

# Analysis of relationship between meteorological and agricultural drought using standardized precipitation index and vegetation health index

U Ma'rufah<sup>1</sup>, R Hidayat<sup>1\*</sup>, I Prasasti<sup>2</sup>

<sup>1</sup>Department of Geophysics and Meteorology, Faculty of Mathematics and Natural Science, Bogor Agricultural University, Bogor, Indonesia

<sup>2</sup>Remote Sensing Application Center, Indonesian National Institute of Aeronautics and Space (LAPAN), Pasar Rebo, Jakarta, Indonesia

E-mail: rahmath@apps.ipb.ac.id

**Abstract.** Agricultural drought is closely related to meteorological drought in which the agricultural drought is an impact of meteorological drought. This study aim to understand the duration, spatial extent, severity and lag time of meteorological and agricultural drought during El Niño years. The data used in this study are monthly data of CHIPRS and MODIS. Meteorological drought and agricultural drought are intensified in the El Niño years. The duration of meteorological drought is different in each region but generally occurs during June to November. Agricultural drought mostly occurs from August to November. Spatially, meteorological drought and agricultural drought in 2015 has wider extent and higher severity (SPI <-2 and VHI <10) than in 2002. Agricultural drought generally intensified in areas that have monsoonal rainfall type such as Java, Bali, Nusa Tenggara, Lampung, southern part of Kalimantan, and southern part of Sulawesi. We found that VHI is significantly correlated with SPI-3 reach 58% of the total area of Indonesia. It means rainfall deficit during three months has a significant impact on agricultural drought in Indonesia. In general, SPI-3 and VHI clearly explain the relationship between meteorological drought and agricultural drought in Indonesia.

## 1. Introduction

El Niño is widely known as global-scale climate anomalies which clearly modulates rainfall anomalies in many parts of the world including Indonesia [1]. Decreasing rainfall during El Niño frequently causes dryness over major part of Indonesia. In Indonesia some extreme drought occurred simultaneously with El Niño event such as in 1997/1998, 2002/2006, 2006/2007, and 2009/2010 [2]. Indonesian National Board for Disaster Management (BNPB) also mentioned that El Niño in 2015 resulted in Puso in 111,000 ha farming area [3].

Drought has a significant impact on the economic, agriculture, environment, and social. Drought event during El Niño 2002 caused a Puso in 42,000 ha of paddy fields in Java [4]. Monitoring drought needs to be done to minimize such drought impact. Various drought indices have been developed by a number of previous studies to determine the duration and severity of drought [5]. These indices have been developed based on different parameters so that each drought indices reflects different drought conditions.



Drought indices is widely applied for monitoring drought are Standardized Precipitation Index (SPI) and the Vegetation Health Index (VHI) [6-8]. SPI is a meteorological drought index that shows level of drought due to rainfall deficit while the VHI is an agriculture drought index based on remote sensing that shows stress level of crops due to drought. VHI is constructed by combining the Vegetation Condition Index (VCI) and the Temperature Condition Index (TCI) [6-10].

Agricultural drought is closely related to meteorological drought in which the agricultural drought is basically the impact of meteorological drought. In general, plants do not respond to the meteorological drought directly. It seems there is a lag time between meteorological drought and agricultural drought. Therefore, the objectives of this study are (1) to identify the duration, spatial extent, and severity of drought and (2) to determine lag time of meteorological drought and agricultural drought in Indonesia during El Niño events.

## 2. Methodology

### 2.1. Standardized Precipitation Index

SPI is calculated in different time scale 1,3,6,9, and 12 months (SPI-, SPI-3, SPI-6, SPI-9 and SPI-12) based on CHIRPS data. CHIRPS is rainfall data set incorporating infrared Cold Cloud Duration (CCD) observation and in-situ station data for trend analysis and seasonal drought monitoring [11]. SPI calculation requires historical precipitation data at least 30 years. So that, CHIRPS data from 1981 to 2015 are used to do the calculation. The SPI calculation is conducted by following McKee *et.al* (1993) [12]. Meteorological drought category based on SPI shown in Table 1.

**Table 1.** Classification of drought severity by WMO [13]

SPI	Category
$-1,0 < \text{SPI} \leq 1,0$	Normal
$-1,5 < \text{SPI} \leq -1,0$	Moderate
$-2,0 < \text{SPI} \leq -1,5$	Severe
$\text{SPI} < -2,0$	Extreme

### 2.2. Vegetation Health Index

MODIS data that consist of Enhanced Vegetation Index (EVI) and Land Surface Temperature (LST) are used to calculate VCI and TCI, respectively [8]. We use time series data from 2001 to 2015 to calculate their minimum and maximum values. The calculation of VCI and TCI is shown by following equations:

$$\text{VCI} = 100 \times \frac{E - E_{\min}}{E_{\max} - E_{\min}} \quad (1)$$

where E is EVI value of a given month.  $E_{\min}$  and  $E_{\max}$  denote the maximum and minimum EVI values, respectively, for the month from multiyear time series.

$$\text{TCI} = \frac{L_{\max} - L}{L_{\max} - L_{\min}} \times 100 \quad (2)$$

where L is LST value of a given month.  $L_{\min}$  and  $L_{\max}$  denote the maximum and minimum LST values, respectively, for the month from multiyear time series. VHI are calculated based on VCI and TCI value using equation 3 [10]

$$\text{VHI} = (0.5 \times \text{VCI}) + (0.5 \times \text{TCI}) \quad (3)$$

Agricultural drought classified into five category shown in Table 2 [10].

**Table 2.** Agricultural drought category based on VHI

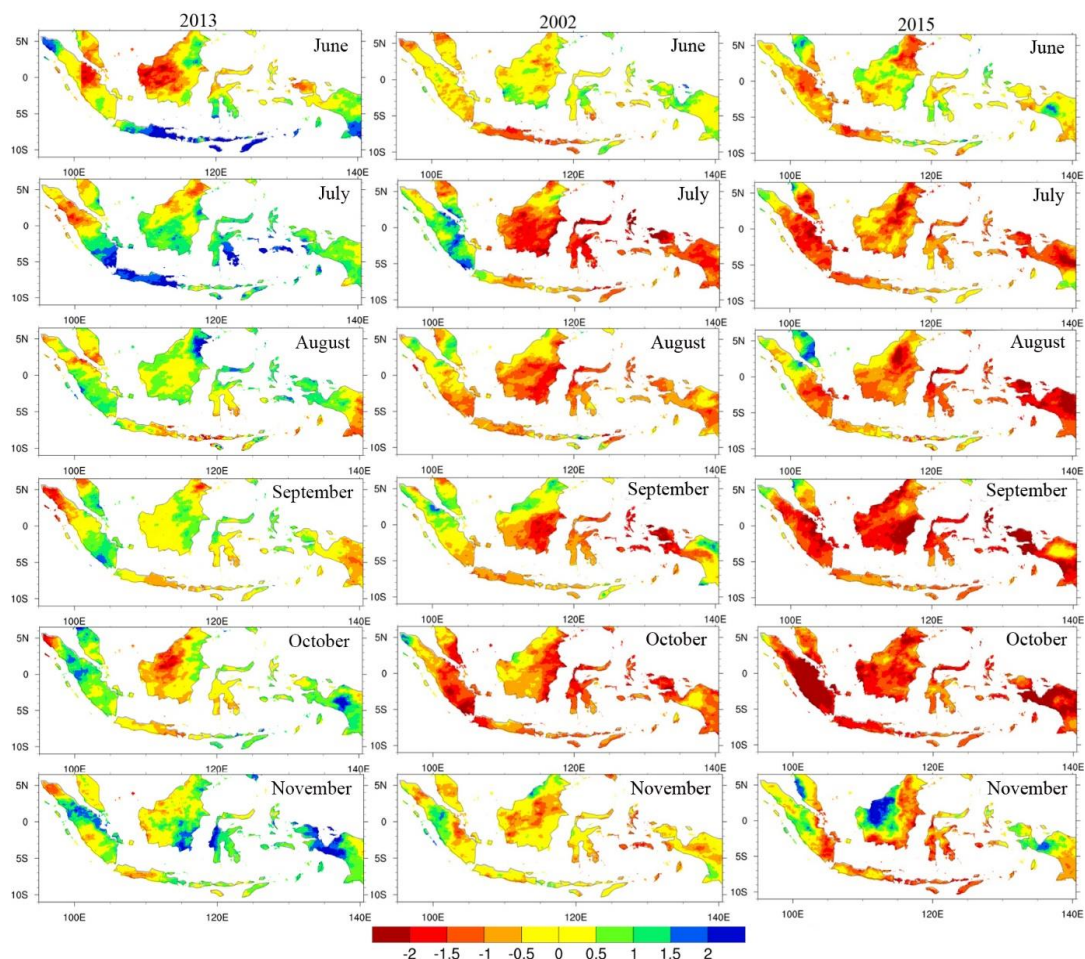
VHI	Category
< 10	Extreme
< 20	Severe
< 30	Moderate
< 40	Mild
>40	No Drought

### 2.3. Correlation analysis between SPI and VHI

Correlation analysis is conducted spatially between SPI and VHI over Indonesia region. This analysis is used to determine the time lag between meteorological drought and agricultural drought over the region.

## 3. Result and Discussion

### 3.1. Meteorological Drought



**Figure 1.** Spatial distribution of SPI in normal year(2013), weak El Niño (2002), and Strong El Niño (2015)

Meteorological drought is analyzed based on 1-month SPI (SPI-1). SPI-1 has significant correlation with rainfall so it is more appropriate to use it to analysis the meteorological drought [14]. The calculation of SPI-1 was done from 1981 to 2015. However, this study focuses on the drought

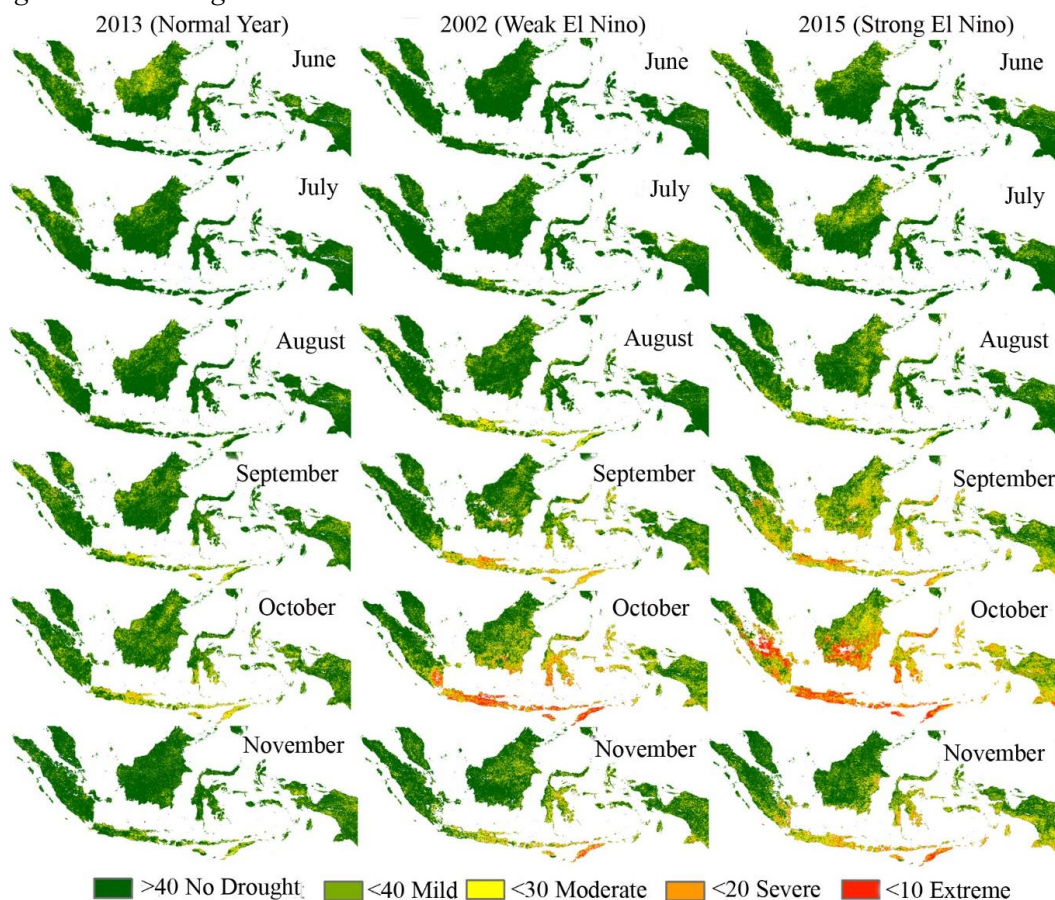
conditions that occurred in 2002, 2013 and 2015. SPI-1 in 2002, 2013 and 2015 was chosen to explain the spatial distribution of drought in Indonesia during weak El Niño (2002), normal (2013), and a strong El Niño (2015).

El Niño has significant impact on the occurrence of meteorological drought in Indonesia [2]. In a normal year (2013), most of regions in Indonesia are not affected by drought (see Figure 1). Drought occurs only in some parts of southern Indonesia (e.g. Java, Bali, Nusa Tenggara) and some part of northern Indonesia (e.g. north Sumatera and west Kalimantan). Southern region of Indonesia experienced drought from August to November, while the northern part of Indonesia such as Sumatera and Kalimantan experience drought from June to November. Moderate to extreme drought occurred in this year.

Drought during weak El Niño (2002) was not spread uniformly across Indonesia (see Figure 1). Southern part of Indonesia experienced drought earlier than other region that is started from June. In addition, Kalimantan, Sulawesi, and Papua experienced drought from July and some part of Sumatera experienced drought from August and reaches peak in October. In general, the drought is in the category of moderate to extreme drought with dominated by strong drought.

Drought in a strong El Niño (2015) occurred almost over the regions during July to October as seen in Figure 1. In June, drought has occurred only in Java and Sumatra, while in November drought occurred only in the southern part of Indonesia. The peak of the drought observed in October. In general, the drought is in the category of moderate to extreme drought.

### 3.2. Agricultural Drought



**Figure 2.** Spatial distribution of VHI in normal year (2013), weak El Niño (2002), and Strong El Niño (2015)

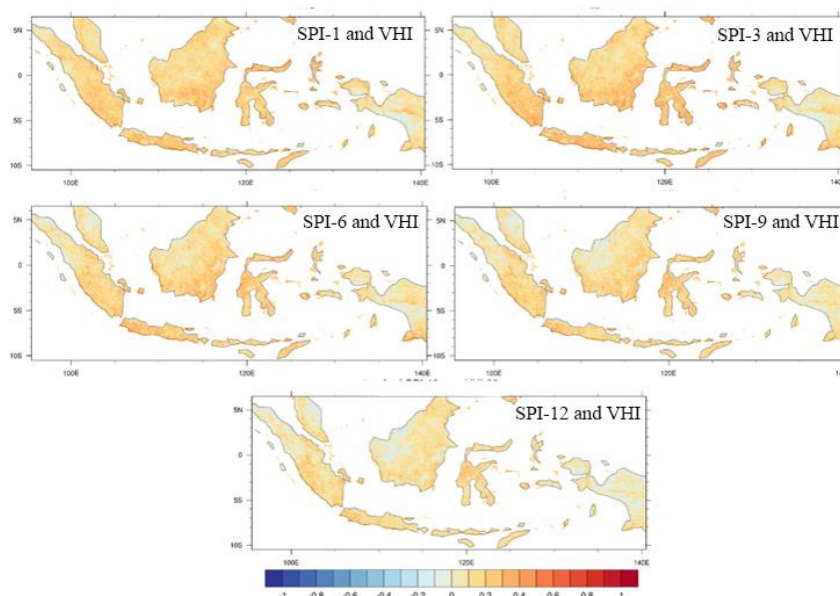


Decreasing levels of soil water or water stress can causes transpiration, photosynthesis, and retrieval of ions decrease that ultimately lead to impair growth and development of plants [15]. Reducing soil water content in constantly lead to agricultural drought [7]. This study used VHI to understand the duration, spatial distribution, and severity or category of agricultural drought. As well as SPI-1, VHI in 2002, 2013 and 2015 were chosen to determine the spatial distribution of agricultural drought in Indonesia during weak El Niño (2002), normal (2013), and strong El Niño (2015). Figures 2 show spatial distribution of VHI during normal year, weak El Niño, and Strong El Niño.

Agricultural drought in Indonesia is generally in the category of moderate to severe in the El Niño year. In 2013, drought observed from September to October in category mild to moderate especially in the southern part of Indonesia. While in 2002, the drought started from August to November where drought dominantly occured in the southern part of Indonesia such as Java, Bali, Nusa Tenggara, southern part of Kalimantan, and southern part of Sulawesi. In this year, drought occurred in the category of mild to extreme in which extreme droughts observed in October. Moreover, drought occurred from August to November in mild to extreme category during 2015 which extreme drought is dominant in October. In all cases (2002, 2013, and 2015) drought were observed in the southern part of Indonesia. Due to most of the region are dominated by crops. Wang and Liu (2014) pointed out that crops area more quickly respond to drought than the forest [9].

### 3.3. Relationship between Meteorological and Agricultural Drought

Meteorological drought and agricultural drought generally has lagged in time [9]. It is exist due to soil water content is not reduced directly when rainfall gradually decreased. Spatially SPI-1, SPI-3, SPI-6, SPI-9, and SPI-12 are positively correlated with VHI in most Indonesian regions as seen in Figure 3. The correlation coefficient has increased from SPI-1 to SPI-3, then tend to be declined following the increase in time scale. Percentage of coverage with significant correlation coefficient also increased from SPI-1 to SPI-3, then gradually decline following the increase in the time scale (Table 3).



**Figure 3.** Spatial correlation between SPI time scales and VHI

The coefficient correlation between the SPI-3 and VHI is 0.63 with significant at 99% level. The high coefficient correlation covers about 58% of Indonesian region. It implies that the total deficit of rainfall for three months giving a significant impact on agricultural drought or it can be speculated

that the lag time between meteorological and agriculture drought is about three months. Significant coefficient correlation values on SPI-3 and VHI are common in the southern part of Indonesia. This might be explained that the southern part of Indonesia are mostly center for crop production. This finding is agreed with previous studies (e.g. Wang et al, 2014; Huang et al, 2015). They pointed out crops respond to the drought after a rainfall deficit for three months [9,16].

**Table 3.** Data of correlation between SPI and VHI

Correlation	Maximum Correlation Value	Minimum Correlation Value	Percentage of coverage with significant correlation		
			Level 95%	Level 99%	Total
SPI-1 and VHI	0,55	-0,33	19%	38%	57%
SPI-3 and VHI	0,63	-0,32	15%	43%	58%
SPI-6 and VHI	0,58	-0,35	16%	31%	47%
SPI-9 and VHI	0,58	-0,47	15%	24%	39%
SPI-12 and VHI	0,61	-0,5	13%	15%	28%

**Table 4.** Coefficient correlation between SPI and VHI in some territory and containing note<sup>a</sup>

Territory	SPI-1 and VHI	SPI-3 and VHI	SPI-6 and VHI	SPI-9 and VHI	SPI-12 and VHI
Bali	0,34 <sup>**</sup>	0,37 <sup>**</sup>	0,32 <sup>**</sup>	0,25 <sup>**</sup>	0,21 <sup>**</sup>
Demak	0,14	0,19 <sup>*</sup>	0,16 <sup>*</sup>	0,13	0,13
Kediri	0,25 <sup>**</sup>	0,30 <sup>**</sup>	0,21 <sup>**</sup>	0,12	0,08
Cianjur	0,19 <sup>*</sup>	0,23 <sup>**</sup>	0,19 <sup>*</sup>	0,19 <sup>*</sup>	0,17 <sup>*</sup>
Progo	0,29 <sup>**</sup>	0,38 <sup>**</sup>	0,32 <sup>**</sup>	0,18 <sup>*</sup>	0,17 <sup>*</sup>
Lamongan	0,09	0,11	0,06	0,01	-0,01
Karawang	-0,01	0,05	0,07	0,05	0,09

<sup>a</sup>Note are referenced using alpha superscripts

<sup>\*(\*\*)</sup>correlation with significant level 95% (99%)

Mostly paddy fields in Indonesia is located in southern part and 40% centered on Java Island [17]. Territories of Java and Bali are chosen to get more detail about the correlation between SPI and VHI in paddy field. The correlation between SPI and VHI in paddy fields in most areas of Java and Bali (see Table 4) also show that the correlation increases in 3-months scale (SPI-3). Moreover, the coefficient correlation gradually decreased in SPI-6, SPI-9, and SPI-12. We noted that Karawang has a different pattern due to well-maintenance irrigation system. Again, Table 4 shows the higher coefficient correlation between SPI dan VHI is in 3-months scale (SPI-3). It shows that the agricultural drought, especially in crop land, occurs after rainfall deficit for three months. We speculated that in generally agricultural drought in the majority of land crop has 3 months behind the occurrence of meteorological drought.

#### 4. Conclusion

Meteorological drought and agricultural drought in Indonesia is more intensive during the El Niño years. The duration of meteorological drought is different in each region but generally observe during June to November. However, agricultural drought mostly occurs from August to November. In strong El Niño (2015), we found wider extent and higher severity of meteorological drought than in weak El Niño (2002). Areal extent and severity of meteorological drought generally is increased in each month and reached its peak in October. Agricultural drought in strong El Niño (2015) also has wider areal

extent and higher severity compared to weak El Niño (2002). The intensive agricultural drought mostly occurs in the areas with monsoonal precipitation such as Java, Bali, Nusa Tenggara, Lampung, southern part of Kalimantan, and southern part of Sulawesi. Moreover, VHI is significantly correlated with SPI-3 which accounted for 58% of the total area of Indonesia. It implies that the deficit of rainfall during three months has a significant impact on agricultural drought in Indonesia. In other words the agricultural drought is seen about 3-month behind the occurrence of meteorological drought. Therefore, SPI-3 and VHI clearly explain the relationship between meteorological drought and agricultural drought in Indonesia.

## References

- [1] Aldrian E and Susanto RD 2003 Identification of three dominant rainfall region within Indonesia and their relationship to sea surface temperature *International Journal of Climatology* **23** 1435-1452
- [2] Adianingsih ES 2014 Tinjauan Metode Deteksi Parameter Kekeringan Berbasis Data Penginderaan Jauh *Seminar Nasional Penginderaan Jauh* 210-220
- [3] BNPB 2015. Dampak EL-Niño Tahun 2015 terhadap Kekeringan di Indonesia. Retrieved from <http://www.bnpb.go.id/berita/2554/dampak-el-nino-tahun-2015-terhadap-kekeringan-di-indonesia> on May 27th, 2016
- [4] Utami AW Jamhari and Hardyastuti S 2011 El Niño, La Niña, dan penawaran pangan di Jawa, Indonesia *Jurnal Ekonomi Pembangunan* **12** 257-271
- [5] Zargar A, Sadiq R, Naser B, and Khan FI 2011 A review of drought indices *Environ Rev* **19** 333-349
- [6] Bhuiyan C, Singh RP, and Kogan FN 2006 Monitoring drought dynamics in the Aravalli region (India) using different indices based on ground and remote sensing data *Elsevier* **8** 289-302
- [7] Dalezios NR, Blanta A, Spyropoulos NV, and Tarquis AM 2014 Risk identification of agriculture drought for sustainable agroecosystems *Natural Hazards and Earth System Science* **14** 2435-2448
- [8] Roswiantiarti O, Sofan P, and Anggraini N 2011 Monitoring of drought-vulnerable area in Java Island, Indonesia using satellite remote-sensing data. *Jurnal Penginderaan Jauh* **8** 21-34.
- [9] Wang H, Lin H, and Liu D 2014 Remotely sensed drought index and its responses to meteorological drought in Southwest China *Remote Sensing Letters* **5** 413-422
- [10] Kogan FN 1995 Application of vegetation index and brightness temperature for drought detection *Adv Space Res* **15** 91-100
- [11] Funk C, Peterson P, Landsfeld M, Pedreros D, Verdin J, Shukla S, Husak G, Rowland J, Harrison L, Hoell A, and Michaelsen J 2015 The climate hazards infrared precipitation with stations-a new environmental record for monitoring extremes *Scientific Data* DOI: 10.1038/sdata.2015.66
- [12] McKee TB, N J Doesken, and J Kleist 1995 Drought monitoring with multiple time scales *American Meteorological Society* 233-236
- [13] WMO 2012 Standardized Precipitation Index User Guide *Geneva (SY): Publication Board World Meteorological Organization*
- [14] Adeogun BK, Nwude MO, Mohammad YS, and Adie DB 2014 Evaluation of suitable Standardized Precipitation Index time scales for meteorological, agricultural, and hydrological drought analyses *FUTA Journal of Research in Science* **2** 140-149
- [15] Arve LE, Torre S, Olsen JE, and Tanino KK 2011 Stomatal responses to drought stress and air humidity *Venkateswarlu*
- [16] Huang S, Huang Q, Chang J, Leng G, and Xing L 2015 The response of agricultural drought to meteorological drought and the influencing factors: A case study in the Wei River Basin, China *Agricultural Water Management* **159** 45-54
- [17] Ministry of Agriculture 2014 Statistics of Agricultural Land 2009-2013 *Jakarta (ID): Center for Agricultural Data and Information System-Ministry of Agriculture*