

A predictive model for Lake Chad total surface water area using remotely sensed and modeled hydrological and meteorological parameters and multivariate regression analysis

Frederick Policelli ^{a,d}, Alfred Hubbard ^{b,c}, Hahn Chul Jung ^{a,c}, Ben Zaitchik ^d, Charles Ichoku ^e

^a NASA Goddard Space Flight Center
8800 Greenbelt Rd
Code 617, Hydrological Sciences Lab
Greenbelt, MD, 20771, USA

^b NASA Goddard Space Flight Center
8800 Greenbelt Rd
Code 618, Biospheric Sciences Lab
Greenbelt, MD, 20771, USA

^c Science Systems and Applications, Inc.
10210 Greenbelt Rd
Lanham, MD, 20706, USA

^d Johns Hopkins University
Dept. of Earth and Planetary Sciences

22 3400 N. Charles St.

23 Baltimore, MD 21218, USA

24

25 ^e Howard University

26 Dept. of Interdisciplinary Studies

27 1840 7th Street, NW

28 Washington, DC 20001, USA

29

30 * Correspondence: frederick.s.policelli@nasa.gov; Tel.: +01-301-614-6573

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Abstract

Lake Chad is an endorheic lake in west-central Africa at the southern edge of the Sahara Desert. The lake, which is well known for its dramatic decrease in surface area during the 1970s and 1980s, experiences an annual flood resulting in a maximum total surface water area generally during February or March, though sometimes earlier or later. People along the shores of Lake Chad make their living fishing, farming, and raising livestock and have a vested interest in knowing when and how extensive the annual flooding will be, particularly those practicing recession farming in which the fertile ground of previously flooded area is used for planting crops. In this study, the authors investigate the relationship between lake and basin parameters, including rainfall, basin evapotranspiration, lake evapotranspiration, lake elevation, total surface k U h Y f ' U f Y U ž ' U b X ' h \ Y ' d f Y j] c i g ' mY U f Đ g ' h c h U ` ' g i f season month (except November) linking total surface water area to the other parameters. The resulting equations allow the user to estimate the December average monthly total surface water area of the lake in late November, and to make the estimates for January to May in early December. Based on the results of a Leave One Out Cross Validation analysis, the equations for lake area are estimated to have an average absolute error ranging from 5.3 percent (for February estimates) to 7.6 percent (for May estimates).

¹Keywords: surface water area, Lake Chad, precipitation, evapotranspiration, lake height

1. Introduction

Lake Chad is a shallow, endorheic lake in the Sahel region of west-central Africa shared by Chad, Nigeria, Niger, and Cameroon. There have been numerous hydrological models developed for Lake Chad. Each of these has limitations relative to the statistical model for lake area we present here.

Bader et al. (2011) developed a hydrological model that simulates the water level in the northern pool, southern pool, and archipelago using riverine and direct rainfall inputs to the lake. It also estimates total water area for each of the pools. According to (Lemoalle et al., 2012), the model results correspond well with satellite measurements of northern pool surface water area and with satellite measurements of the total surface water area from 1980 onward, though these

Abbreviations: Inter-Tropical Convergence Zone (ITCZ), above sea level (ASL), Climate Hazards Group InfraRed Precipitation with Station data (CHIRPS), Famine Early Warning Systems Network (FEWS NET) Land Data Assimilation System (FLDAS), Land Surface Temperature (LST), evapotranspiration (ET), Leave One Out Cross Validation (LOOCV), precipitation (P). For reading regression results: Pws is total south basin wet season precipitation, ETws is total south basin wet season evapotranspiration, LakeETws is wet season lake ET percent variation from the 1988 to 2016 wet season mean, H is November lake height variation from the 1993-2002 mean, A is the average surface water area for the given month, A- is h \ Y ' d f Y j average surface water area for the given month.

results are not quantified. One disadvantage of this model is that it requires input of the Chari River and Komodougu-Yobe River discharges, data that is not publicly available.

Coe and Foley (2001) report the results of a hydrological model of Lake Chad. They describe

U' a c X Y` ' k] h \ ' í [c c X ' U [f Y Y a Y b h ' k] h \ ' h \ Y 1954 b Z Y f f Y h c ' %- * + " ' ' H \ Y ' í] b Z Y f f Y X ' ` U _ Y ' a, but rather is derived from a relationship between lake area and lake level. It is important to note that the hydrology of @U _ Y ' 7 \ U X ' W \ U b [Y X ' g] [b] Z] W U b h ` m '] b ' 7 \ U X %- h \$ Ð g í G a U ` ` ' @U _ Y ' 7 \ U b X í based on a coarse (10 km resolution) digital elevation model, may not be adequate to define the lake after this transition.

Gao et al. (2011) developed a hydrological model of Lake Chad. They compared images of the lake extent from the model with images derived from remote sensing. Three image pairs were shown, for October 31, 1963; December 25, 1972; and January 31, 1987. Each pair of images Z f c a ' h \ Y ' h k c ' Y U f `] Y f ' X U h Y g ' fl V Y Z é G a U h \ Y ' h f U b g] @U _ Y ' 7 \ U X í is similar, though no numerical value was provided. The model-observation pair for the post transition period did not look very similar. This raises the question of the utility of the model for producing lake area in the current, post transition period.

Lemoalle (2004) developed a crude expression for Lake Chad surface area based on a g] a d `] Z] Y X ' k U h Y f ' V U ` U b W Y ' a c X Y ` ' U b X ' X a s s u m e s no seepage from the lake and requires knowledge of the streamflow to the lake, which as previously noted is not publicly available. Delclaux et al. (2008) developed a hydrological

model of the Lake Chad Basin, however the results they presented included streamflow and elevation, but not surface area.

Coe and Birkett (2004) used upstream measurement of river height along with in-situ stream flow and gauge height to estimate river discharge 500 km downstream and wet season height of Lake Chad, greater than 600 km downstream. Their method, though, clearly relies on hard-to-obtain in-situ measurements and does not include lake area.

The first objective of this paper is to assemble and examine a set of satellite- and model-based data for the southern Lake Chad Basin relevant to developing a statistical model for the area of the lake during its flooding season. This data set includes time series of satellite- and gauge-based precipitation and modeled ET for the southern part of the basin, modeled lake ET data, satellite-based lake elevation data, and satellite-based estimated lake total surface water area. Given the limitations of the existing models described above, the second objective is to develop a predictive statistical model for total lake surface water area using regression methods on the data set. The regression method includes backward elimination variable selection and a Leave One Out Cross Validation (LOOCV) analysis to optimize the resulting statistical model.

2 Study Area

The Lake Chad basin is approximately 2.5 million square kilometers, about eight percent of the African continent, and the largest endorheic basin in the world (Gao et al., 2011). Lake Chad is the terminal lake of this basin. The northern part of the basin lies within the Sahara and does not generate runoff that reaches Lake Chad (Delclaux et al., 2008).

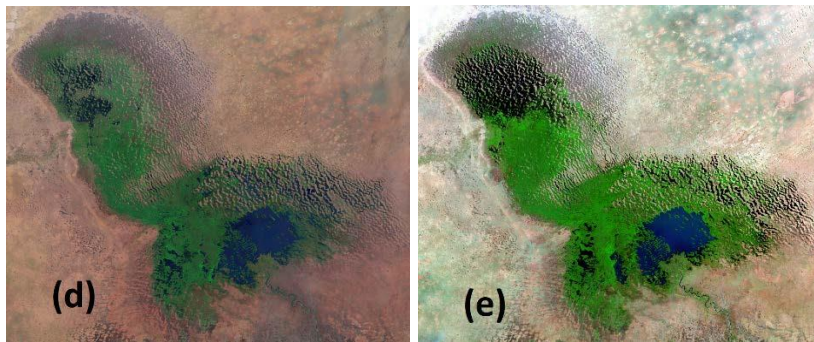


Fig. 2 Evolution of Lake Chad. Optical imagery from (a) 1963, Argon satellite (b) 1973, Landsat 1 (c) 1987, Landsat 5 (d) 2003, Landsat 7 (e) 2013, Landsat 8. Images provided by U.S. Geological Survey, Department of the Interior/USGS

Below about 280 m ASL, the lake separates into a southern pool and a northern pool divided by islands (Lemoalle, 2004).

The population of the lake shore is around 2 million (Magrin, 2016) and the people make a living through a combination of fishing, farming, and raising livestock (Sarch and Birkett, 2000).

hydrology, it is difficult to provide farmers with information on the timing of the floods and a sense of how large the flood is going to be in any given year. This can be a serious problem for farmers who grow crops near the lake shore and periodically lose crops to flooding (Okpara et

al., 2016).

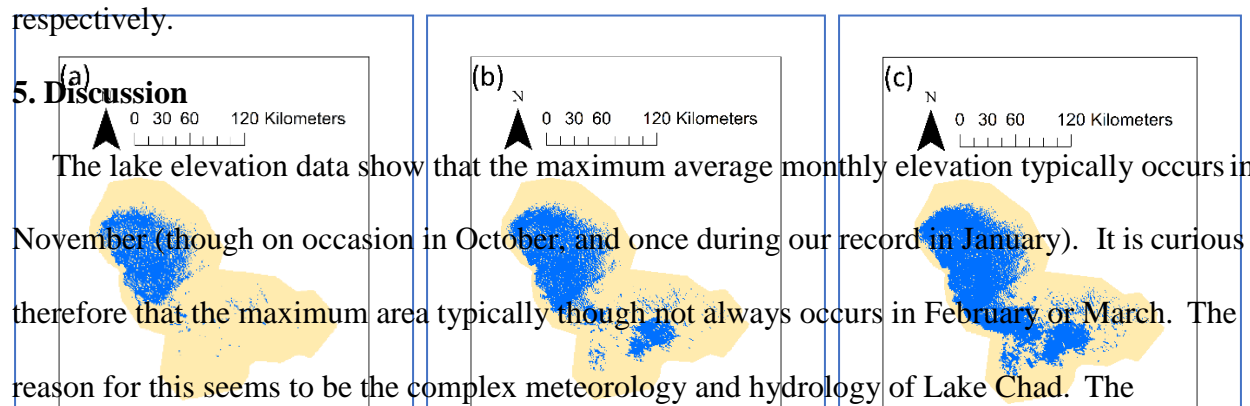
The southwest shore of Lake Chad begins in mid-January to late February. It seems reasonable to conclude that predictions of lake surface area made in late November or early December could be used for agricultural decision making for locations with similar start dates.

There are a number of dams and irrigation schemes in the river basins that drain into Lake Chad, such as those of the Chari-Logone and Komadougou-Yobe river systems. However, at the large scale, they do not amount to a substantial revision of the natural seasonal hydrological patterns, which are largely determined by the West African Monsoon and the position of the ITCZ (Birkett, 2000).

The surface of Lake Chad is not flat and level; the barriers to flow from the southern pool to the northern pool and archipelago result in the water in these areas frequently being at different elevations. Additionally, the local winds and flow of water from the Chari River into the lake (addressed further in the Discussion section) combined with the complex shape of the lake lead to an evolving lake surface (Carmouze et al., 1983). Lake surface elevation time series data used in this study refer to satellite-based measurements made in the relatively small open water portion of the lake in the southern pool.

Figure 3 presents areas of the lake at or below selected elevations. The topographic data used to create figure 3 is from a 1 arc-second (~30m) resolution Digital Elevation Model (DEM) from the NASA Shuttle Radar Topography Mission (SRTM). The blue areas would correspond to lake levels (limited by the accuracy of the DEM) if the lake surface were flat and the lake filled

7.6%, while however, using this average May area, the area should only be 4.1% and water providing
 13,363 sq. km, these percentages are equivalent to 1017 sq. km, 1751 sq. km, and 2085 sq. km



movement of the water in Lake Chad is influenced by both the winds and the Chari-Logone Fig. 3 Blue areas have elevations at or below (a) 279m ASL, (b) 280m ASL, (c) 281m ASL water supply. Monsoon winds drive the displacement of the southern waters toward the north, and movement begins around the northeastern end of the Great Barrier in June, at the end of the According to (Leblanc et al., 2011) most of the flooded area of Lake Chad is covered by low water. The Chari-Logone flood waters begin in August and provide half of their water in aquatic vegetation including rooted and floating plants. This area is not readily measured with optical remote sensing, but must be accounted for to get an accurate estimate of total surface back toward the southern pool. This is also when the satellite radar altimeters (which collect data over the southern pool) typically record the highest levels. During the peak of the riverine flooding, water again reaches the northern pool and also spreads into the archipelago from the south basin. Following the end of the movement of water to the north pool in January, there is a water arrives from such a distance that the peak lake level and lake area occur months after the general spreading of water in the southern pool to the periphery until April. (Carmouze et al., 1983). These movements result in a complex and changing lake surface topography, and are Lake Chad from the southern portion of the basin where precipitation rates are highest (Leblanc et al., 2011). It is the reason that the Lake Chad peak level and area take place in the dry season.

The main limitation of the model we present here is that it is not a physically-based model and may not perform well for conditions outside of those in the database we have developed. For example, we are not able to do regressions and provide predictions for wet season months because we do not have surface water area data for those conditions (Policelli et al., 2018). Also, the performance may degrade if it is used outside of the range of areas for which the regression models were established (8,700 sq. km – 16,800 sq. km) or when the hydrology of the lake changes substantially. There is no fixed date at which the model becomes unusable. However, if it is used for operational forecasting, it would be wise to regularly update the model with new data as it becomes available.

6. Conclusions

We have built a record of remote sensing data and model products (precipitation, evapotranspiration, lake height, and lake area) for the Lake Chad Basin and used this record to run a correlation analysis and a regression analysis. From the correlation analysis (see Appendix), we have found (1) the highest correlation between basin evaporation and total lake surface water area is 0.43 and occurs with a seven month lag, (2) the highest correlation between lake height anomaly and total lake surface water area is 0.57 and occurs with a four month lag, (3) the highest correlation between precipitation and total lake surface water area is 0.39 and occurs with a seven month lag, (4) the highest correlation between percent of 1988-2016 average lake ET and the lake total surface water area is 0.65 at zero lag time, and (5) there is a correlation

The predictions using all of the equations derived from regression in this study can be made with only remotely sensed data and model outputs; no in-situ data is required. Any improvements in the measurement of the parameters we use in this analysis would likely improve the desired end result – the prediction of the Lake Chad surface area in time to be used for agricultural decisions.

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Appendix

We examined the correlation between P \pm ET for the southern Lake Chad Basin and the lake elevation variation relative to 1993-2002 for 1992 to 2017, and found the maximum correlation of 0.69 at four months lag time between these variables. Next, we examined the correlation between the P \pm ET for the southern part of the basin and the lake surface area and found the maximum correlation of 0.37 at eight months lag time. We also examined the correlation

between the lake elevation variation and the lake surface water area and found the maximum correlation of 0.57 at four months lag time. We found a maximum correlation of 0.43 for ET vs

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To determine the value added by the ET data to our analysis, we examined the correlation between precipitation (without subtracting ET) for the southern Lake Chad Basin and the lake elevation and found an increase in the maximum correlation to 0.80 at four months lag time. We found the correlation between the percentage of the 1988-2016 average lake ET and the lake total surface water area to be 0.65 at zero lag time. We also examined the correlation between precipitation and total lake surface area and found an increase in the maximum correlation to 0.39 at seven months lag time. This is the same lag time as for the maximum correlation between ET and total lake surface area and represents the time it takes for much of the net precipitation to make its way from run off in the southernmost part of the Lake Chad basin, to flowing through the Chari-Logone River system, to reaching the lake and causing an increase in the lake area. A surprising result of the correlation analysis is that the use of FLDAS ET data in the analysis to produce (P-ET) causes a small decrease in the correlation numbers relative to what is achieved with precipitation alone. Note however that ET is somewhat more closely correlated with total lake surface water area than is precipitation (.43 vs. .39, both at 7 months lag time). Precipitation and ET have a maximum correlation coefficient of 0.93 with a one month delay between the two.

