

Microscopic Controls on the Desorption/Dissolution of Sorbed U(VI) and Their Influence on Reactive Transport

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Research Objectives:

A sizable groundwater U plume exists in Hanford's 300 A resulting from the disposal of fuel rod dissolution wastes containing Al, Cu, and U to the vadose zone. This project is studying U-contaminated samples collected along a flow path from the waste source to the Columbia River. Three primary objectives are being pursued:

- To develop microscopic models for U desorption/adsorption in sediments along the flow path including both geochemical reaction and diffusive mass transport processes.
- To parameterize the microscopic models with appropriate laboratory measurements and data within context of a dual continuum, reactive transport model (DCM).
- To apply the parameterized DCM to laboratory columns of different size and sediment texture for testing of scaling hypotheses.

Research Progress and Implications:

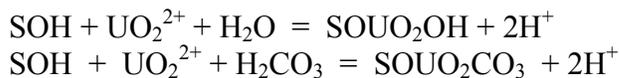
Molecular speciation measurements on sediments collected by excavation (Figure 1) using X-ray absorption (XAS – XANES and EXAFS) and cryogenic laser induced fluorescence spectroscopy (CLIFS) have shown that the chemical speciation of U(VI) changes from precipitated forms in the upper vadose zone near the historic source term to adsorption complexes in the deeper vadose zone and aquifer sediments (Bond et al., 2006; Catalano et al., 2006; Wang et al., 2005, Arai et al., 2006). Precipitated U(VI) exists within secondary mineral coatings on mm-sized lithic fragments in the form of coprecipitates with calcite, and as discrete uranyl phases including metatorbernite [Cu(UO₂PO₄)₂·8H₂O] and autunite. Adsorption occurs within aggregates of fine-grained phyllosilicates that include chlorite, vermiculite, and smectite, with variable ferrihydrite in low concentration. Adsorption is dominated by surface complexes to phyllosilicates in some locations and depths, and by surface complexes to ferrihydrite in others. Adsorption is significant within the fines fraction (e.g., log K_d ≈ 0.5 to 2.5 L/kg), but its overall effect is diluted strongly by mass-dominant gravel.

A critical aspect of U(VI) desorption/sorption in 300 Area sediments is that it is strongly mass transfer limited at the pore scale [Qafoku et al., (2005); and other papers in preparation]. Contaminant U is released from the sediments slowly, requiring long contact times and/or repeated pore volumes of leaching to accomplish significant U removal, regardless of speciation. The pore scale mass transfer process is best described with a distributed rate model. In preliminary studies comparing the <2mm fraction with whole sediments from the same location,

we have observed that adsorption reactions scale with texture (e.g., decrease in proportion to the mass content of fines), while mass transfer properties do not (e.g., they are slower than expected; Figure 2). The rate distribution of uranium mass transfer in 300 Area sediment is a function of reactive grain size distribution and is, consequently, scale-dependent.

In comparing different sediments from the 300 Area, mass transfer rates were observed to vary with U molecular speciation, sediment texture, and aqueous pH/carbonate content. For sediments containing adsorbed U(VI), the mass transfer rate decreases with increasing sorption strength and ferrihydrite content. Desorption is invariably slower than adsorption. These findings indicate that pore scale mass transfer rates will exhibit significant heterogeneity in the field depending on sediment properties and whether desorption or adsorption is occurring. Pore scale mass transfer is thought to result from diffusive transport within particle coatings and phyllosilicate aggregates of limited porosity.

In the deeper vadose zone and groundwater where adsorption is the primary retardation process, sorption strength is found to vary strongly with groundwater geochemical composition (e.g., pH and carbonate content) and mineral properties of the fines fraction, such as sediment surface area, ferrihydrite content, and phyllosilicate distribution (Bond et al., 2006). Accordingly, the distribution coefficient (K_d) varies by two orders of magnitude (Figure 3), over the range in water composition and sediment characteristics present at the site. A surface complexation model (SCM) has been developed for the 300 Area sediments (Bond et al., 2006) that explicitly includes aqueous speciation effects through surface complexation reactions:



Where SOH is an undifferentiated surface site with concentration of $3.84 \mu\text{mole}/\text{m}^2$. This model semi-quantitatively describes carbonate and pH effects on U(VI) adsorption and allows semi-quantitative prediction of K_d . However, different surface-area normalized, SCM model parameters are required for each sediment indicating the presence of uncharacterized variability in sediment mineralogy or site concentration.

Planned Activities:

- Research over the next two years of project performance will address the following issues:
- Mineralogic and physical causes for mass-transfer controlled desorption/sorption, and the specific sediment properties that control rates.
 - Characterization and measurement of microscopic transport parameters (e.g., diffusivity and tortuosity) and physical characteristics (e.g., porosity and thickness) of U-sorption domains.
 - Parameterization of the DCM and application to column adsorption/desorption experiments of different sizes with different sediment textures.

Information Access:

Publications

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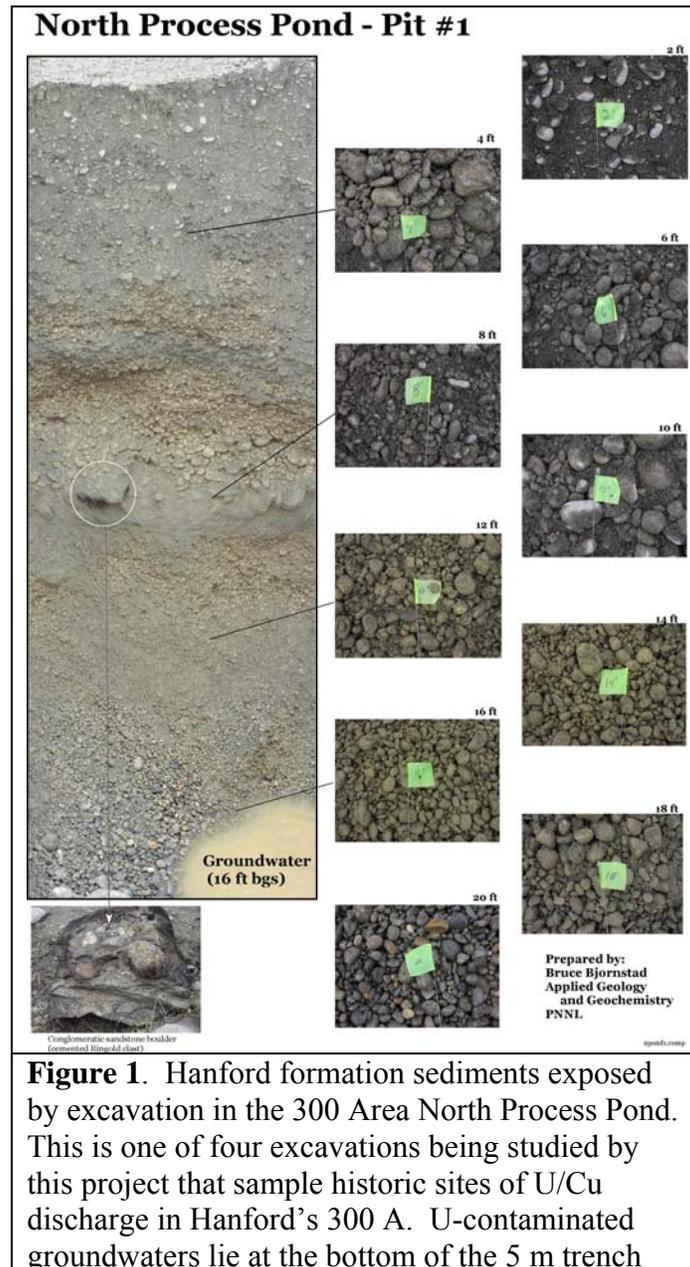
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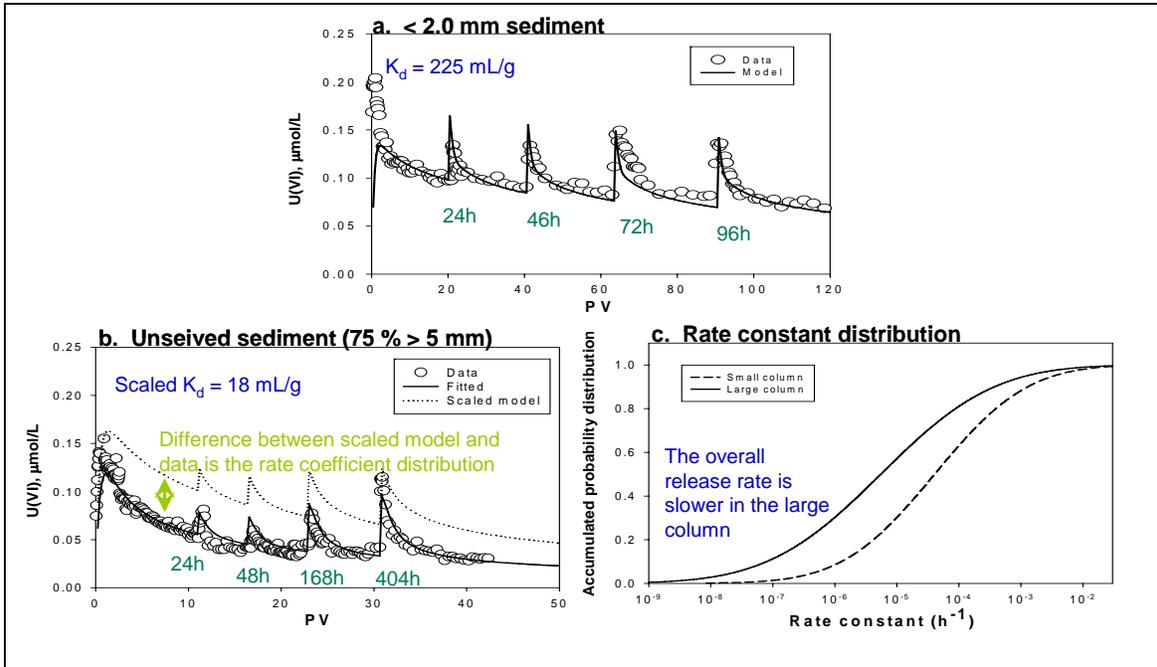


Figure 2. Desorption of contaminant U(VI) from sieved (< 2.00 mm) (a.) and unseived (b.) 300 Area sediment in columns as modeled with a distributed rate mass transfer model (DRM). Spikes in the effluent curves are stop-flow events applied to quantify mass-transfer rates. Estimated error in c.) increases to the left. The DRM is a simplified precursor to the more robust dual continuum model (DCM) that is the ultimate focus of research parameterization by this project.

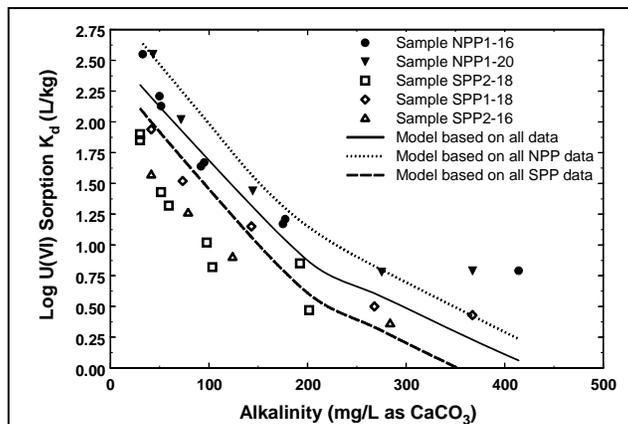


Figure 3. Alkalinity dependence of $\log K_d$ values for U(VI) sorption to 300 Area sediments. Solid curves show the fits to the data with surface complexation models calibrated with all of the data, or separately with NPP or SPP sediment data.