocity areas within the junction disappear and only the average flow regime remains. For instance, if the flow in the crossroad is mainly subcritical with a small supercritical area in a corner (where the flood mark is located), then the 4-points cross-section case will compute an average subcritical flow at the junction and thus overestimate the peak water depth (compared to the flood marks and the reference case results). This phenomenon observed at five comparison cells locations is presented in Fig. 14. The opposite phenomenon can also occur if the flow is globally supercritical in the junction with just a small subcritical area. Besides, the local modifications in the computed flow also influence drastically the flow at larger scales and in the 4-points section cases, it alters the flow distribution to the downstream boundary streets as shown in Fig. 15.

To sum up, it appears that the 7-points cross-section computation results show a very small peak water depths modification on the whole studied area whereas these results with 4-points and 5-points cross-section computation are very affected by the street profile mesh simplification (Table 1).

Validation of the model calibration on the 2002 flood event

Apart from the October 1988 flood, the September 2002 event was among the largest flood observed in the past decades in Nîmes. The event process was roughly similar to the

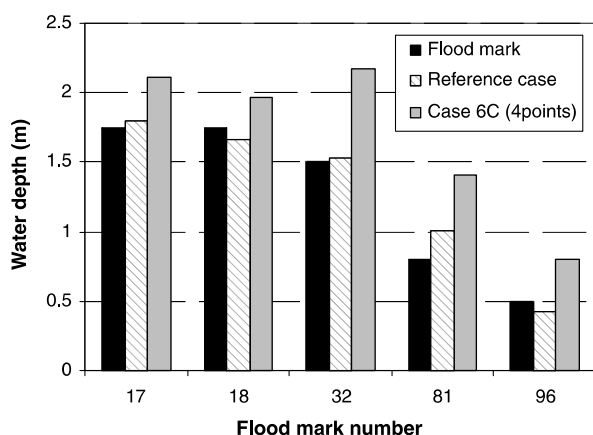


Figure 14 Comparison between flood mark and computed peak water depth for reference and 4-points cross-section calculation at supercritical small scale zones of the crossroads.

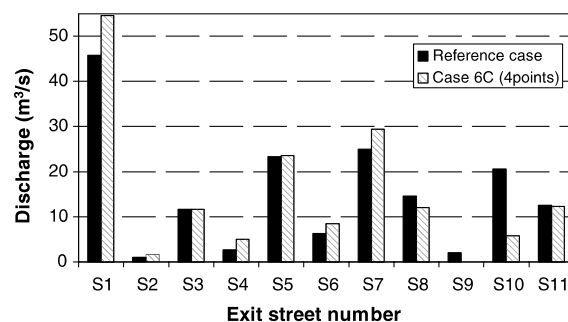


Figure 15 Comparison of flow rate distribution in the downstream end streets for reference and 4-points cases.

1988 case with a rainfall intensity much lower and thus an available volume also reduced. BCEOM company applied a similar hydrological method in order to obtain both input hydrographs. In this case, the 2002 sewer network capacities is of the order of the flow rates coming from the upstream watersheds (6 and 7 m³/s compared to 13 and 30 m³/s) and both pipes capacities were subtracted from the original input hydrographs. Twenty-eight flood marks were measured on building walls after this event, with an average peak water recorded of about 30 cm and a maximum of 85 cm in *Faita* street.

The parameters set used in the 1988 reference case is used to simulate the 2002 event using the same mesh with the appropriate input hydrographs (Fig. 16). It appears from the computed results that the flood dynamics is quite similar as the one described for the 1988 case, even though all water depths are much lower. The comparison between the computed peak water depths and the flood marks is given in Fig. 17.

It appears that the errors in peak water depth computation are quite homogeneously distributed in the studied area, and no spatial tendency can be observed. Indeed, contrary to the former case, the water depths in the narrow streets network of the Central zones are not underestimated. Furthermore, the average error between computed and recorded peak water depths is $\bar{d}_h = -3$ cm proving that the calibration of the model is accurate for the simulation of the September 2002 flood event. However, the large standard deviation ($\sigma_{d_h} = 22$ cm) shows that some local phenomenon must have modified the flow locally. Indeed, with the lower

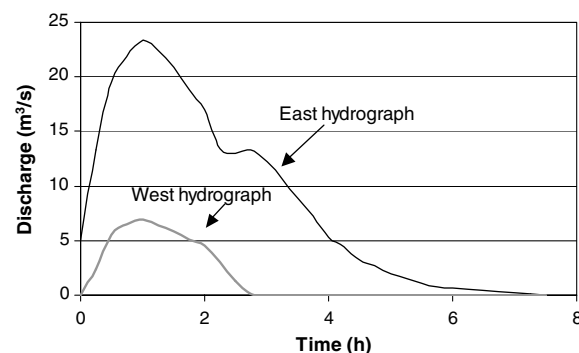


Figure 16 Flood hydrograph for September 2002 flood event after subtracting the sewer capacities.

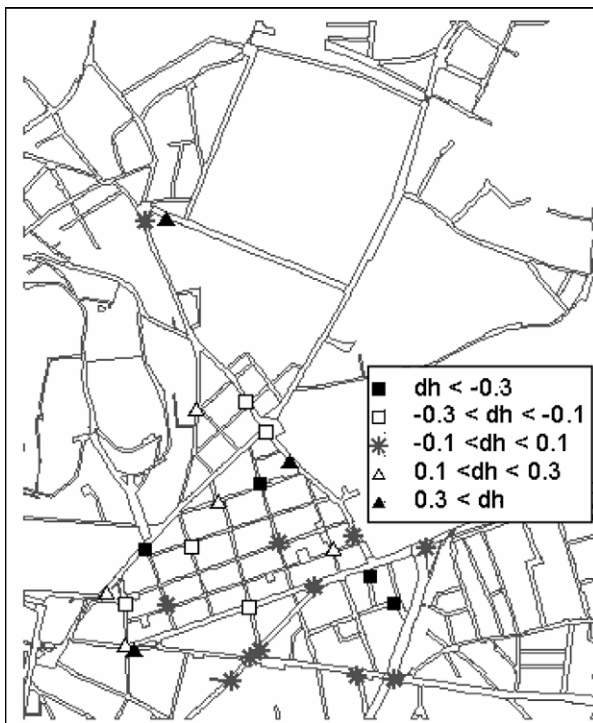


Figure 17 Comparison between peak water depths and measured flood marks.

water depths in the streets, the small scale obstacles, topographical irregularities and storage areas have a larger influence on the flow for this less extreme flood.

Conclusions and recommendations

2D shallow water equations in their complete form were solved to simulate the flood propagation through an urban area represented as a network of streets. Street profiles including its main structural elements were used to mesh the streets and crossroads. Other features such as obstacles (cars, bus stops, etc.), flood water storage spaces (cellars, parkings, etc.) were not considered but could be introduced in order to obtain more realistic local results (Haider, 2001; Haider et al., 2003). A first case of severe urban flood (1988 case) for which the exchanges between the streets and the potential storage areas or the sewage network are globally negligible was used for calibration and in order to check the influence of various numerical and topographical parameters on the computed flow. The numerical results obtained show a standard deviation of about 50 cm. This is rather high but reflects the uncertainty in the flood marks, the insufficiency of the topographical data provided initially, and, may be, the local influence of somewhat hazardous events such as blockages by cars or building walls irregularities for instance. Concerning the sensitivity analysis, it appeared that some parameters had very limit effects and others had larger effects on the flow structure and especially the computed peak water depths. Furthermore, some parameter modifications have global, zonal or local influences. Table 2 sums-up the impacts of the various parameters tested in this article but the conclu-

sions will not necessarily be the same in other flood situations.

Consequently, the following recommendations could be useful for modelers wishing to simulate a severe urban flood in a dense urban area:

- (1) Special attention should be paid during the hydrological calculations to establish the input hydrographs. Especially, the peak discharge of the hydrographs should be precisely calculated if the first objective is to establish the map of the peak water depths on the area.
- (2) If the rain volume falling on the studied area or the storage volume available in the domain are small compared to the input hydrographs volumes, their effects will be very small and it is not worth representing them.
- (3) Different friction coefficients should be applied for the various homogeneous urban feature zone, depending on the width of the streets and the fixed and mobile obstacles that may increase the resistance to the flow.
- (4) Collecting information about the flooded areas just downstream from the studied zone is important in order to set accurate limit boundary conditions at the downstream streets.
- (5) The refinement of the mesh to be used for the calculation depends strongly on the objectives of the study. If the main objective is to have an overview of the flood dynamics, then a rough description of the street profile seems accurate enough. In the case of the 1988 flood, the 5-points cross-section points with three wetted cells longitudinally between two crossroads seem a good compromise. Nevertheless in this case, the modeler should keep in mind that a slight global water depth overestimation (about 10 cm) occurs on the whole studied domain due to the increased average street profile bottom elevation. The main advantages of considering a coarse mesh is both to simplify the topographical data collection on the area and to reduce the calculation time, often prohibitive when trying to simulate precisely an urban flood. However, if information about local velocities and/or water depths exists in order to use this data for model calibration, then a more precise description of the streets is required. In this case, it seems that the representation of the gutters is not required and the use of 7-points crossing-sections (four wetted cells in the section), thus with 49 nodes and 32 wetted cells in each crossroad should be accurate enough except in the complex crossroads where specific attention have to be paid to mesh the local topography more precisely.

Finally, the calibrated model was validated on a second less extreme event: the September 2002 flood. The results show a standard deviation of about 20 cm, which is better than for the 1988 flood case but is high compared to the average peak water depth in the area (about 30 cm). One reason for this may be the higher sensitivity of the model to the small scale irregularities of the topography.

These results on one particular flood case should be confirmed on other field or experimental cases showing a similar complicated pattern of streets and crossings. However, it was shown that the use of a code solving 2D shallow water equations to assess flood risk in a city seems a convenient

Table 2 Impact of the main parameters in the sensitivity analysis with: (1) global change of peak water depths (with ↑ = increase and ↓ = decrease of peak water depth); (2) change in flow distribution at Faita/Sully crossroad and to the downstream limit areas; (3) local flow regime change

Parameter modification	(1)	(2)	(3)	Other impacts and remarks
Increase input hydrographs <i>Case 1A</i>	X (↑)			Average increase of 12.5 cm if the flow rates are increased by 20%
Increase friction coefficient <i>Case 2B</i>	X (↑)	X	X	
Friction coefficient defined by zone <i>Case 2C</i>				The calibrated coefficient value improves the results in every zone
Change of downstream boundary condition numerical treatment <i>Case 3A</i>				Modification of discharge distribution in the downstream streets
Backwater effects from downstream zones: blocking some exit streets. <i>Case 3B</i>				Global flow modifications in the whole southern area
Flat street profile <i>Case 4A</i>	X (↑)		X	Similar increase of water depth in the three zones compared to reference case
Simplified main crossroad <i>Case 4B</i>		X		Backwater effects upstream from the crossroad and flow pattern modifications
Local obstacle <i>Case 4C</i>				Strong local effects and slight downstream consequences
Simplified longitudinal mesh <i>Case 5A</i>		X	X	Modification of backwater effects in downstream filling zones
Simplified street profile (4- and 5-points cross-sections) <i>Cases 6B and 6C</i>	X (↑)	X	X	Simplification of the flow patterns in the street junctions
Considering the rain <i>Case 1B</i>	X (↑)			Limited impact due to small rain volumes compared to input hydrographs
Considering the storage areas <i>Case 1C</i>	X (↓)			Limited impact due to small volumes stored compared to input hydrographs
Diffusion coefficient <i>Case 2A</i>			X	Limited impact
Neglecting gutters (7-points cross-sections) <i>Case 6A</i>	X (↑)			Limited impact

solution. In the case of Nîmes, the results of such calculations are being used for the reorganization of buildings in the Richelieu area for urban projects ("Hoche – Sernam" operations). The new urban planning will be taken into account to modify the mesh of the studied area and calculations with the calibrated model of a flood event equivalent to the 1988 case will permit to check the peak water depths modifications.

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