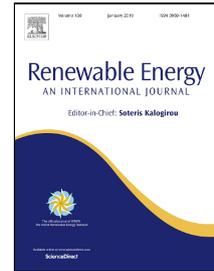


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Assessment of Surface Wind Datasets for Estimating Offshore Wind Energy along the Central California Coast

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1 **Assessment of Surface Wind Datasets for Estimating Offshore Wind Energy along the**  
2 **Central California Coast**

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15  
16 **Highlights**

17 Comprehensive assessment of near-surface wind datasets for Central California

18 Seasonal and diurnal error metrics and spatial differences analyzed

19 Framework for assessing wind data products for offshore wind energy

20

21

22

## 23 **Abstract**

24 In the United States, Central California has gained significant interest in offshore wind energy due  
25 to its strong winds and proximity to existing grid connections. This study provides a  
26 comprehensive evaluation of near-surface wind datasets in this region, including satellite-based  
27 observations (QuikSCAT, ASCAT, and CCMP V2.0), reanalysis (NARR and MERRA), and  
28 regional atmospheric models (WRF and WIND Toolkit). This work highlights spatiotemporal  
29 variations in the performance of the respective datasets in relation to in-situ buoy measurements  
30 using error metrics over both seasonal and diurnal time scales. The two scatterometers (QuikSCAT  
31 and ASCAT) showed the best overall performance, albeit with significantly less spatial and  
32 temporal resolution relative to other datasets. These datasets only slightly outperformed the next  
33 best dataset (WIND Toolkit), which has significantly greater temporal and spatial resolution as  
34 well as estimates of winds aloft. Considering tradeoffs between spatiotemporal resolution of the  
35 underlying datasets, error metrics relative to in-situ measurements, and the availability of data  
36 aloft, the WIND Toolkit appears to be the best dataset for this region. The framework and tradeoff  
37 analysis this research developed and demonstrated to assess offshore wind datasets can be applied  
38 in other regions where offshore wind energy is being considered.

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43 **Keywords:** Offshore Wind Energy, Scatterometers, Reanalyses, Regional Atmospheric Models,  
44 Surface Winds, Tradeoff Analysis

45

## 46        **1. Introduction**

47            Over the last few decades, renewable energy sources have become an increasingly  
48 important component of broader energy portfolios. Costs of renewable energy have decreased  
49 substantially, and more governments recognize the importance of reducing greenhouse gas  
50 emissions. As a result, governments at many levels have set targets for increasing renewable  
51 energy generation. For example, across the European Union, the European Parliament and Council  
52 has set a target of 20% for energy consumption from renewables by the year 2020 (2020 Climate  
53 & Energy Package). Additionally, many states within the United States have adopted increased  
54 renewable energy portfolio targets. This includes California, which has set a goal to supply 50%  
55 of energy through renewable sources by the year 2030 (SB350-Clean Energy and Pollution  
56 Reduction Act of 2015).

57            In response to governmental initiatives and decreases in costs, deployment of renewable  
58 energy projects has been increasing rapidly, with an emphasis on photovoltaic solar and land-based  
59 wind turbines [1]. Offshore wind turbines also have received considerable interest and investment,  
60 particularly in Europe [2]. Offshore wind energy has several advantages over solar and land-based  
61 energy sources since offshore winds tend to be stronger and more consistent than land-based winds  
62 [3] and are less likely to directly conflict with other land-use activities. Additionally, offshore wind  
63 energy production may be able to reduce discrepancies in production and demand that are difficult  
64 to alleviate with solar output because of its diurnal cycle.

65            To best guide the evaluation and planning of offshore wind energy in a particular area,  
66 accurate wind datasets with sufficient temporal and spatial resolution are needed. Offshore winds  
67 typically exhibit temporal variability on interannual, seasonal, synoptic, and diurnal time scales.  
68 Furthermore, wind power is proportional to the cube of the wind speed, meaning that small changes

69 in wind speed (e.g., over the course of the day or with different seasons) can lead to drastic  
70 differences in power output. Also, for power generation to be most valuable, it will need to match  
71 grid demands and base load needs, which vary daily and seasonally. Thus, wind datasets long  
72 enough to capture interannual variability and with sufficient temporal resolution to resolve diurnal  
73 variability are required for estimating wind energy power production and value. In addition,  
74 understanding spatial variation in offshore wind power can help support site planning and  
75 assessment by highlighting areas with the greatest potential to generate power and therefore areas  
76 with the greatest potential value. Despite the importance of understanding temporal and spatial  
77 variations in offshore winds for assessing this renewable energy resource, previous work has rarely  
78 resolved both daily and seasonally cycles at multiple sites and/or over a large area. Moreover, the  
79 utilization of temporally-averaged (mean) wind speeds over an annual cycle can lead to large errors  
80 and mismatches in grid demand and production estimates over shorter (seasonal and daily) time  
81 scales.

82         The lack of detailed assessments across a range of time scales and over broad spatial  
83 domains is mainly attributable to the absence of a single perfect offshore wind dataset with the  
84 appropriate temporal and spatial resolution. In-situ near-surface wind measurements from moored  
85 buoys are often available over long time periods (decades) with a very high temporal resolution  
86 (hourly or better), but these buoys are usually sparse (often >10-100 km apart). Remote sensing  
87 measurements of near-surface winds obtained from satellites equipped with scatterometers can  
88 measure vector wind fields across large areas that are more spatially resolved than buoy platforms,  
89 but the measurements are only available during satellite passes, at most several times per day [4].  
90 Reanalysis products, which objectively combine both observations and numerical models, often  
91 have consistent temporal resolution over decades and contain winds at various vertical levels above

92 the surface, but have coarser spatial resolution compared to satellite-derived data. Finally, regional  
93 atmospheric models have some of the highest spatial and temporal resolution, including data aloft  
94 at various vertical levels; however, they often experience substantial error relative to in-situ  
95 observations and are sensitive to local parameterizations [5].

96 Previous studies have evaluated the performance of various wind datasets in different  
97 regions (see Carvalho et al. [3] and the references therein). Pickett et al. [6] and Tang et al. [7]  
98 assessed the performance of QuickSCAT satellite observations relative to local buoys along the  
99 West Coast of the United States, but they did not assess other datasets. Carvalho et al. [8]  
100 conducted a comprehensive comparison of satellite-based observations, reanalysis products, and  
101 the Weather Research and Forecast (WRF) regional model with five buoys in the Iberian Peninsula  
102 coast. Carvalho et al. [3] extended the analyses of Carvalho et al. [8] by including newer  
103 scatterometers (e.g. ASCAT). However, these studies focused on error metrics over one year and  
104 did not consider longer time periods or seasonal and diurnal variability. Alvarez et al. [9] used a  
105 longer time period (10 years) to evaluate satellite-based products and reanalysis products against  
106 in-situ buoy measurements in the southern Bay of Biscay. They found that QuikSCAT had the  
107 lowest bias in wind speed and wind direction and the Cross-Calibrated Multi-Platform (CCMP,  
108 blended satellite product) had the lowest error, but they did not include an analysis of the diurnal  
109 signal.

110 Collectively, these studies and others (see Carvalho et al. [3] and the references therein),  
111 also suggest that the performance of different wind products varies by study region, indicating the  
112 need for site-specific analyses. The majority of site-specific evaluations of offshore wind data have  
113 focused on coastal waters along Europe, typically in association with existing or planned offshore  
114 wind farms (e.g. [3,8–10]). To date, all but one of the world's offshore wind farms in operation

115 consist of fixed-bottom wind turbines located in shallow waters of less than 100 m. Yet, as  
116 technology advances, the cost of building floating wind turbines in water greater than 100 meters  
117 deep may be less than that of fixed-bottom platforms by 2030 [11]. The first MW-scale floating  
118 turbine was successfully deployed in the North Sea in 2009 [45]. In 2017, the world's first floating  
119 offshore wind turbines were successfully launched with the Hywind project in Scotland, paving  
120 the way for future wind farms in deeper waters further from the coast [46]. With improvements in  
121 floating turbine technology, deployment of offshore wind farms is likely to increase in the future,  
122 particularly in areas with deeper shelf waters. Understanding wind patterns (both spatially and  
123 temporally) in these environments will be key to guiding and assessing marine renewable energy  
124 production.

125         Along the West Coast of the United States the continental shelf is narrow, such that waters  
126 are often >100 m deep only a few kilometers from shore. As a result, the majority of the ocean  
127 area with the potential for wind power production is located in deep waters where floating turbines  
128 would be necessary [12]. The Central California region considered in this study, spanning from  
129 south of Monterey Bay to Point Conception is characterized by moderately strong winds  
130 throughout the year (e.g., [13]; Figure 1). Additionally, this region is located in the vicinity of  
131 several existing connections to the state's electrical grid, including the Morro Bay power plant  
132 (closed in 2014) and the Diablo Canyon nuclear power plant (California's last remaining nuclear  
133 power plant slated to close in 2025). Finally, the study domain is outside of National Marine  
134 Sanctuary areas, where restrictions on disturbance to the seabed will likely preclude floating  
135 turbine deployment. Attracted by these features, private industry has shown great interest in  
136 pursuing permits from government agencies for the development of deep water, floating offshore  
137 wind farms (BOEM: <https://www.boem.gov/California/>). Therefore, a detailed analysis of the

138 available wind products in this region is needed. However, aside from a few simple analyses of  
139 winds (e.g., [14]), there are no comprehensive assessments of long-range, high-resolution wind  
140 products in this region. Without this information, it is difficult to accurately evaluate the power  
141 production potential of this region.

142 To address this knowledge gap, we conducted a comprehensive evaluation of near-surface  
143 winds from various datasets (satellite-based, model, reanalysis) and compared them to local buoy  
144 measurements. We used these datasets, which span nearly a decade and with up to 2-km spatial  
145 resolution, to assess error metrics (bias and root-mean-square-error) over seasonal and diurnal time  
146 scales. Using the results of these point-to-point comparisons, and consideration of the  
147 spatiotemporal resolution of each dataset and whether it provides data aloft, we examined tradeoffs  
148 between various dataset attributes (e.g. bias, error, spatial and temporal resolution, availability of  
149 data aloft) to identify the best dataset for offshore wind energy application. We then explored  
150 characteristics of the chosen dataset to reveal temporal changes in near-surface wind speeds across  
151 the domain along the Central Coast of California. The framework we developed to evaluate the  
152 various products is readily applicable to other regions where similar analyses are needed, and the  
153 wind dynamics we reveal for the Central Coast can be used to support the generation of accurate  
154 and detailed estimates of potential power production in the region.

155

## 156 **2. Data and Methods**

### 157 **2.1 Study Domain**

158 The Central Coast of California is located along the eastern boundary of the Pacific Ocean  
159 and features steeply sloping bathymetry. In this study, we considered the domain bounded by the

160 Monterey Bay National Marine Sanctuary to the north, the Channel Islands National Marine  
161 Sanctuary to the south, and the 1000 m isobath in the offshore direction, generally west (Figure 1).  
162 The offshore limit is the maximum water depth for offshore wind turbine installation based on  
163 current technology and industry experience [12]. Along this stretch of coastline, there are three  
164 existing connections to the state's electrical grid: the Morro Bay power plant (closed in 2014), the  
165 Diablo Canyon nuclear plant (slated to close in 2025), and Vandenberg Air Force Base. This region  
166 is characterized by moderately strong and consistently equatorward winds throughout much of the  
167 year, particularly for the region north of Point Conception (e.g., [13,15]). A previous study  
168 suggested that the annual average of wind speed at hub height exceeds  $7 \text{ m s}^{-1}$ , highlighting the  
169 potential for offshore wind farms [12].

## 170 **FIGURE 1 LOCATION**

171

## 172 **2.2 Wind Datasets**

### 173 **2.2.1 Buoy observations**

174 Near-surface winds in this study domain were obtained from moored buoys measuring  
175 winds at 5 m above the surface and reporting an average wind speed every ten minutes (i.e., the  
176 National Data Buoy Center (NDBC) continuous wind product, <http://www.ndbc.noaa.gov/>). We  
177 employed buoy data as a reference to represent true characteristics of near-surface winds, as is  
178 commonly done in the existing literature (e.g., [16]). While buoy measurements are the best  
179 available in-situ data, buoy measurements may be less reliable under strong winds [17], but these  
180 measurements are still likely the best estimates of true wind speeds. Among all datasets considered,  
181 the buoy dataset is the only to output near-surface winds at 5 m above the sea surface, with the

182 other datasets outputting near-surface winds at 10 m above the sea surface. Thus, to enable a direct  
183 comparison, we converted the buoy-measured wind speeds from 5 m to 10 m assuming a neutrally  
184 stable atmosphere following the method of Liu and Tang [18]. This is a reasonable assumption  
185 given that calculated atmospheric stabilities show a neutrally stable atmosphere during most  
186 seasons and hours of the day. Potential errors in 10-m winds speeds when atmospheric stability  
187 deviates slightly from neutral conditions are expected to be small [19]. Buoys 46028 and 46011  
188 are located north of Point Conception, and buoy 46054 is located just to the south of Point  
189 Conception, at the western edge of the Santa Barbara Channel (red dots in Figure 1).

190

### 191 **2.2.2 Satellite-based observations**

192 We evaluated two scatterometers, which measure surface wind stress by sending  
193 microwave signals and then recording the backscattered signal in response to ocean roughness  
194 (e.g. [18]). Surface wind stress is converted to equivalent neutral winds 10 m above grounds based  
195 on the assumption of a nearly neutral atmosphere [18]. Vector wind fields are produced at  
196 approximately the same geographical location during ascending and descending passes of the  
197 satellite (i.e., twice per day). Here, we opted to use the swath data with 12.5 km spatial resolution  
198 because this high resolution product can contain small-scale features [20]. We downloaded both  
199 scatterometer-derived datasets from the NASA's Jet Propulsion Laboratory Physical  
200 Oceanography Distributed Active Archive Center site (<https://podaac.jpl.nasa.gov>).

201 The first scatterometer dataset we evaluated was QuikSCAT, which measures the  
202 backscattered signal using the Ku-band frequency and passes through our study domain around 5  
203 and 18 h every day. QuikSCAT data were available from June 1999 to November 2009. This  
204 widely-used product has been validated for accuracy against in-situ buoy observations over various

205 forcing regimes (e.g., [21]). We adopted the latest version of the Level 2 product (QuikSCAT  
206 Level 2B Version 3), which uses the improved geophysical model function and corrected rain  
207 contaminated wind speeds with a neural network approach [22].

208 The second scatterometer dataset we evaluated was ASCAT, which is a new-generation  
209 scatterometer launched in October 2006. It agrees well with QuikSCAT especially when wind  
210 speeds range between  $3 \text{ m s}^{-1}$  and  $20 \text{ m s}^{-1}$  [23]. ASCAT passes a local point around 9 and 20 h  
211 and uses the C-band frequency operation, which is less sensitive to rain contamination than the  
212 Ku-band frequency operation [24]. Because of its narrower swath width, ASCAT is limited to  
213 approximately 60% of the coverage of QuikSCAT during the same period [25]. The ASCAT Level  
214 2-Coastal product applies a boxcar filtering to yield more wind data close to the coast [26]. We  
215 used the Level 2 product's Climate Data Record version, which was reprocessed using consistent  
216 calibration from January 2007 to March 2014.

217 The last satellite-based product we assessed was the Cross-Calibrated Multi-Platform  
218 Version 2 (CCMP V2.0, a continuation of CCMP Version 1.1) [27]. We obtained this dataset from  
219 Remote Sensing Systems (<http://www.remss.com/>). This blended product combines satellite-  
220 derived wind fields from microwave radiometers and scatterometers, with moored buoys and  
221 ERA-Interim model data using a Variational Analysis Method. It provides global and gap-free  
222 wind fields on a  $0.25^\circ$  grid four times per day from 1987 to the present. Previous studies in the  
223 southern Bay of Biscay [9] and the Iberian Peninsula coast [3, 8] demonstrated that CCMP  
224 accurately captured offshore winds.

225

### 226 **2.2.3 Reanalysis datasets**

227 We also assessed two reanalysis products, which combine in-situ observations with  
228 numerical models: 1) Modern-Era Retrospective Analysis (MERRA,  
229 <http://disc.sci.gsfc.nasa.gov/mdisc/>), a global reanalysis product [28], and 2) North American  
230 Regional Reanalysis (NARR, <https://www.esrl.noaa.gov/psd/>), a regional reanalysis product [29].  
231 MERRA is a commonly-used global reanalysis product for wind resource evaluations (e.g., [30]).  
232 It provides hourly data on a grid of  $2/3^\circ$  by  $1/2^\circ$  from 1979 to 2016. NARR outputs data every three  
233 hours (since 1979) and has a spatial resolution of 32 km. Both products yield wind data at various  
234 pressure levels above the surface. In part because it assimilates more observations into its model,  
235 NARR data yield more accurate results relative to global reanalysis products [29]. Previous studies  
236 in other regions have also shown good agreement between NARR and in-situ measurements near  
237 the surface and aloft (e.g., [31–33]).

238

#### 239 **2.2.4 Regional atmospheric model simulations**

240 We analyzed simulated near-surface wind speeds from two regional model datasets. The  
241 first dataset covers the entire U.S. West Coast and was carried out using WRF model version 3.6  
242 [34,35], which is initialized and forced at boundaries with the Climate Forecast System Reanalysis.  
243 The model is configured with two nested grids, where the outer domain has a horizontal resolution  
244 of 18 km, and the inner domain has a resolution of 6 km. It is set up with a full set of  
245 parameterization schemes including the Mellor-Yamada-Nakanishi-Niino planetary boundary  
246 layer scheme [36], which is one of the best planetary boundary layer schemes to simulate realistic  
247 cloud cover and wind. More details can be found in Renault et al. [35]. Hourly 10-m wind fields  
248 above the ground level are available from 2004 to 2013 and used for this study.

249 The second regional model dataset is from the WIND Toolkit  
250 (<https://www.nrel.gov/grid/wind-toolkit.html>), developed by the National Renewable Energy  
251 Laboratory (NREL) for the purpose of wind power application [30]. The results were generated  
252 by the WRF model version 3.4.1, which is initialized and forced at boundaries by the European  
253 Center for Medium-Range Weather Forecasts Interim Reanalysis. This model uses three nested  
254 grids with resolutions of 18 km, 6 km, and 2 km, respectively, with the inner 2 km grid covering  
255 the entire contiguous United States. The optimal model configuration is the best one from the eight  
256 model configurations tested by NREL. This configuration outputs simulations with small overall  
257 bias and in complex terrain, realistic hourly and diurnal wind variations, and highly resolved wind  
258 fields near the surface. More details can be found in Draxl et al. [30,37]. We analyzed hourly 10-  
259 m wind fields (2007-2013). In addition to near-surface wind fields, wind data at higher altitudes  
260 up to 200 m are also available.

261

### 262 **2.3 Comparisons and Statistics**

263 In order to compare the various datasets to the buoy observations, we obtained the closest  
264 point in space and time from each wind dataset relative to each of the three buoys. We included  
265 observations only if they met our collocation criteria with buoy data: measurements must have  
266 been recorded within 30 minutes of a buoy measurement and no more than 12.5-km from the buoy  
267 for all datasets except WIND Toolkit. We use a more restrictive spatial criterion of 2 km for WIND  
268 Toolkit because of its higher resolution of 2 km. Unlike gridded datasets, the closest swath point  
269 of the scatterometer data to a local buoy is not fixed and its measurement time is slightly different  
270 each day. In line with previous studies (e.g., [6]), we found no connection between the separation  
271 distance and the bias in QuikSCAT/ASCAT relative to a local buoy. Between 2000 and 2008 (time

272 period used for comparison in this study), the mean separation distance between the closest  
273 QuikSCAT point and buoy was 5.59, 4.45, and 4.93 km for buoy sites 46028, 46011, and 46054,  
274 respectively. Between 2007 and 2013, the mean separation distance between the closest ASCAT  
275 point and buoy was 4.91, 3.55, and 3.77 km. The distance between a local buoy and the closest  
276 point of a comparative gridded dataset is shown in Table 2.

277 We evaluated the seven aforementioned wind datasets in relation to buoy measurements  
278 using the collocation criteria described above. To summarize the performance of each wind data,  
279 we utilized the statistical metrics of the bias and the root-mean-square-error (RMSE) between one  
280 dataset and buoy measurements. To illustrate the relationship between two variables, we fitted the  
281 paired data to a linear regression line and provided its intercept, its slope, and the coefficient of  
282 determination ( $R^2$ ) of the model fit in Tables 2, 3 and 4.

283 Complete annual data were available for at least 7 years for all datasets (see Table 1 for  
284 details), thereby reducing the impact of interannual variability on our analysis. To display  
285 climatological characteristics of near-surface winds, we used buoy data from 1998-2016 and  
286 compared these winds between the buoys and other datasets for each year of overlap.

287

## 288 **2.4 Tradeoff Analysis**

289 In order to evaluate the relative merits of the datasets and identify the best dataset for  
290 offshore wind power applications, we applied a tradeoff analysis to our results. Tradeoff analysis  
291 is a useful graphical tool for comparing the relative performance of a set of options in relation to  
292 multiple objectives [38]. We considered five key objectives, or factors, in the tradeoff analysis of  
293 the seven wind datasets: temporal and spatial resolution (higher better), the absolute value of bias

294 and RMSE (lower better), and availability of wind speed data aloft (better). We then conducted  
295 visual inspection of pairwise tradeoff plots of the seven datasets in relation to the five factors in  
296 order to compare and contrast the relative merits of the datasets and identify the most appropriate  
297 one(s) for offshore wind power applications.

298

### 299 **3. Results**

#### 300 **3.1. Buoy climatology**

301 Climatological characteristics of buoy winds are shown in Figure 2. Each curve represents  
302 the composite day average wind speed for a particular month (i.e., average wind speed calculated  
303 using data over all years from a particular hour during each month). All times referenced are  
304 Pacific Standard Time (PST).

#### 305 **FIGURE 2 LOCATION**

306 The three different buoy sites display similar diurnal structure with daily minimums in the  
307 late morning and peaks in the early evening (Figure 2). There is also a slight seasonality in both  
308 the timing of the daily minimums and peaks, as well as the daily range. During months with  
309 stronger wind forcing (e.g., spring/summer upwelling months, cf. Walter et al. [15]), the daily  
310 peaks arrived slightly later compared to during other months. For example, at 46028, wind speed  
311 peaks around 20 h in May and at 18 h in January. Notably, the diurnal variability is comparable to  
312 that of the seasonal variability. There is also considerable buoy-to-buoy (i.e., spatial) variability at  
313 various time scales. Among the three sites, buoy 46054 displayed the strongest diurnal variations  
314 in wind speed with differences as large as  $3 \text{ m s}^{-1}$  between the daily minimum and maximum in  
315 some months.

316 Seasonal cycles also varied among buoys. The 10-m wind speeds at buoys 46028 and  
317 46011 reached their maxima in spring, whereas the 10-m wind speed at buoy 46054 reached its  
318 maximum in the summer (see Walter et al. [15] for a discussion of the seasonality at buoy 46011).  
319 This seasonal variation is closely connected to large-scale pressure systems, which fluctuate  
320 seasonally, but tend to produce equatorward winds near the surface along the coastline (see  
321 Fewings et al. [13] for a detailed description). Among the three sites, buoy 46054 had the strongest  
322 and most variable winds, which is strongly impacted by the interaction between the marine  
323 boundary layer and coastal capes (i.e., Point Conception) [13].

324

### 325 **3.2 Paired comparisons with buoy measurements**

326 Direct comparison between the wind speed calculated from each buoy at each site and each  
327 respective dataset were made using all data available over the selected time period for all points in  
328 each dataset that met collocation criteria. Figure 3 shows scatter plots and the linear regression line  
329 between each wind product's wind speed and the buoy site's wind speed. Statistics from the linear  
330 regression and collocation criteria are shown in Table 2. The error metrics (bias and RMSE) are  
331 displayed in Table 3.

### 332 **FIGURE 3 LOCATION**

333 Based on the performance of the linear regression (Table 2, Figure 3) and the error metrics  
334 (Table 3), ASCAT had the lowest bias and RMSE, and the largest coefficient of regression with  
335 buoy-based site measurements, slightly outperforming the other scatterometer-based observation,  
336 QuikSCAT. This is not surprising, given previous validations of the product in other regions (e.g.,  
337 [3]). We note, however, that the scatterometers (particularly ASCAT) have the smallest number of

338 points used for comparison with the buoy data because of the temporal resolution (typically only  
339 two measurements per day) and a shorter time period relative to other datasets. Following  
340 scatterometer-based observations, the WIND Toolkit showed the best correspondence with buoy  
341 data; this dataset even outperformed the scatterometers slightly with respect to bias at buoys 46011  
342 and 46054 and had relatively low error as well (Table 3). We note that the WIND Toolkit is also  
343 the most spatially (2 km) and temporally (1 hr) resolved dataset, and it contains wind data at  
344 various levels about the sea surface. While the WIND Toolkit, a version of the WRF regional  
345 model, displayed some of the best results, the other WRF model considered (denoted WRF here,  
346 a model developed for the West Coast of the United States) showed the worst correspondence to  
347 local buoys in this region. Given the sensitivity of the performance of the WRF model in wind  
348 simulation to various configurations and parameterizations, (e.g., [5]), it is possible that the better  
349 performance of the WIND Toolkit than its counterparts is associated with its configuration  
350 particularly optimized for simulating wind for wind energy applications. The largest error (RMSE)  
351 among the three sites is generally found at buoy 46054, which is located just south of Point  
352 Conception, highlighting the difficulty of resolving the wind field near complex land topography.  
353 Among these seven wind datasets, five (ASCAT, QuikSCAT, CCMP, NARR, and MERRA)  
354 display the worst correspondence at buoy 46054, while model simulations (WRF and WIND  
355 Toolkit) show relatively consistent correspondence across all buoys. The greater biases in the  
356 reanalysis datasets at buoy 46054 are likely due to their coarser spatial resolution, which is not  
357 able to accurately capture small-scale coastal orography near Point Conception and its impact on  
358 the velocity field.

359 To further investigate the differences between the various datasets and the local buoys, we  
360 examined the wind speed difference between a particular dataset and the local buoy as a function

361 of the buoy wind speed (Figure 4). In general, data products overestimated winds relative to the  
362 buoy at low wind speeds and underestimated at high wind speeds, with varying degrees of  
363 magnitude. This feature and negative relationship is consistent with the findings of previous  
364 studies using less than two years of data [3,6,7]. Statistics of the linear regression between wind  
365 speed difference and buoy wind speed are shown in Table 3. Both scatterometer-based  
366 observations (ASCAT and QuikSCAT) and WIND Toolkit exhibit smaller slopes among the three  
367 local buoys, indicating less functional dependence of the errors on wind speed relative to other  
368 datasets. We note that at the lowest and highest wind speeds, wind speed differences are less  
369 important for estimating wind power production due to turbine mechanical constraints that require  
370 cut-in and cut-out wind speed restrictions at low and high wind speeds, respectively.

371 **FIGURE 4 LOCATION**

372

### 373 **3.3 Seasonal and diurnal bias**

374 We examined the diurnal and seasonal dependence of bias and error (RMSE) as a function  
375 of both the time of day (i.e., diurnal signal) and month (i.e., seasonal signal) (bias: Figure 5; RMSE:  
376 Figure 6). To ensure that one dataset and reference buoy have the same sample size, we used paired  
377 data for the comparison analysis from Section 3.2. Here, a positive (negative) bias indicates that  
378 the respective dataset overestimates (underestimates) the buoy wind speed.

379 **FIGURE 5 LOCATION**

380 **FIGURE 6 LOCATION**

381 Overall, QuikSCAT and ASCAT show some of the smallest biases among the datasets,  
382 although there are only two hours per day for comparison. Generally, both datasets show different  
383 performance between the early morning and evening. While the bias is consistently low at 46028,  
384 the bias at 46011 is more positive in the mornings, whereas the bias at 46054 is negative in the  
385 evenings. The other satellite-based product, CCMP, is more temporally resolved (6 hour  
386 resolution), but shows much higher bias. Similar to QuikSCAT and ASCAT, CCMP tends to  
387 overestimate buoy-measured wind speeds near 46011. In contrast to 46011, CCMP underestimates  
388 wind speeds near 46028 and 46054.

389 The reanalysis product NARR exhibits consistently low biases at 46028 and 46011, yet  
390 strongly underestimates wind speed (i.e., negative bias) at 46054. Such low and homogeneous  
391 biases at the two northernmost buoy sites (46028 and 46011) are not seen in the other reanalysis  
392 product, MERRA, which displays weaker wind speeds compared to buoy measurements (i.e.,  
393 negative bias) in the morning. The weaker winds in the morning, along with no difference (46028)  
394 or relatively higher wind speeds (46011) in the evening, particularly from May to September, lead  
395 to stronger predicted diurnal cycles than observed at the buoys.

396 Both atmospheric regional model simulations used in this study display lower biases at  
397 46054, compared to other datasets. At 46028 and 46011, WRF overestimates wind speed  
398 throughout the day in summer months. For the WIND Toolkit, wind speed is underestimated (i.e.,  
399 negative bias) close to buoy measurements at 46028. It tends to overestimate wind speed (i.e.,  
400 positive bias) from 00:00 to 12:00 PST at 46011 in contrast to slight underestimates in the evening.

401 Overall, QuikSCAT, ASCAT, and WIND Toolkit are the best performing datasets with the  
402 lowest bias, and hence smallest discrepancies from local buoys. The bias appears to be tied to the  
403 RMSE in which the greater bias corresponds to the greater RMSE. Since the diurnal and seasonal

404 patterns in bias (RMSE) are different across the three buoys, a simple correction of the underlying  
405 dataset is likely to lead to more uncertainties spatially.

406

### 407 **3.4 Tradeoff analysis for seven datasets**

408 Although scatterometer-based observations were the best performing datasets relative to  
409 buoy measurements in this study domain, their temporal resolution is too coarse to fully resolve  
410 the diurnal cycle of near-surface winds. By contrast, the next performing dataset, WIND Toolkit,  
411 provides hourly wind fields with much higher spatial resolution. To evaluate the relative merits of  
412 the datasets and identify the best dataset for offshore wind power applications, we conducted  
413 tradeoff analysis to illustrate important differences in the characteristics of the seven datasets in  
414 relation to five factors: the absolute value of bias, RMSE, data availability aloft, temporal  
415 resolution, and spatial resolution. Here, we considered the overall performance (the absolute value  
416 of bias and RMSE) at the three local buoy sites in this domain (Table 3), but the performance at  
417 individual sites can be obtained in a similar fashion.

418 Figure 7a shows the mean bias and RMSE over the three buoy sites along with error bars  
419 representing one standard deviation from the mean. ASCAT, QuickScat, and WIND Toolkit all  
420 have similarly low levels of bias and RMSE, consistently at the three buoy sites; however, among  
421 these datasets, only WIND Toolkit contains data aloft (Figure 7a). Furthermore, WIND Toolkit  
422 contains a far superior spatial and temporal resolution, compared with ASCAT and QuickScat  
423 (Figure 7b). Only WRF contains spatial and temporal resolution comparable with that by the  
424 WIND Toolkit, but WRF is otherwise inferior because it has a much larger RMSE. Collectively,

425 these tradeoff analysis results indicate that WIND Toolkit is the most appropriate dataset for  
426 supporting offshore wind power applications in this region.

#### 427 **FIGURE 7 LOCATION**

428

### 429 **3.5 Spatial and temporal variations of wind speed over a wide area**

430 Based on the point-to-point comparison and the tradeoff analysis, the WIND Toolkit  
431 appears to be the best dataset for offshore wind power applications and can better estimate wind  
432 speeds daily and seasonally over a wide area. Figure 8 displays the average 10-m wind speeds at  
433 different hours and over four seasons from 2007-2013 using WIND Toolkit. Similar to the three  
434 buoy sites, other areas across the central California region are characterized by strong diurnal  
435 (weaker in the morning and stronger in the evening) and seasonal (stronger in spring and weaker  
436 in fall) variability in the wind speed. The diurnal cycle is enhanced during spring and summer  
437 months, relative to fall and winter months, consistent with data from the three buoy sites shown in  
438 Figure 2. Figure 8 also highlights the local maxima of wind speed near the complex topography of  
439 Point Conception.

#### 440 **FIGURE 8 LOCATION**

441

### 442 **3.5 Characteristics of wind direction**

443 We also assessed the climatology of near-surface wind direction at the local buoys and in  
444 comparison with the other wind datasets. We present wind direction in terms of where the wind is  
445 coming from in degrees clockwise from true north (i.e., 0° wind direction indicates a wind coming

446 from the north and blowing to the south). To account for the direction difference due to the  
447 discontinuity between  $0^\circ$  and  $360^\circ$ , and to quantify the direction difference between  $-180^\circ$  and  $180^\circ$ ,  
448 the wind direction from the respective data set ( $\theta_O$ ) relative to the buoy data ( $\theta_B$ ) was modified  
449 following Pensieri et al. [16]. First, we computed the wind direction difference ( $\theta_O - \theta_B$ ). When  $\theta_O$   
450  $- \theta_B > 180^\circ$ ,  $\theta_O = \theta_O - 360^\circ$  and when  $\theta_O - \theta_B < -180^\circ$ ,  $\theta_O = \theta_O + 360^\circ$ . With the modified  $\theta_O$ , the  
451 wind speed difference was calculated as  $\theta_O - \theta_B$ .

452 Based on the time period of 1998-2016, winds measured at the three buoys are  
453 predominately northwesterly (i.e., along-shore equatorward) (see wind rose histograms in Figure  
454 9). Persistent, but variable in magnitude, northwesterly winds are closely linked to large-scale  
455 pressure systems and the interaction between air flows and topography along the coast (e.g., [13]).  
456 At the site 46054, near-surface winds have more westerly components than the other two sites,  
457 resulting from steering by the adjacent coastline that is oriented in the E-W direction near the Santa  
458 Barbara Channel. Examination of the diurnal cycle shows a more northerly component in the early  
459 morning, followed by a more westerly component in the afternoon (not shown), consistent with  
460 local sea breezes along the Central Coast (e.g., [39]).

#### 461 **FIGURE 9 LOCATION**

462 The error metrics (bias and RMSE) of wind direction from paired data are shown in Table  
463 4. Most of the datasets reveal a positive (i.e., clockwise) bias, with the exception of ASCAT at  
464 46011 and 46054 and NARR at 46054. Similar to the wind speed analysis, the two scatterometers  
465 (QuikSCAT and ASCAT) and WIND Toolkit display the best overall performance in terms of bias  
466 and RMSE for wind direction. QuikSCAT has the lowest bias in wind direction at 46011 and  
467 46054. ASCAT has the lowest bias in wind direction and the second lowest RMSE at 46028.  
468 WIND Toolkit has the second lowest bias and the lowest RMSE at all three sites.

469

470 **4. Discussion and Conclusion**

471 This study provides a comprehensive evaluation of near-surface wind datasets along the central  
472 region of the California coast, ranging from south of Monterey Bay to north of Point Conception.  
473 This particular region has received considerable interest in the development of offshore wind farms  
474 due to its strong, steady winds and existing connections to the state's electrical grid. This study  
475 provides the first known assessment of various wind datasets in this region over both seasonal and  
476 diurnal time scales, both of which are critical for accurate assessment of offshore wind power  
477 production but are seldom considered at the same time by previous studies. In addition, this study  
478 provides a framework by which to assess spatiotemporal variations among various datasets for a  
479 particular region, including comparison of error metrics over both seasonal and diurnal time scales  
480 and tradeoff analysis. This framework can be applied to other regions – using the five factors we  
481 focused on and possibly others of importance – where accurate estimates of wind speed are needed  
482 to evaluate wind energy potential as well as other needs.

483 We examined near-surface wind fields from seven datasets, including satellite  
484 observations, reanalysis products, and regional model output. For each dataset considered, we  
485 found no common pattern of bias and RMSE at all local buoy sites on certain hours of the day or  
486 months of the year. Overall, the two scatterometers, QuikSCAT and ASCAT, showed the best  
487 performance relative to the in-situ buoy measurements. However, the coarse temporal resolution  
488 (i.e., two measurements per day) and spatial resolution (12.5 km) of these datasets limits their  
489 applicability for offshore wind power assessment, particularly since this region experiences strong  
490 diurnal wind forcing and strong spatial gradients in the wind field. On the other hand, WIND  
491 Toolkit was one of the most highly resolved datasets (1 hr temporal and 2 km spatial resolution),

492 and performed nearly as well as the scatterometers in the various error metrics we assessed.  
493 Moreover, the WIND Toolkit has wind data available above the surface and at potential turbine  
494 hub heights, which could obviate interpolation and extrapolation techniques needed with other  
495 data products [3]. Site-specific assessments should consider tradeoffs between spatiotemporal  
496 resolution of the underlying dataset, error metrics relative to local buoy measurements, and the  
497 availability of data at hub height when assessing various data products for offshore wind energy  
498 assessments and power calculations. With consideration of these factors, the WIND Toolkit  
499 appears to be the best dataset for the central California region. Due to the lack of wind observations  
500 at altitudes greater than 5 or 10 m in this region, it is challenging to evaluate offshore wind power  
501 potential at hub height (i.e., heights of at least 100 m above the sea surface based on current  
502 technologies), which is a critical factor considered for future offshore wind siting and  
503 development. Since the surface wind distribution can provide the implications for wind distribution  
504 at hub height, future work will focus on the calculation of wind power generation at hub height  
505 from the WIND Toolkit under different scenarios both spatially and temporally.

506 Finally, tradeoff analysis is a useful graphical tool for comparing the relative performance  
507 of a set of options in relation to multiple objectives. Grounded in Portfolio Theory (maximize  
508 return, minimize risk of financial investments; [40]), we applied tradeoff analysis to factors  
509 important to offshore wind power applications. This analysis revealed the overall superior value  
510 of WIND Toolkit (in relation to the prescribed factors), and more generally demonstrated a  
511 framework that could be used for evaluating wind datasets in other regions. Furthermore, the  
512 tradeoff analysis framework is adaptable, allowing for integration of additional factors important  
513 to offshore wind power applications, including potential impacts of wind energy development on  
514 the marine ecosystem [41, 42]. In such cases the tradeoff analysis axes can be expanded to include

515 these factors, and relative weights can be applied to the factors, in order to help identify  
516 development options that most effectively represent the socio-economic priorities in the system  
517 [43].

518

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524

## 525 **References**

- 526 [1] I. Graabak, M. Korpås, Variability Characteristics of European Wind and Solar Power  
527 Resources—A Review, *Energies*. 9 (2016) 449. doi:10.3390/en9060449.
- 528 [2] X. Sun, D. Huang, G. Wu, The current state of offshore wind energy technology  
529 development, *Energy*. 41 (2012) 298–312. doi:10.1016/j.energy.2012.02.054.
- 530 [3] D. Carvalho, A. Rocha, M. Gómez-Gesteira, C. Silva Santos, Offshore winds and wind  
531 energy production estimates derived from ASCAT, OSCAT, numerical weather prediction  
532 models and buoys – A comparative study for the Iberian Peninsula Atlantic coast, *Renew.*  
533 *Energy*. 102 (2017) 433–444. doi:10.1016/j.renene.2016.10.063.
- 534 [4] W.T. Liu, X. Xie, Measuring ocean surface wind from space, *Remote Sens. Mar. Environ.*  
535 *Man.* 6 (2006) 149–178.

- 536 [5] D. Carvalho, A. Rocha, M. Gomez-Gesteira, C. Santos, A sensitivity study of the WRF  
537 model in wind simulation for an area of high wind energy, *Environ. Model. Softw.* 33  
538 (2012) 23–34. doi:10.1016/j.envsoft.2012.01.019.
- 539 [6] M.H. Pickett, W. Tang, L.K. Rosenfeld, C.H. Wash, QuikSCAT satellite comparisons  
540 with nearshore buoy wind data off the U.S. West Coast, *J. Atmos. Ocean. Technol.* 20  
541 (2003) 1869–1879. doi:10.1175/1520-0426(2003)020<1869:QSCWNB>2.0.CO;2.
- 542 [7] W. Tang, W.T. Liu, B.W. Stiles, Evaluation of high-resolution ocean surface vector winds  
543 measured by QuikSCAT scatterometer in coastal regions, *IEEE Trans. Geosci. Remote*  
544 *Sens.* 42 (2004) 1762–1769. doi:10.1109/TGRS.2004.831685.
- 545 [8] D. Carvalho, A. Rocha, M. Gomez-Gesteira, C. Silva Santos, Comparison of reanalyzed,  
546 analyzed, satellite-retrieved and NWP modelled winds with buoy data along the Iberian  
547 Peninsula coast, *Remote Sens. Environ.* 152 (2014) 480–492.  
548 doi:10.1016/j.rse.2014.07.017.
- 549 [9] I. Alvarez, M. Gomez-Gesteira, M. deCastro, D. Carvalho, Comparison of different wind  
550 products and buoy wind data with seasonality and interannual climate variability in the  
551 southern Bay of Biscay (2000-2009), *Deep. Res. Part II Top. Stud. Oceanogr.* 106 (2014)  
552 38–48. doi:10.1016/j.dsr2.2013.09.028.
- 553 [10] E. Sharp, P. Dodds, M. Barrett, C. Spataru, Evaluating the accuracy of CFSR reanalysis  
554 hourly wind speed forecasts for the UK, using in situ measurements and geographical  
555 information, *Renew. Energy.* 77 (2015) 527–538. doi:10.1016/j.renene.2014.12.025.
- 556 [11] P. Beiter, W. Musial, A. Smith, R. Damiani, M. Maness, T. Stehly, V. Gevorgian, M.  
557 Mooney, G. Scott, P. Beiter, W. Musial, A. Smith, L. Kilcher, R. Damiani, M. Maness, S.

- 558           Sirnivas, T. Stehly, A Spatial-Economic Cost- Reduction Pathway Analysis for U . S .  
559           Offshore Wind Energy Development from 2015 – 2030 A Spatial-Economic Cost-  
560           Reduction Pathway Analysis for U . S . Offshore Wind Energy Development from 2015 –  
561           2030, (2016).
- 562 [12] W. Musial, D. Heimiller, P. Beiter, G. Scott, C. Draxl, 2016 Offshore Wind Energy  
563           Resource Assessment for the United States, (2016). doi:NREL/TP-5000-66599.
- 564 [13] M.R. Fewings, L. Washburn, C.E. Dorman, C. Gotschalk, K. Lombardo, Synoptic forcing  
565           of wind relaxations at Pt. Conception, California, *J. Geophys. Res. Ocean.* 121 (2016)  
566           5711–5730. doi:10.1002/2016JC011699.
- 567 [14] Q. Jiang, J.D. Doyle, T. Haack, M.J. Dvorak, C.L. Archer, M.Z. Jacobson, Exploring wind  
568           energy potential off the California coast, *Geophys. Res. Lett.* 35 (2008) 1–6.  
569           doi:10.1029/2008GL034674.
- 570 [15] R.K. Walter, K.J. Armenta, B. Shearer, I. Robbins, J. Steinbeck, Coastal upwelling  
571           seasonality and variability of temperature and chlorophyll in a small coastal embayment,  
572           *Cont. Shelf Res.* 154 (2018) 9–18. doi:10.1016/j.csr.2018.01.002.
- 573 [16] S. Pensieri, R. Bozzano, M.E. Schiano, Comparison between QuikSCAT and buoy wind  
574           data in the Ligurian Sea, *J. Mar. Syst.* 81 (2010) 286–296.  
575           doi:10.1016/j.jmarsys.2010.01.004.
- 576 [17] P.K. Taylor, E.C. Kent, M.J. Yelland, B.I. Moat, The accuracy of marine surface winds  
577           from ships and buoys, *CLIMAR 99, WMO Work. Adv. Mar. Climatol.* (1999) 59–68.
- 578 [18] W.T. Liu, W. Tang, Equivalent neutral wind, *JPL Publ.* (1996) 17–96.

- 579 [19] S.B. Capps, C.S. Zender, Global ocean wind power sensitivity to surface layer stability,  
580 Geophys. Res. Lett. 36 (2009) 1–5. doi:10.1029/2008GL037063.
- 581 [20] A. Verhoef, A. Stoffelen, Validation of ASCAT 12 . 5-km winds, (2009) 1–11.
- 582 [21] N. Ebuchi, H.C. Graber, M.J. Caruso, Evaluation of wind vectors observed by  
583 QuikSCAT/SeaWinds using ocean buoy data, J. Atmos. Ocean. Technol. 19 (2002) 2049–  
584 2062. doi:10.1109/IGARSS.2001.976753.
- 585 [22]Physical Oceanography Distributed Active Archive Center ( PO . DAAC ): QuikSCAT  
586 Level 2B Version 3 Guide Document. Version 1 7 March (2013) 1–16.
- 587 [23] A. Bentamy, D. Croize-Fillon, C. Perigaud, Characterization of ASCAT measurements  
588 based on buoy and QuikSCAT wind vector observations, Ocean Sci. 4 (2008) 265–274.  
589 doi:10.5194/os-4-265-2008.
- 590 [24] D.E. Weissman, M.A. Bourassa, J. Tongue, Effects of rain rate and wind magnitude on  
591 SeaWinds scatterometer wind speed errors, J. Atmos. Ocean. Technol. 19 (2002) 738–  
592 746. doi:10.1175/1520-0426(2002)019<0738:EORRAW>2.0.CO;2.
- 593 [25] S. Kako, A. Isobe, M. Kubota, High-resolution ASCAT wind vector data set gridded by  
594 applying an optimum interpolation method to the global ocean, J. Geophys. Res. Atmos.  
595 116 (2011) 1–16. doi:10.1029/2010JD015484.
- 596 [26] G. Osi, EUMETSAT Advanced Retransmission Service ASCAT Wind Product User  
597 Manual, (2016) 1–23.
- 598 [27] R. Atlas, R.N. Hoffman, J. Ardizzone, S.M. Leidner, J.C. Jusem, D.K. Smith, D. Gombos,  
599 A Cross-calibrated, Multiplatform Ocean Surface Wind Velocity Product for

- 600 Meteorological and Oceanographic Applications, *Bull. Am. Meteorol. Soc.* 92 (2011)  
601 157–174. doi:10.1175/2010BAMS2946.1.
- 602 [28] M.M. Rienecker, M.J. Suarez, R. Gelaro, R. Todling, J. Bacmeister, E. Liu, M.G.  
603 Bosilovich, S.D. Schubert, L. Takacs, G.K. Kim, S. Bloom, J. Chen, D. Collins, A.  
604 Conaty, A. Da Silva, W. Gu, J. Joiner, R.D. Koster, R. Lucchesi, A. Molod, T. Owens, S.  
605 Pawson, P. Pegion, C.R. Redder, R. Reichle, F.R. Robertson, A.G. Ruddick, M.  
606 Sienkiewicz, J. Woollen, MERRA: NASA's modern-era retrospective analysis for  
607 research and applications, *J. Clim.* 24 (2011) 3624–3648. doi:10.1175/JCLI-D-11-  
608 00015.1.
- 609 [29] F. Mesinger, G. DiMego, E. Kalnay, K. Mitchell, P.C. Shafran, W. Ebisuzaki, D. Jović, J.  
610 Woollen, E. Rogers, E.H. Berbery, M.B. Ek, Y. Fan, R. Grumbine, W. Higgins, H. Li, Y.  
611 Lin, G. Manikin, D. Parrish, W. Shi, North American regional reanalysis, *Bull. Am.*  
612 *Meteorol. Soc.* 87 (2006) 343–360. doi:10.1175/BAMS-87-3-343.
- 613 [30] C. Draxl, A. Clifton, B.M. Hodge, J. McCaa, The Wind Integration National Dataset  
614 (WIND) Toolkit, *Appl. Energy.* 151 (2015) 355–366.  
615 doi:10.1016/j.apenergy.2015.03.121.
- 616 [31] B. Bylhouwer, D. Ianson, K. Kohfeld, Changes in the onset and intensity of wind-driven  
617 upwelling and downwelling along the North American Pacific coast, *J. Geophys. Res.*  
618 *Ocean.* 118 (2013) 2565–2580. doi:10.1002/jgrc.20194.
- 619 [32] X. Li, S. Zhong, X. Bian, W.E. Heilman, Climate and climate variability of the wind  
620 power resources in the Great Lakes region of the United States, *J. Geophys. Res. Atmos.*  
621 115 (2010) 1–15. doi:10.1029/2009JD013415.

- 622 [33] G.W.K. Moore, R.S. Pickart, I.A. Renfrew, Buoy observations from the windiest location  
623 in the world ocean, Cape Farewell, Greenland, *Geophys. Res. Lett.* 35 (2008) 3–7.  
624 doi:10.1029/2008GL034845.
- 625 [34] W.C. Skamarock, J.B. Klemp, J. Dudhi, D.O. Gill, D.M. Barker, M.G. Duda, X.-Y.  
626 Huang, W. Wang, J.G. Powers, A Description of the Advanced Research WRF Version 3,  
627 Tech. Rep. (2008) 113. doi:10.5065/D6DZ069T.
- 628 [35] L. Renault, A. Hall, J.C. McWilliams, Orographic shaping of US West Coast wind  
629 profiles during the upwelling season, *Clim. Dyn.* 46 (2016) 273–289. doi:10.1007/s00382-  
630 015-2583-4.
- 631 [36] M. Nakanishi, H. Niino, An improved Mellor-Yamada Level-3 model: Its numerical  
632 stability and application to a regional prediction of advection fog, *Boundary-Layer  
633 Meteorol.* 119 (2006) 397–407. doi:10.1007/s10546-005-9030-8.
- 634 [37] C. Draxl, B. Hodge, A. Clifton, C. Draxl, B. Hodge, A. Clifton, Overview and  
635 Meteorological Validation of the Wind Integration National Dataset Toolkit, (2015) 1–87.  
636 doi:NREL/TP-5000-61740.
- 637 [38] S.E. Lester, C. Costello, B.S. Halpern, S.D. Gaines, C. White, J.A. Barth, Evaluating  
638 tradeoffs among ecosystem services to inform marine spatial planning, *Mar. Policy.* 38  
639 (2013) 80–89. doi:10.1016/j.marpol.2012.05.022.
- 640 [39] R.K. Walter, E.C. Reid, K.A. Davis, K.J. Armenta, K. Merhoff, N.J. Nidzieko, Local  
641 diurnal wind-driven variability and upwelling in a small coastal embayment, *J. Geophys.  
642 Res. Ocean.* 122 (2017) 955–972. doi:10.1002/2016JC012466.

- 643 [40] H. Markowitz, Portfolio Selection, *J. Finance*. 7 (1952) 77–91. doi:10.1111/j.1540-  
644 6261.1952.tb01525.x.
- 645 [41] C. White, B.S. Halpern, C. V Kappel, Ecosystem service tradeoff analysis reveals the  
646 value of marine spatial planning for multiple ocean uses., *Proc. Natl. Acad. Sci.* 109  
647 (2012) 4696–4701. doi:10.1073/pnas.1114215109/-  
648 /DCSupplemental.www.pnas.org/cgi/doi/10.1073/pnas.1114215109.
- 649 [42] H. Farr, Y. Wang, B. Ruttenberg, R. Walter, C. White, Environmental impacts of  
650 deepwater floating offshore wind and wave energy facilities, *Western Society of*  
651 *Naturalists*, Pasadena, CA, November 2017.
- 652 [43] S.E. Lester, J.M. Stevens, R.R. Gentry, C. V. Kappel, T.W. Bell, C.J. Costello, S.D.  
653 Gaines, D.A. Kiefer, C.C. Maue, J.E. Rensel, R.D. Simons, L. Washburn, C. White,  
654 Marine spatial planning makes room for offshore aquaculture in crowded coastal waters,  
655 *Nat. Commun.* 9 (2018) 945. doi:10.1038/s41467-018-03249-1.
- 656 [44] Dong, C., L. Renault, Y. Zhang, J. Ma, and Y. Cao, 2017: Expansion of West Coast  
657 Oceanographic Modeling Capability. US Department of the Interior, Bureau of Ocean  
658 Energy Management, Pacific. OCS Study BOEM 2017-055. 83 pp  
659 (<https://www.boem.gov/epis/5/5636.pdf>)
- 660 [45] Patel, Prachi (2009, June 22nd). Floating Wind Turbines to Be Tested. *IEEE Spectrum*.  
661 Retrieved from [https://spectrum.ieee.org/green-tech/wind/floating-wind-turbines-to-be-](https://spectrum.ieee.org/green-tech/wind/floating-wind-turbines-to-be-tested)  
662 [tested](https://spectrum.ieee.org/green-tech/wind/floating-wind-turbines-to-be-tested) (accessed 30 July 2018).
- 663 [46] Frith, J. (2017, October 18). The world's first floating wind farm, Statoil's Hywind project,  
664 has been officially opened by Scotland's First Minister Nicola Sturgeon. *Maritime Journal*.

665 Retrieved from <http://www.maritimejournal.com/news101/marine-renewable->  
 666 energy/words-first-floating-offshore-wind-farm-now-open (accessed 2 March 2018).

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Dataset	Type of dataset	Spatial resolution	Temporal resolution	Time used in this study
QuikSCAT	Satellite (Swath)	12.5km	2 times per day	2000-2008
ASCAT	Satellite (Swath)	12.5km	2 times per day	2007-2013
CCMP V2	Satellites and analyses	0.25°lat/lon	4 times per day	2004-2013
NARR	Regional reanalysis	32 km	8 times per day	2004-2013
MERRA	Global reanalysis	1/2°lat-2/3°lon	Hourly	2004-2013

WRF	Regional model	6 km	Hourly	2004-2013
WIND Toolkit	Regional model	2 km	Hourly	2007-2013

672

673 Table 1: Characteristics of wind datasets considered for comparison with buoy observations.

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Buoy	Dataset	Slope	Intercept	R <sup>2</sup>	Distance from buoy (km)	Number of valid pairs
46028	QuikSCAT	0.90	1.01	0.92	5.59	5654
	ASCAT	0.94	0.25	0.94	4.91	2153
	CCMP V2	0.68	1.37	0.77	9.80	13013
	NARR	0.76	1.59	0.75	9.07	26018
	MERRA	0.59	1.69	0.74	26.85	78014
	WRF	0.38	4.88	0.15	3.64	78012
	WIND	0.79	1.13	0.83	0.62	55957

46011	QuikSCAT	0.85	1.56	0.84	4.45	5449
	ASCAT	0.89	1.02	0.84	3.55	2241
	CCMP V2	0.75	2.00	0.67	13.22	12368
	NARR	0.72	1.40	0.69	7.62	24695
	MERRA	0.62	2.04	0.64	29.07	74049
	WRF	0.39	4.77	0.14	3.24	74046
	WIND	0.77	1.65	0.73	0.89	51953
46054	QuikSCAT	0.68	2.23	0.82	4.93	3875
	ASCAT	0.81	1.16	0.88	3.77	1670
	CCMP V2	0.53	1.66	0.62	15.41	7928
	NARR	0.49	1.13	0.53	7.62	15754
	MERRA	0.40	1.76	0.58	31.41	47392
	WRF	0.37	4.91	0.16	2.10	47385
	WIND	0.80	1.45	0.79	0.99	34001

676

677 Table 2: Statistical metrics from the linear regression between buoy data and each of the  
 678 comparison datasets.

679

680

Buoy	Dataset	Bias	RMSE	Slope	Intercept	R <sup>2</sup>
46028	QuikSCAT	0.26	1.22	-0.10	1.01	0.12
	ASCAT	-0.18	1.03	-0.06	0.25	0.06
	CCMP V2	-0.94	2.20	-0.32	1.37	0.42
	NARR	-0.15	2.03	-0.24	1.59	0.23
	MERRA	-1.28	2.54	-0.41	1.69	0.58
	WRF	0.41	4.47	-0.62	4.88	0.32
	WIND	-0.43	1.76	-0.21	1.13	0.26
46011	QuikSCAT	0.67	1.50	-0.15	1.56	0.14
	ASCAT	0.38	1.40	-0.11	1.02	0.08
	CCMP V2	0.52	2.02	-0.25	2.00	0.19
	NARR	-0.29	1.90	-0.28	1.40	0.26

	MERRA	-0.21	2.04	-0.38	2.04	0.40
	WRF	1.17	4.07	-0.61	4.77	0.28
	WIND	0.28	1.82	-0.23	1.65	0.18
46054	QuikSCAT	-0.49	2.03	-0.32	2.23	0.50
	ASCAT	-0.33	1.47	-0.19	1.16	0.31
	CCMP V2	-1.99	3.29	-0.47	1.66	0.57
	NARR	-2.78	4.00	-0.51	1.13	0.55
	MERRA	-2.87	4.08	-0.60	1.76	0.75
	WRF	0.04	4.43	-0.63	4.91	0.36
	WIND	-0.07	1.90	-0.20	1.45	0.19

681

682 Table 3: Statistics from the comparison between the buoy data and comparison datasets, including  
683 error metrics (bias and RMSE), as well as outputs (slope, intercept, coefficient of determination)  
684 from the linear regression between the wind speed difference and the buoy wind speed.

685

Buoy	Dataset	Bias	RMSE
46028	QuikSCAT	6.74	38.74

	ASCAT	1.49	37.36
	CCMP V2	6.51	41.86
	NARR	3.77	40.85
	MERRA	5.95	41.98
	WRF	7.10	68.88
	WIND	3.84	36.76
46011	QuikSCAT	0.44	45.83
	ASCAT	-3.09	47.30
	CCMP V2	2.02	44.47
	NARR	0.26	41.96
	MERRA	0.28	44.49
	WRF	14.50	68.05
	WIND	0.99	39.97
46054	QuikSCAT	1.39	46.85
	ASCAT	-3.86	46.73
	CCMP V2	7.78	42.85
	NARR	-8.32	45.03

MERRA	11.12	45.30
WRF	14.18	69.37
WIND	2.70	38.65

686

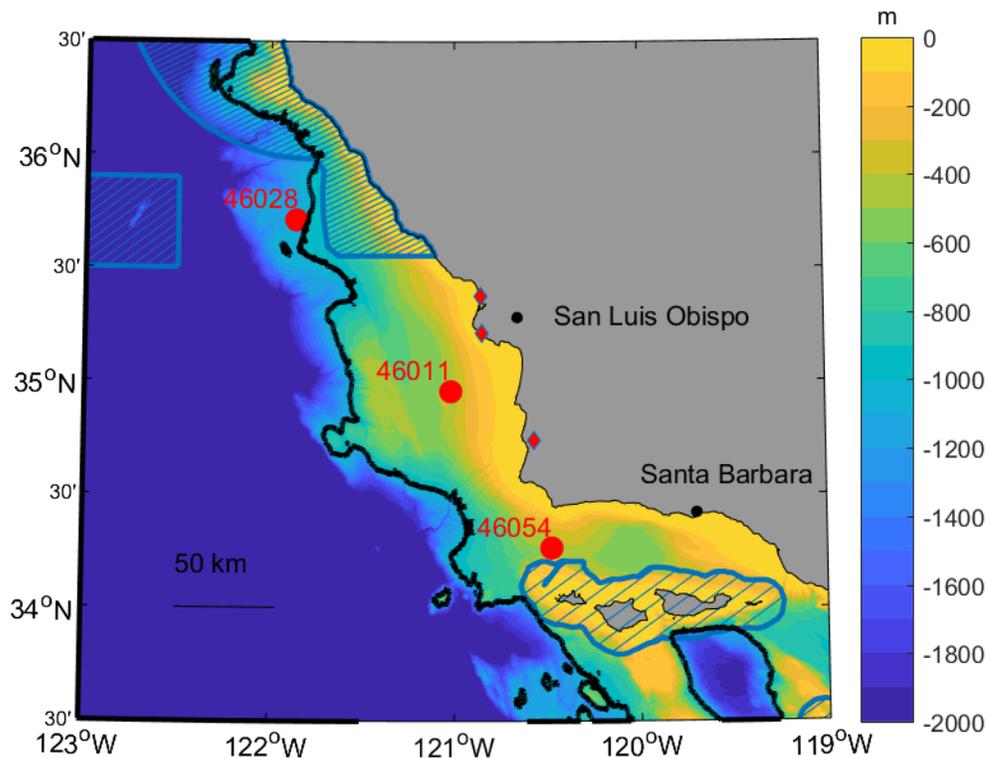
687 Table 4: Statistical metrics of wind direction ( $^{\circ}$ ) from paired data. A positive bias indicates a  
688 clockwise bias.

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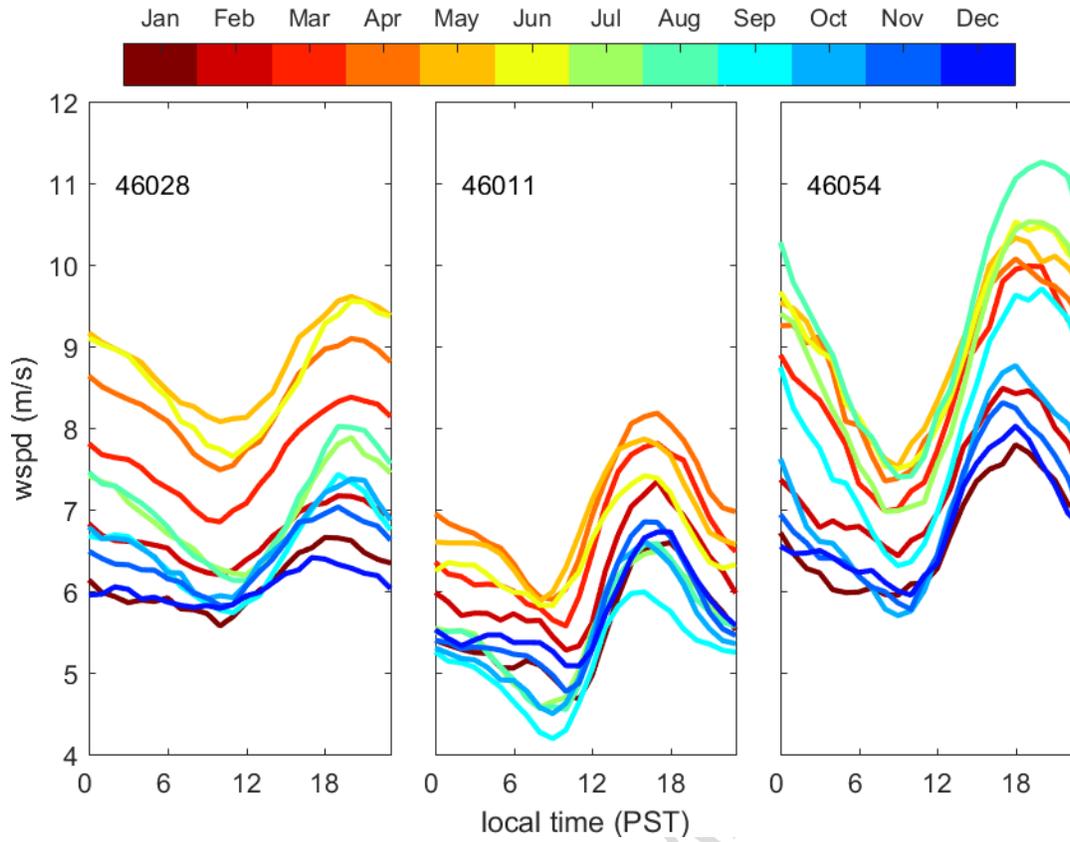
694 Figure 1: Bathymetry of the Central California Coast highlighting the locations of buoy platforms  
 695 (red circles, representing buoys 46028, 46011, 46054 from north to south), existing state electrical  
 696 grid connections (red diamonds), National Marine Sanctuaries (dashed blue lines; Monterey Bay  
 697 Sanctuary to the north and Channel Islands Sanctuary to the south), and the 1000 m isobath (solid  
 698 black line). The state electrical grid connections from north to south are the Morro Bay power  
 699 plant, Diablo Canyon nuclear power plant, and Vandenberg Air Force Base.

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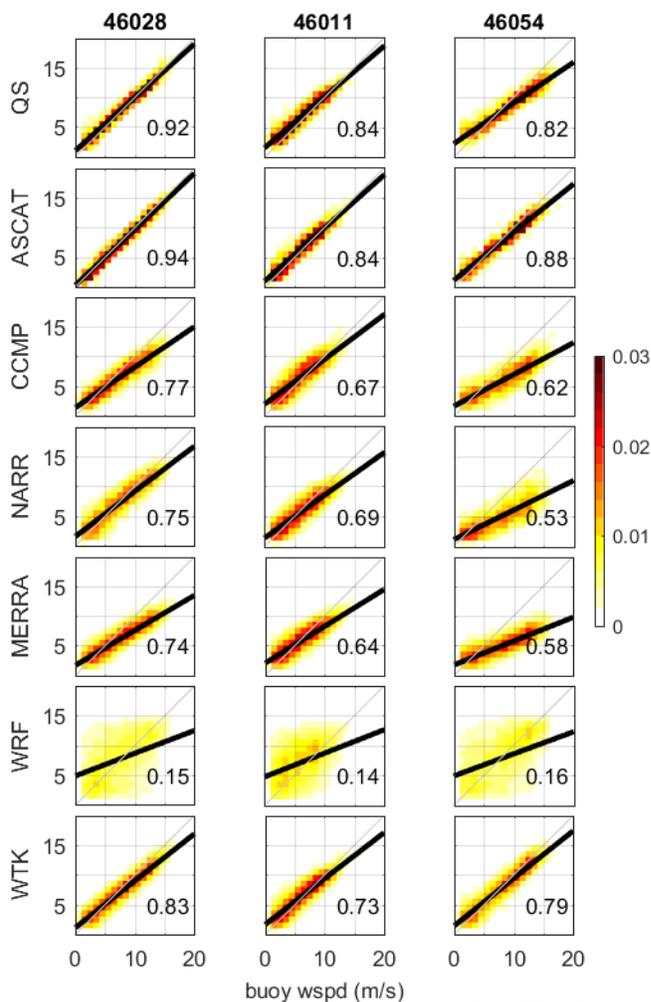


704

705 Figure 2: Composite day average buoy wind speed for a particular month (colors) using data

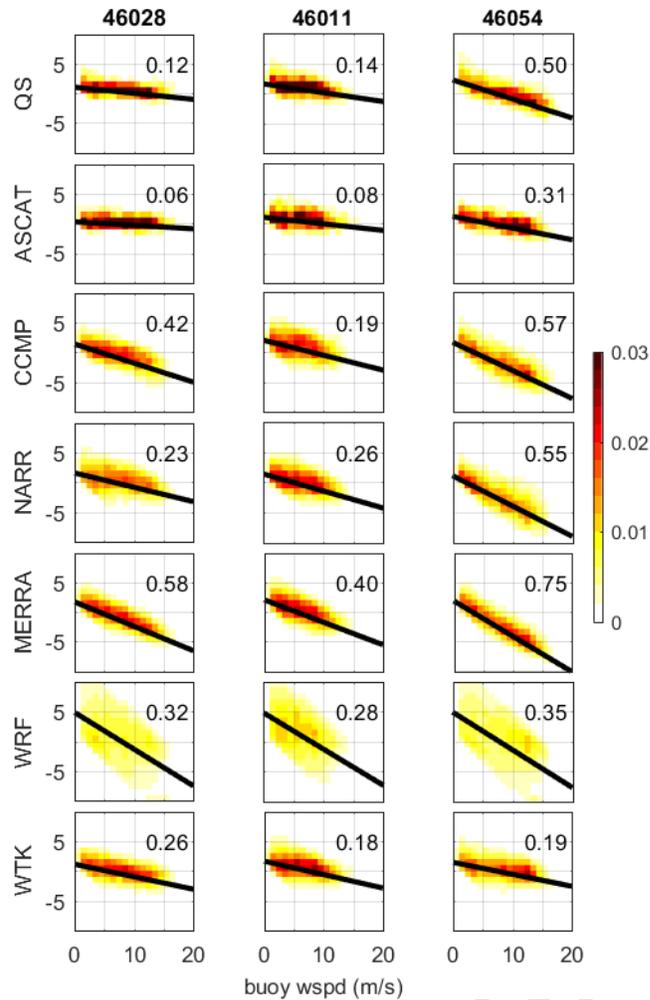
706 calculated over all years (1998-2016) for each buoy (46028, 46011, and 46054 from left to right).

707



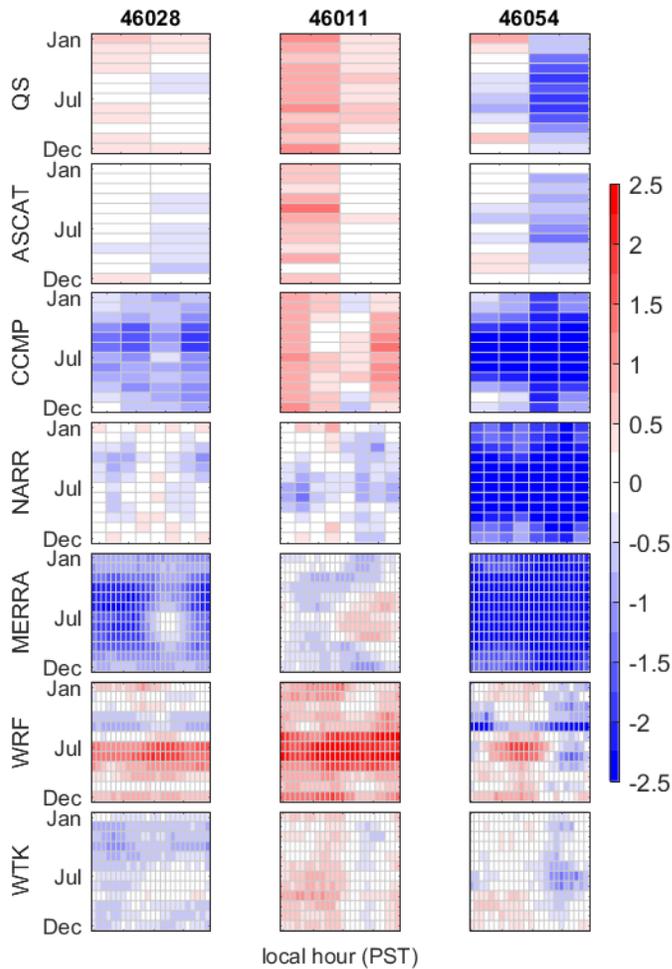
708

709 Figure 3: Comparisons of wind speed between the buoy measurements and each respective dataset  
 710 (all in  $\text{m s}^{-1}$ ). The value on each subplot shows the coefficient of determination ( $R^2$ ) from a linear  
 711 regression model (fit shown as bold black line). The one-to-one line is also shown for reference  
 712 (thin gray line). Wind speed is binned by  $1 \text{ m s}^{-1}$  along both of the x-axis and y-axis, and then  
 713 divided by the total number of data pairs to yield the frequency of data points in a particular bin  
 714 (colors). The rows from the top to the bottom are QuikSCAT (QS), ASCAT, CCMP V2.0 (CCMP),  
 715 NARR, MERRA, WRF, and WIND Toolkit (WTK). The columns from the left to the right  
 716 represent the local buoy 46028, 46011, and 46054, respectively. Note that the time period used for  
 717 analysis depends on data availability.



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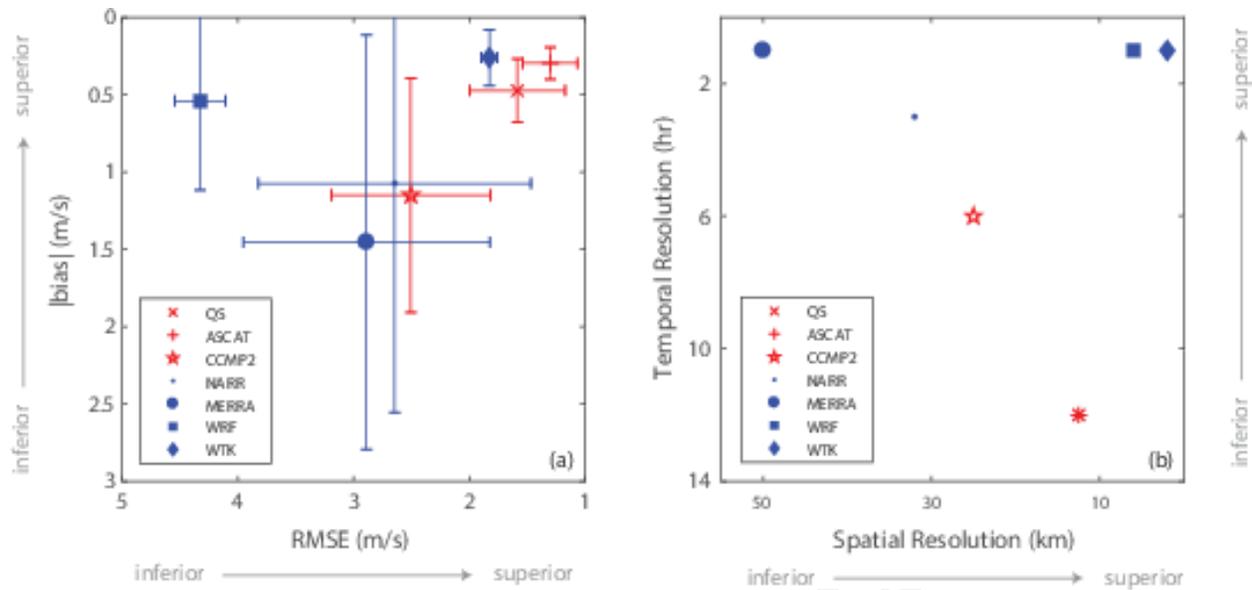
719 Figure 4: Differences in wind speed between the buoy and the other respective datasets as a  
 720 function of the buoy wind speed (all in  $\text{m s}^{-1}$ ). The value on each subplot shows the coefficient of  
 721 determination ( $R^2$ ) from a linear regression model (black line). Wind speed is binned by  $1 \text{ m s}^{-1}$   
 722 along both of the x-axis and the y-axis, and then divided by the total number of data pairs to yield  
 723 the frequency of data points in a particular bin (colors). The rows from the top to the bottom are  
 724 QuikSCAT (QS), ASCAT, CCMP V2.0 (CCMP), NARR, MERRA, WRF, and WIND Toolkit  
 725 (WTK). The columns from the left to the right represent the local buoy 46028, 46011, and 46054,  
 726 respectively. Note that the time period used for analysis depends on data availability.



727

728 Figure 5: Bias ( $\text{m s}^{-1}$ ) in the hourly near-surface wind speed in each month for all available paired  
 729 data in relation to the buoy measurements at 46028 (left), 46011 (middle), and 46054 (right). A  
 730 positive (negative) bias indicates that the respective dataset overestimates (underestimates) the  
 731 buoy wind speed. The white color indicates zero bias. The rows from the top to the bottom are  
 732 QuikSCAT (QS), ASCAT, CCMP V2.0 (CCMP), NARR, MERRA, WRF, and WIND Toolkit  
 733 (WTK). The following hours (in PST) are shown for the respective dataset: QS (05 and 18);  
 734 ASCAT (9 and 20); CCMP (04, 10, 16, and 22); NARR (01, 04, 07, 10, 13, 16, 19, and 22); and  
 735 for MERRA, WRF, and WTF (hourly from 00 to 23). See Table 1 for the time period used for  
 736 analysis of individual datasets.



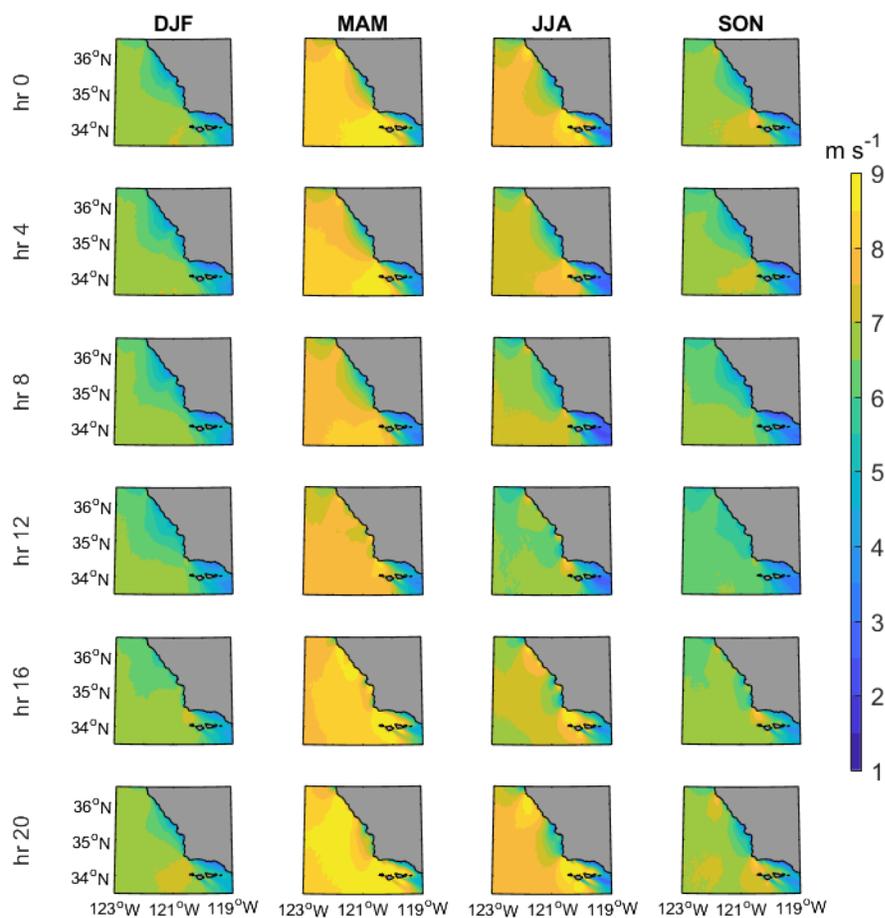


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746 Figure 7: Pairwise tradeoffs in relation to different factors for seven datasets. Blue color represents  
 747 the data availability aloft, while red color represents no data available aloft. (a) The absolute value  
 748 of the bias and RMSE. The markers represent the mean and the error bars represent one standard  
 749 deviation from the mean. (b) The temporal and spatial resolution. QS and ASCAT have the same  
 750 temporal and spatial resolution so they are overlapping in the panel (b). For the MERRA data, we  
 751 show the spatial resolution in the latitudinal direction (see Table 1).

752



753

754 Figure 8: Averages of the hourly 10-m wind speed from WIND Toolkit over 2007-2013 at different  
 755 hours and four seasons. Each column from the left to the right represents winter (December-  
 756 January-February, DJF), spring (March-April-May, MAM), summer (June-July-August, JJA), and  
 757 fall (September-October-November, SON). Each row from the top to the bottom represents 00  
 758 PST, 04 PST, 08 PST, 12 PST, 16 PST, and 20 PST.

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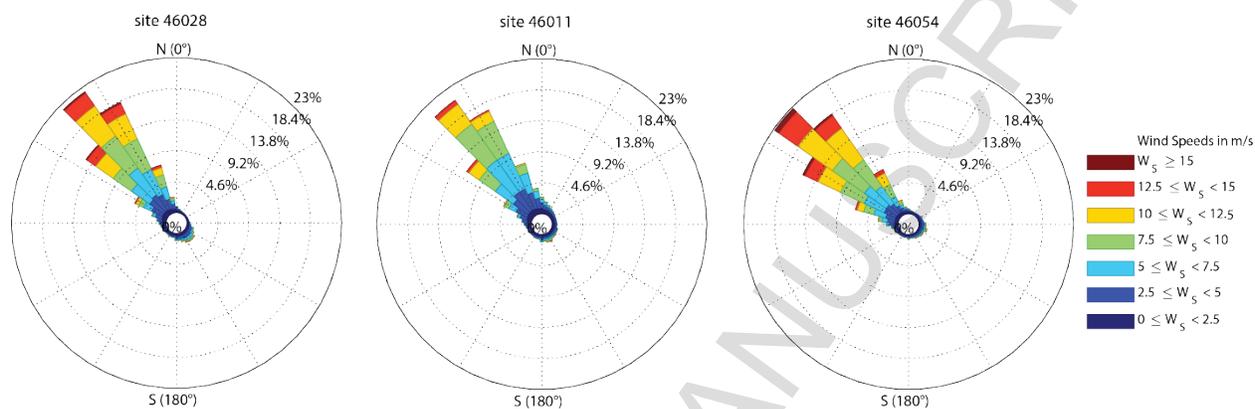
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767 Figure 9: Wind rose histograms using data from 1998-2016 for the three buoys considered in this  
 768 study (46028, 46011, and 46054 from left to right, respectively). The direction shown is the  
 769 direction from which the wind is coming from in degrees clockwise from true north (i.e., 0° wind  
 770 direction indicates a wind coming from the north and blowing to the south).